

Statistical method for the weighting function in the thermal radiation

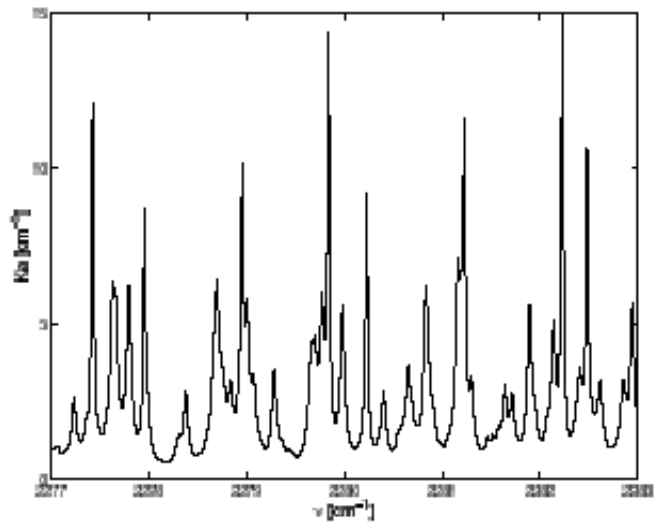
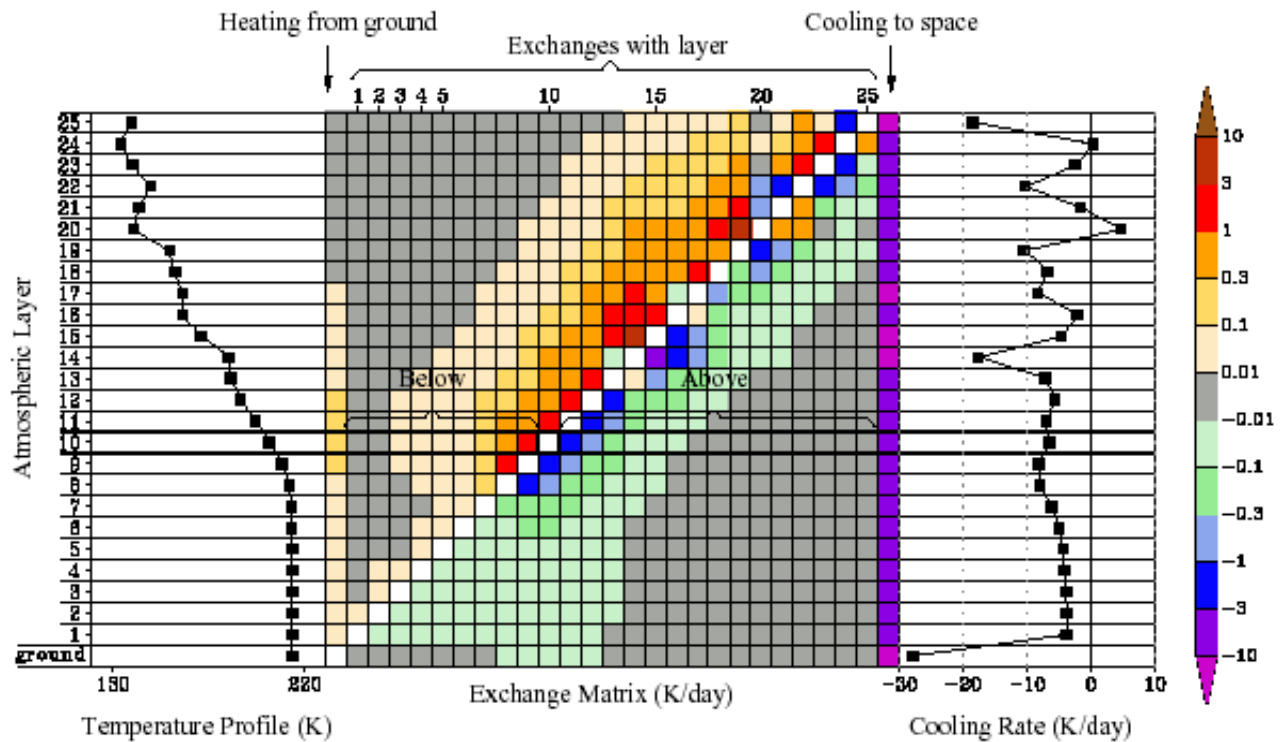
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Radiative transfer scheme

- δ -two stream approximation of radiative transfer equation for solar and thermal bands
- Adding method (solving linear system of equations)
- Economical computation (a good quality cost ratio)
- Separation of cloud and gaseous effects

Thermal radiative fluxes

- Improvements on the basis of Net Exchanged Rate (NER) formalism
- Better estimation of optical thickness of gaseous absorption for every layer in a simplified geometry and re-inject them into computation together with grey body effects



NER

- The atmosphere is divided in homogeneous layers
- Thermal exchange terms
 - ↪ Primary
 - Cooling to space (CTS)
 - Exchange with surface (EWS)
 - Exchange with adjacent layers (EAL)
 - ↪ Secondary
 - Exchange with other layers

NER

- The atmosphere is divided in homogeneous layers
- Thermal exchange terms
 - ↪ Primary - *as accurate as possible*
 - Cooling to space (CTS)
 - Exchange with surface (EWS)
 - Exchange with adjacent layers (EAL)
 - ↪ Secondary - *approximate treatment*
 - Exchange with other layers

Algorithm

- a calculation [I] with profile A and $\delta\tau_{\text{gas}}$ (CTS)
- a calculation [II] with profile B and $\delta\tau_{\text{gas}}$ (EWS)
- 3 calculations [III,IV,V] with profiles A,B,C and $\delta\tau_{\text{gas}}$ (EBL) = $\delta\tau_{\text{min}}$

Result is obtained:

$$[I] + [II] - [III] - [IV] + [V]$$

(multiplying by ΠB except [V] and [VIII])

Algorithm – vertical temperature profile

- Profile A:
 $\Pi B = 1$ at the ground and everywhere in the atmosphere => allows to suppress all other exchanges than 'cooling to space' (CTS)
- Profile B
 $\Pi B = 1$ at the ground and $\Pi B = 0$ everywhere in the atmosphere => allows to suppress all other exchanges than 'exchange with surface' (EWS)
- Profile C
The one corresponding to the physical truth => it mixes CTS, EWS with the 'exchanges between layers' (EBL)

Algorithm

- a calculation [I] with profile A and $\delta\tau_{\text{gas}}$ (CTS)
- a calculation [II] with profile B and $\delta\tau_{\text{gas}}$ (EWS)
- 3 calculations [III,IV,V] with profiles A,B,C and $\delta\tau_{\text{gas}}(\text{EBL})=\delta\tau_{\text{min}}$

- 3 calculations [VI,VII,VIII] with profiles A,B,C and $\delta\tau_{\text{gas}}(\text{EBL})=\delta\tau_{\text{max}}=\delta\tau_{\text{prox}}$

More accurate result is obtained:

$$[I] + [II] - \alpha \cdot ([III] + [IV] - [V]) - (1 - \alpha) \cdot ([VI] + [VII] - [VIII]) + \gamma$$

Method

$$[I] + [II] - \alpha ([III]+[IV]-[V]) - (1-\alpha) ([VI]+[VII]-[VIII]) + \gamma$$

How to calibrate α and γ ?

$\gamma = 0$ in the statistical model; this implies

$0 \leq \alpha \leq 1$ (0 corresponds to $\delta\tau_{\min}$; 1 to $\delta\tau_{\max}$)

Method

- vertical profiles over globe (ARPEGE)
- for gas only atmosphere (clear sky)
- computation of EBL fluxes* using $\delta\tau_{\min}$, $\delta\tau_{\max}$ and exact computation

* Two possibilities EAL accurate or inside statistical fitting

Method

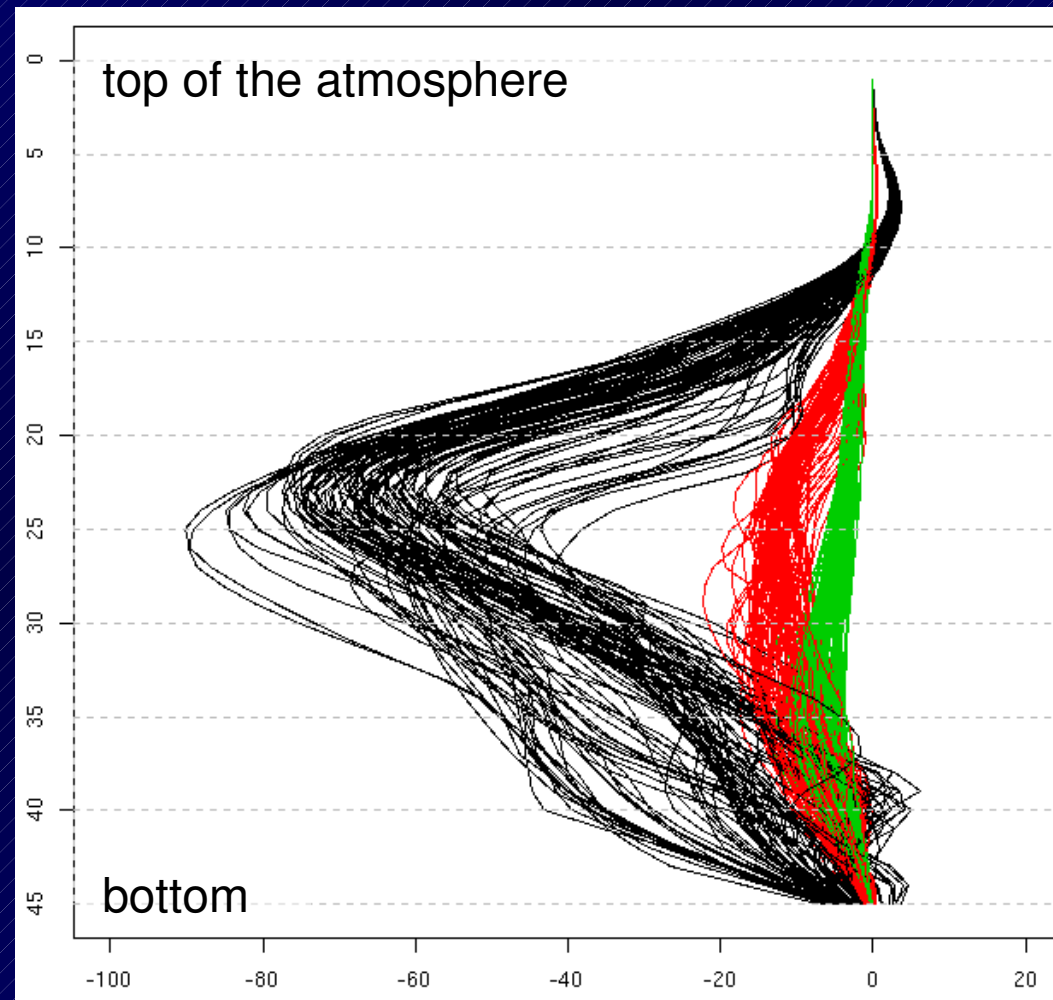
EBL fluxes using

$\delta\tau_{\min}$ - green

$\delta\tau_{\max}$ - black

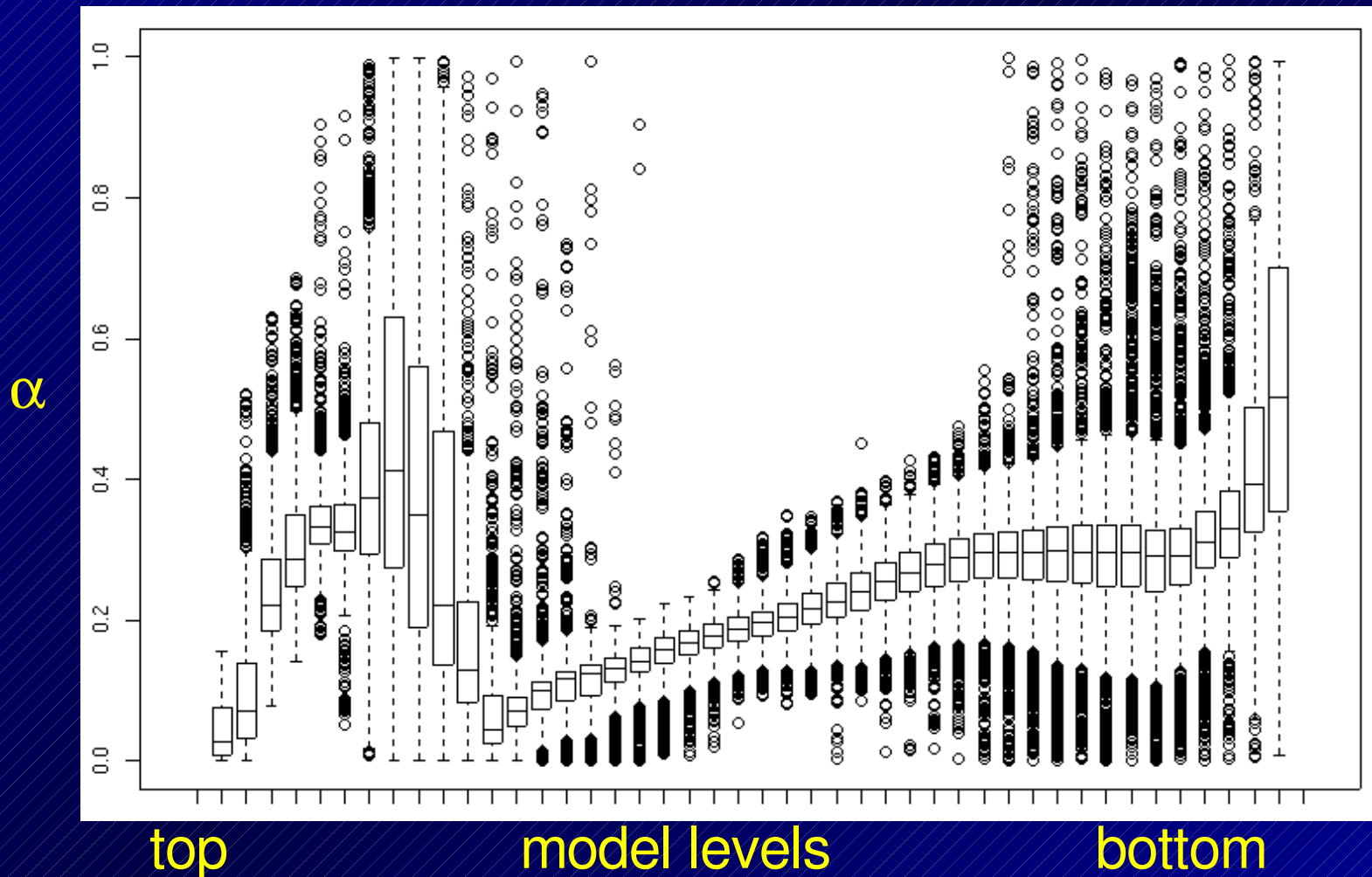
$\delta\tau$ - red

for some vertical
profiles



Method

Computation of α



Method

α increases:

in the lower atmosphere

in regions with temperature inversions

$$\alpha = N \left(\frac{p}{p_s} \right) + K \left(\frac{\partial c_p \Theta}{\partial \varphi} \right)$$

Method

1. version:

$$\alpha = 0.3 \frac{p}{p_s} + 0.1 \left(\frac{\partial c_p \Theta}{\partial \varphi} \right)$$

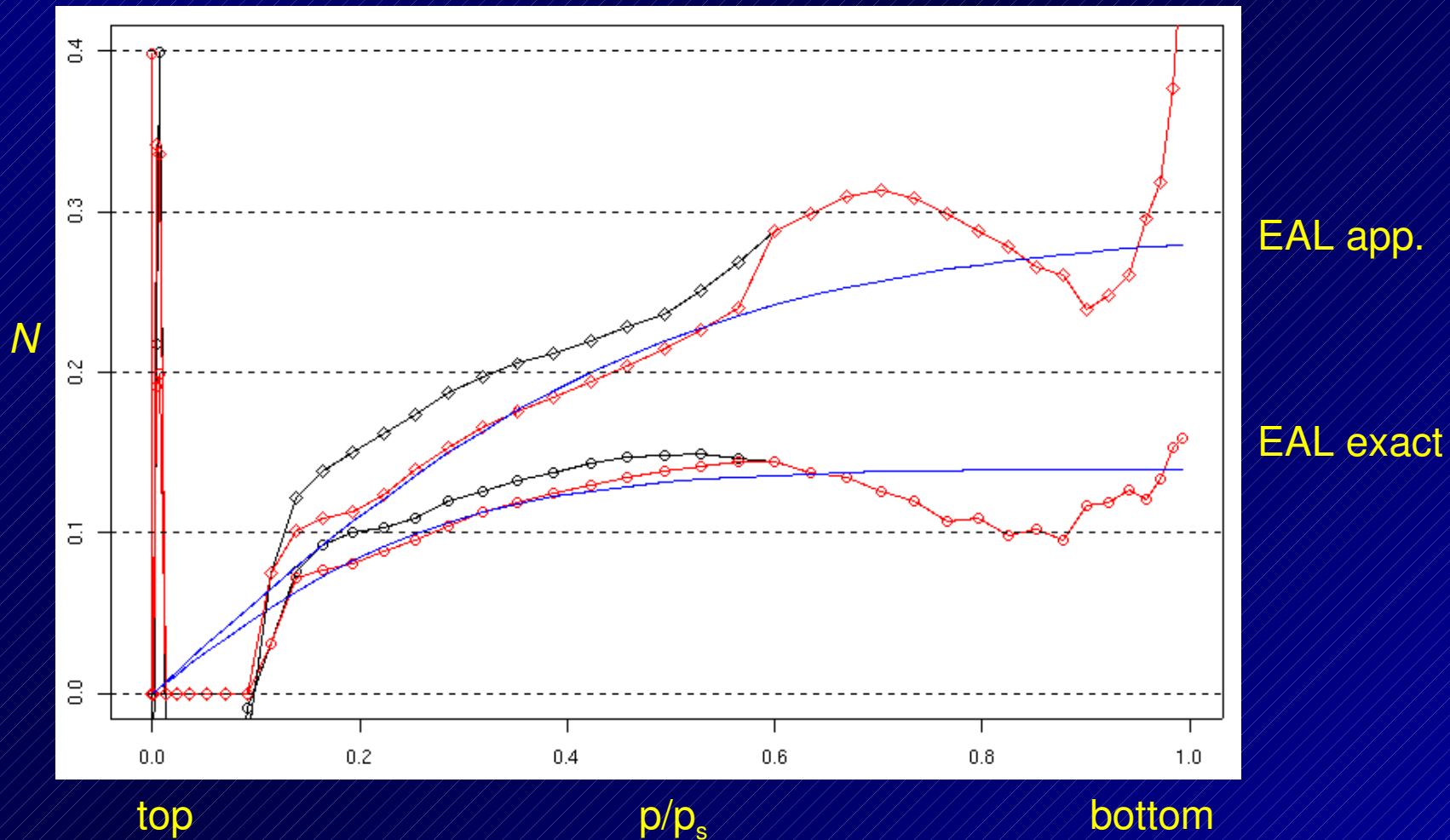
no vertical dependency
for stability part

2. version:

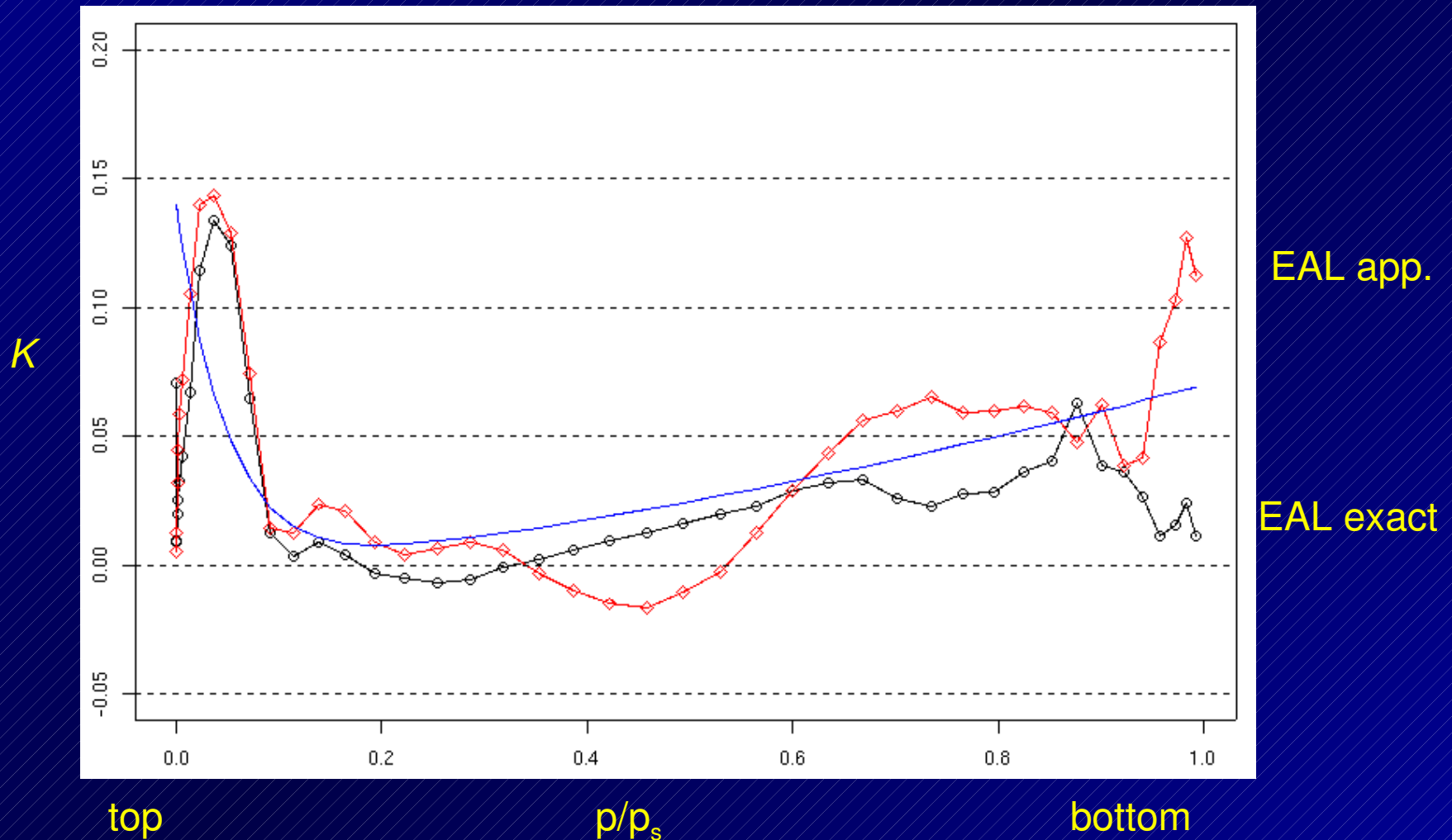
$$\alpha = N \left(\frac{p}{p_s} \right) + K \left(\frac{p}{p_s} \right) \frac{\partial c_p \Theta}{\partial \varphi}$$

allows different treatment of tropospheric and
stratospheric inversions

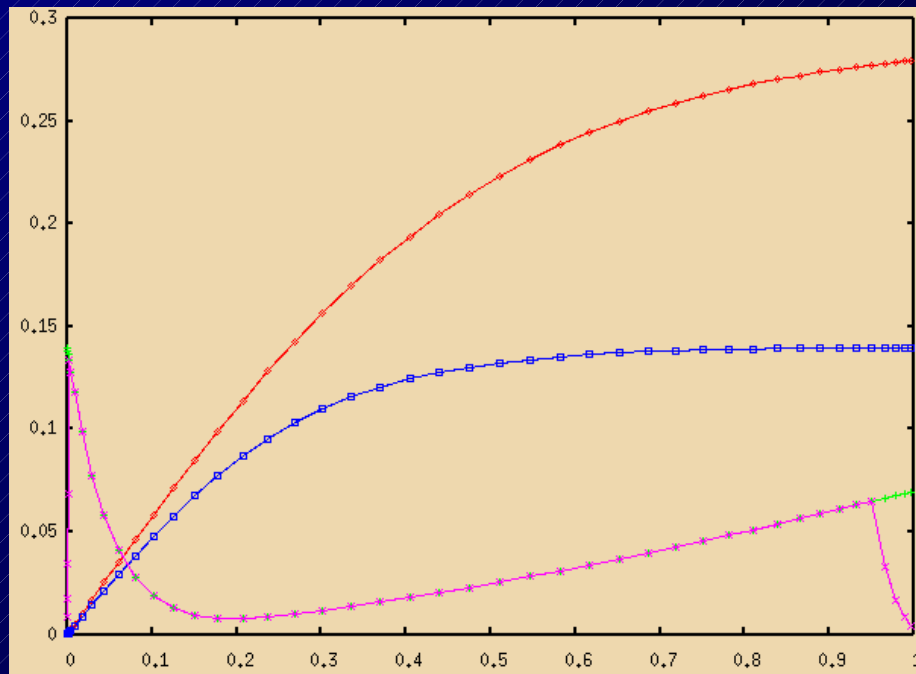
Fitting function for N



Fitting function for K



Fitting function for N and K



N

K

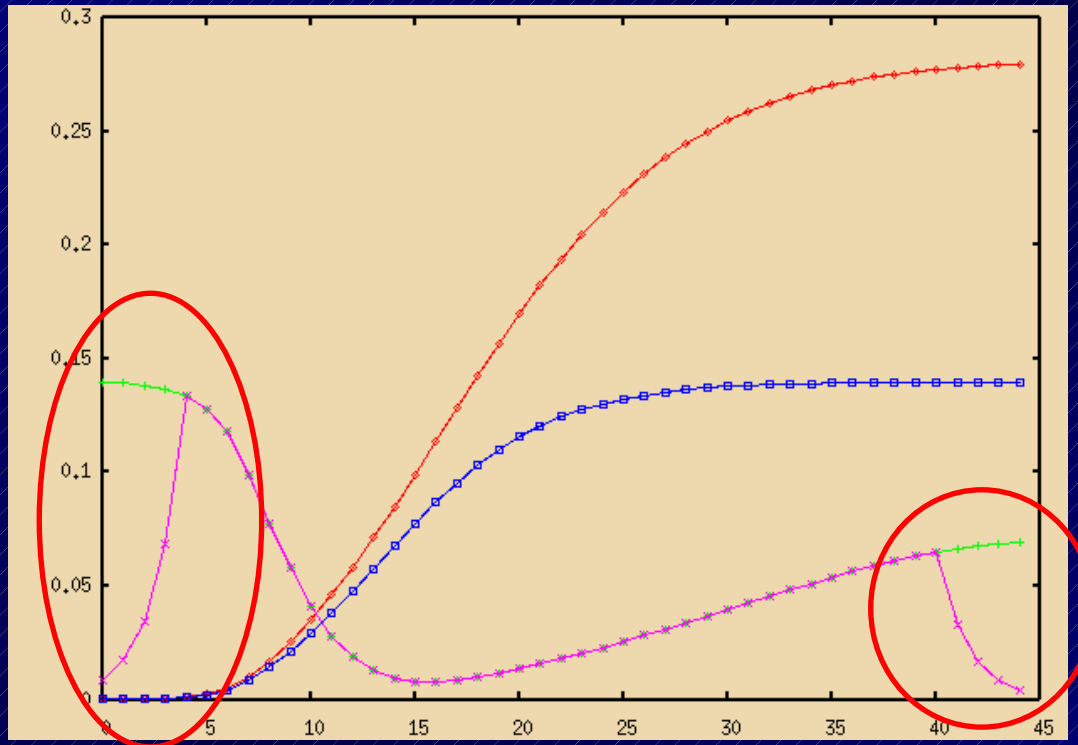
$$\alpha = 0.29 \tanh\left(2 \frac{p}{p_s}\right) + \left(0.07 \left(\frac{p}{p_s}\right)^{1.5} + 0.14 \left(1 - \frac{p}{p_s}\right)^{20}\right) \left(\frac{\partial c_p \Theta}{\partial \varphi}\right)$$

EAL app.

$$\alpha = 0.14 \tanh\left(3.5 \frac{p}{p_s}\right) + \left(0.07 \left(\frac{p}{p_s}\right)^{1.5} + 0.14 \left(1 - \frac{p}{p_s}\right)^{20}\right) \left(\frac{\partial c_p \Theta}{\partial \varphi}\right)$$

EAL exact

Fitting function for N and K

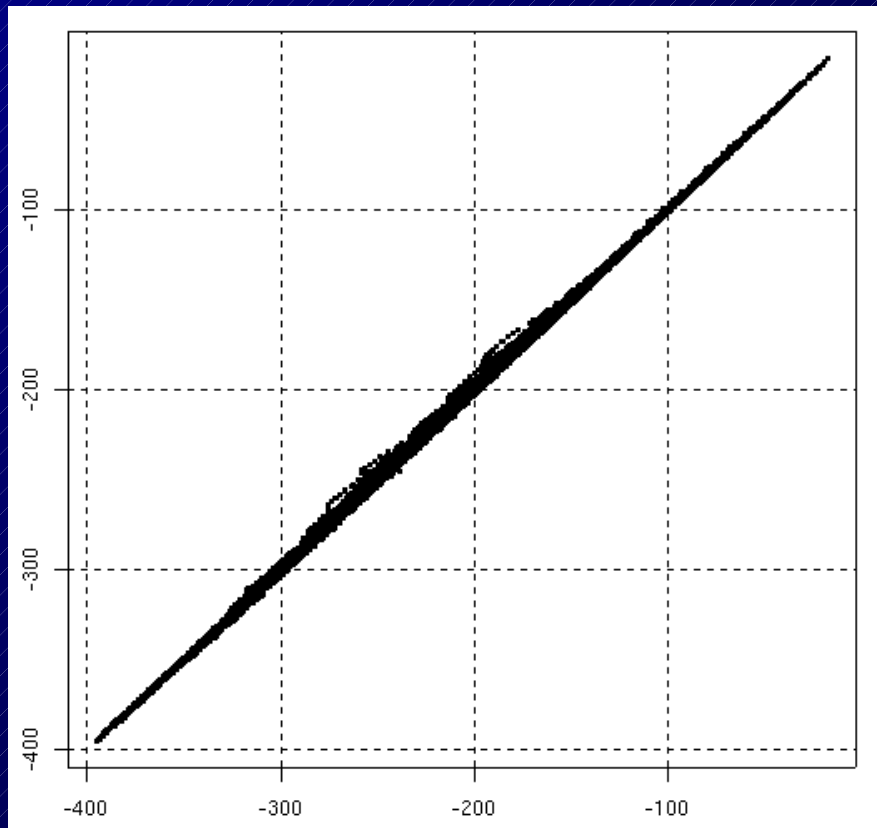


Pushed towards 0 at the edges to prevent double counting of local effect for $EAL=exact$

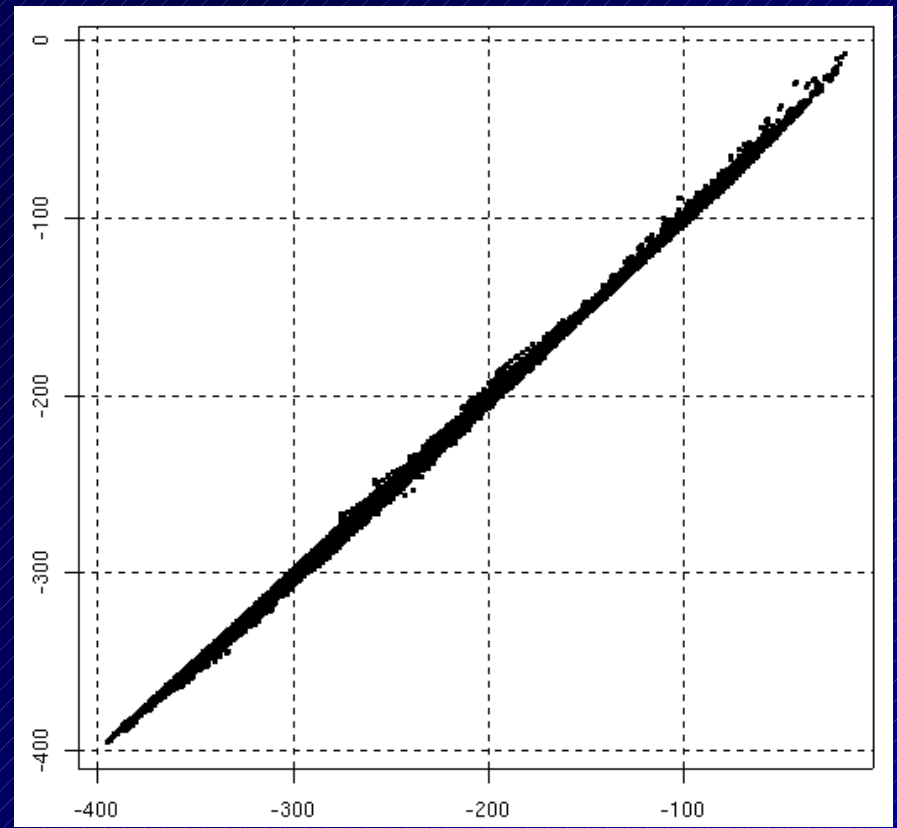
Results

parameterized versus computed thermal flux

Version 2



Version 1



Geopotential - RMSE

Comparison between statistical fit 1 and 2

Only in geopotential

Slight improvements in stratosphere

Only 4 days

