

*Regional Cooperation for
Limited Area Modeling in Central Europe*



Towards the 3D turbulence in the ALARO CMC

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- ▶ Motivation and objectives
- ▶ Introduction: addressing 3D turbulence effects
- ▶ The existing 3D approaches within the TOUCANS scheme
- ▶ Results:
 - ▶ Adjusting the 3D turbulence settings
 - ▶ The role of the deep convection scheme
- ▶ Conclusions and future steps

- ▶ The ultimate goal is to have a $\Delta x \sim < 1\text{km}$ model providing an added value to the existing operational configurations
- ▶ Therefore, we first addressed the settings of model dynamics (mostly horizontal diffusion) and concluded that its tuning is needed at these scales - not easy to achieve
- ▶ The 3D turbulence was considered a necessary addition to provide the added value but its role seems to change
- ▶ By the above increase in resolution, we move across the deep convection grey zone and enter the grey zone of shallow convection and turbulence (some adjustments are expected)

- ▶ There is an ongoing related activity in the ACCORD consortium (in the AROME CMC)
- ▶ Thereby, the following 3D effects are considered:
 - ▶ In production of turbulence energies after Goger et al. (2018, 2019)
 - ▶ The contribution of largest sub-filter scale eddies after Moeng et al. (2010); so-called Leonard terms (originates from LES)

$$\overline{w'\phi'} = -K_\phi \frac{\partial \bar{\phi}}{\partial z} + 2 \left(\frac{\Delta_f^2}{12} \right) \left(\frac{\partial \bar{w}}{\partial x} \frac{\partial \bar{\phi}}{\partial x} + \frac{\partial \bar{w}}{\partial y} \frac{\partial \bar{\phi}}{\partial y} \right) \quad (1)$$

- ▶ Additionally, we can utilize the 1D+2D scheme in TOUCANS, solving 2D diffusion equation for the horizontal direction

- TOUCANS is a two prognostic energy scheme (Baštak Ďuran et al. 2014, 2018):

$$\frac{\partial e_k}{\partial t} + ADV(e_k) = \frac{\partial}{\partial z} \left(K_{e_k} \frac{\partial e_k}{\partial z} \right) + I + II - \frac{2e_k}{\tau_k} \quad (2)$$

$$\frac{\partial e_t}{\partial t} + ADV(e_t) = \frac{\partial}{\partial z} \left(K_{e_t} \frac{\partial e_t}{\partial z} \right) + I - \frac{2e_t}{\tau_t} \quad (3)$$

$$I = -\overline{u'w'} \frac{\partial \bar{u}}{\partial z} - \overline{v'w'} \frac{\partial \bar{v}}{\partial z}, \quad II = E_{s_L} \overline{s_L'w'} + E_{q_t} \overline{q_t'w'} \quad (4)$$

$$s_L = c_{pd} \left(1 + \left[\frac{c_{pv}}{c_{pd}} - 1 \right] q_t \right) T + gz - (L_v q_l + L_s q_i), \quad q_t = q_t + q_l + q_i \quad (5)$$

I, II - source terms; s_L - static energy; q_t - total specific moisture

- In Goger et al. (2018, 2019), the TKE/TTE shear production term (I) is extended into 3D:

$$I_{3D} = \underbrace{-\overline{u'w'}\frac{\partial \bar{u}}{\partial z} - \overline{v'w'}\frac{\partial \bar{v}}{\partial z}}_{I_{\text{vert}}} - \underbrace{\overline{u'u'}\frac{\partial \bar{u}}{\partial x} - \overline{u'v'}\frac{\partial \bar{u}}{\partial y} - \overline{u'v'}\frac{\partial \bar{v}}{\partial x} - \overline{v'v'}\frac{\partial \bar{v}}{\partial y}}_{I_{\text{horiz}}} \quad (6)$$

- The I_{horiz} term is parameterized following Smagorinsky (1963):

$$I_{\text{horiz}} = \left(L_K^H\right)^2 \cdot \left[\left(\frac{\partial u}{\partial x}\right)^2 + \left(\frac{\partial v}{\partial y}\right)^2 + \frac{1}{2} \left(\frac{\partial u}{\partial y} + \frac{\partial v}{\partial x}\right)^2 \right]^{\frac{3}{2}} \quad (7)$$

- The difference between Goger et al. (2018) and Goger et al. (2019) is in L_K^H treatment:

$$L_{K-G18}^H = c_s \Delta x, \quad L_{K-G19}^H = f(U, \sigma_{u,v}, z_i) \quad (8)$$

c_s - Smagorinsky's constant; U - velocity scale; $\sigma_{u,v}$ - horizontal wind variances; z_i - PBL height

- The 1D+2D scheme Eqs. are derived assuming $\partial K_{M/H,hor}/\partial x + \partial K_{M/H,hor}/\partial y = 0$:

$$\frac{\partial u_i}{\partial t} + \dots = -K_{M,hor} \frac{\partial^2 u_i}{\partial x^2} - K_{M,hor} \frac{\partial^2 u_i}{\partial y^2} + \frac{\partial}{\partial z} (\overline{w' u_i'}) \quad (9)$$

$$\frac{\partial \Psi}{\partial t} + \dots = -K_{H,hor} \frac{\partial^2 \Psi}{\partial x^2} - K_{H,hor} \frac{\partial^2 \Psi}{\partial y^2} + \frac{\partial}{\partial z} (\overline{w' \Psi'}) \quad (10)$$

$$\overline{w' u_i'} = -K_{M,ver} \frac{\partial u_i}{\partial z}, \quad \overline{w' \Psi'} = -K_{H,ver} \frac{\partial \Psi}{\partial z} + \text{TOMs} \quad (11)$$

- The exchange coefficients for horizontal and vertical directions have the same "shape":

$$K_{M,ver} = L_K C_K \sqrt{e_k} \chi_3(Ri), \quad K_{H,ver} = L_K C_K C_3 \sqrt{e_k} \phi_3(Ri) \quad (12)$$

$$K_{M,hor} = L_K^H C_K^H \sqrt{e_k} \chi_{3,hor}(Ri), \quad K_{H,hor} = L_K^H C_K^H C_3 \sqrt{e_k} \phi_{3,hor}(Ri) \quad (13)$$

- The shear and stretching HTLS are computed after Wang et al. (2021):

$$L_{\text{Hshr}} = sW \left[\left(\frac{\partial v}{\partial x} \right)^2 + \left(\frac{\partial u}{\partial y} \right)^2 \right]^{-\frac{1}{2}} \quad (14)$$

$$L_{\text{Hstr}} = sW \left[\left(\frac{\partial u}{\partial x} \right)^2 + \left(\frac{\partial v}{\partial y} \right)^2 \right]^{-\frac{1}{2}}, \quad s = \left(\frac{\Delta_0}{\Delta} \right)^\alpha \quad (15)$$

- The combined HTLS is obtained by averaging L_{Hshr} and L_{Hstr} and protected as follows:

$$L_K^{\text{H}'} = \sqrt{L_{\text{Hshr}} L_{\text{Hstr}}}, \quad L_K^{\text{H}} = \max \left(L_K^{\text{H}'}, c_s \sqrt{\Delta x \Delta y} \right) \quad (16)$$

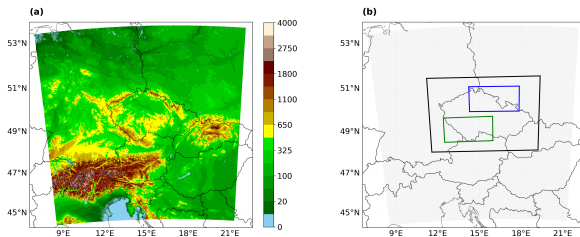
s - resolution-dependent scaling factor; $W = \sqrt{u^2 + v^2}$ - wind speed

- ▶ Following the preliminary tests (small impact of 3D turbulence), we "adjusted" the 1D + 2D scheme's code:

$$K_{M,hor} = L_K^H C_K' C_K \sqrt{e_k} \chi_{3,hor}(Ri), \quad K_{H,hor} = L_K^H C_K' C_K C_3 \sqrt{e_k} \phi_{3,hor}(Ri) \quad (17)$$

Results: The preliminary analysis across resolutions

- ▶ Central European domain: 1024 x 1024 km, linear grid, 87 vert. lev., $\Delta x=4, 2, 1$ and 0.5 km
- ▶ NHYD dyn., ICI scheme (1 iter.), horiz. diffusion and ALARO-1 phy. (both as in operational ALADIN-CZ)
- ▶ INIT: ALADIN-CZ, LBC: ARPEGE (3-h freq.)

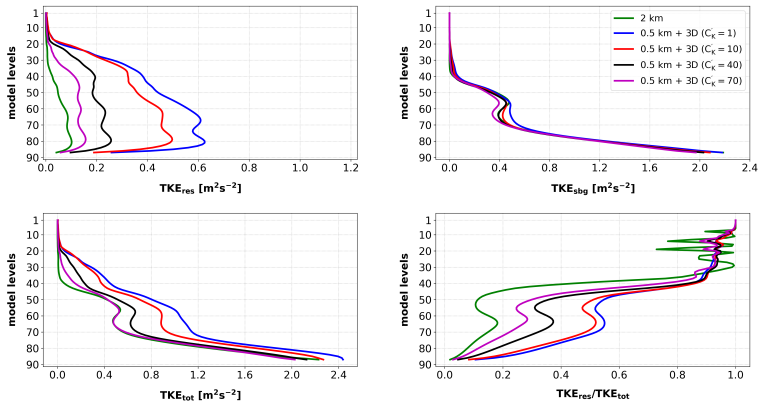


- ▶ The results from two convection cases: (i) with considerable large-scale forcing and (ii) more locally driven
- ▶ During the analysis, we rely on: KE spectra, averaged profiles of resolved and subgrid TKE (TKE_{res} and TKE_{subg}), convective fluxes and noise-indicative fields (e.g., pressure departure)
- ▶ The TKE_{res} is computed by applying the Reynolds averaging to predicted wind components ($u_i = \bar{u}_i + u'_i$):

$$TKE_{res} = \frac{1}{2} \sum_{i=1}^3 u_i'^2; \quad \bar{u}_{i,MA} = \frac{1}{N} \sum_{j=n-k}^{n+k} u_{i,j}, \quad N = 2k + 1 \quad (18)$$

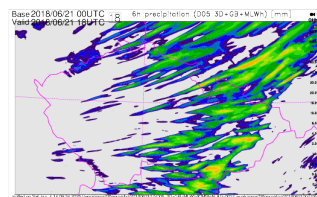
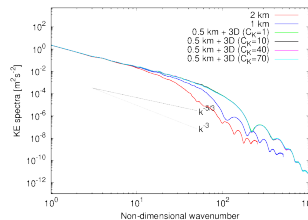
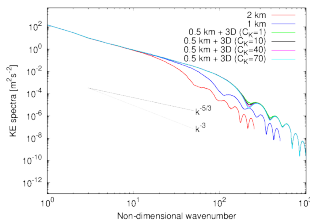
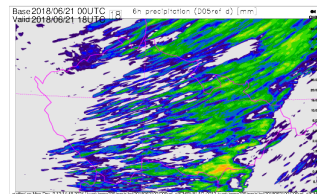
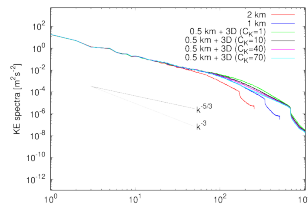
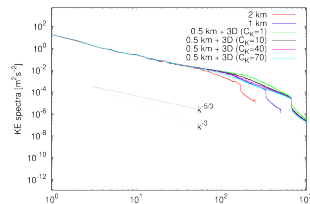
Results: Adjusting the 3D turbulence settings (case 1)

- The sensitivity tests to the C'_K parameter (the strength of the 1D+2D scheme):



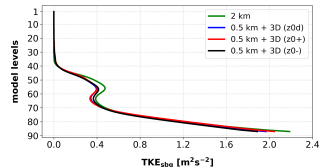
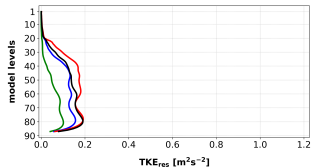
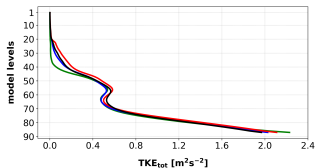
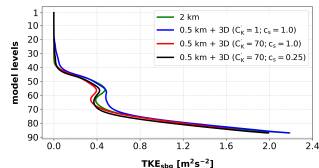
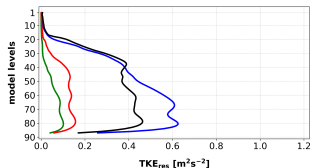
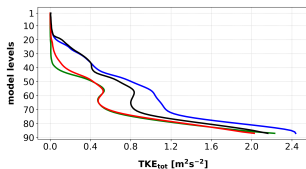
Results: Adjusting the 3D turbulence settings (case 1)

- The impact off the 3D turbulence at $\Delta x = 0.5$ km:

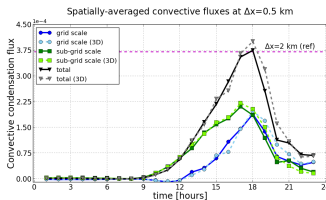
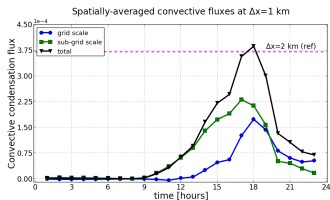
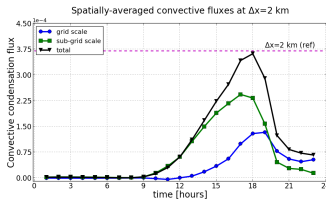
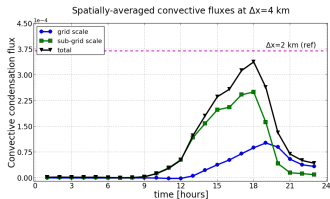


Results: Adjusting the 3D turbulence settings (case 1)

- The sensitivity of complementary diagnostics to L_K^H protection and z_{0m}^{oro} :



► How 3MT scheme adapts with a decrease in Δx ?

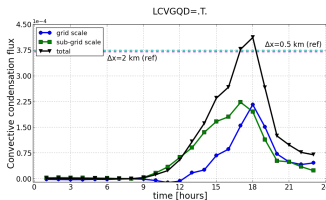
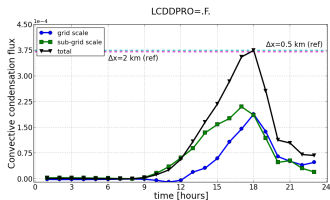
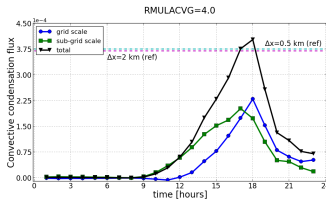
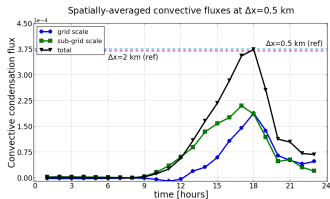


condensation fluxes
in updraught

1. 3MT scheme adapts only
slightly with Δx

2. Total condensation flux
increases with the
grid-scale part

► Can we help the 3MT scheme to adapt a bit more?

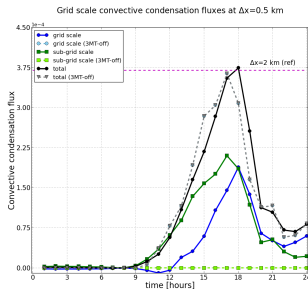
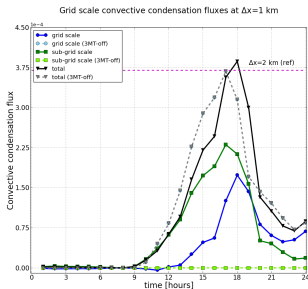


condensation fluxes
in updraught

1. The largest sensitivity
is to RMULACVG↓

2. The downdraught part
has very small impact

- The impact off 3MT scheme at $\Delta x = 1.0$ km and $\Delta x = 0.5$ km:

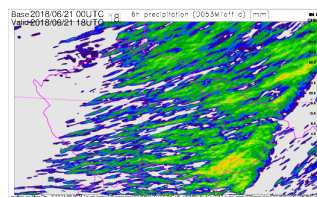
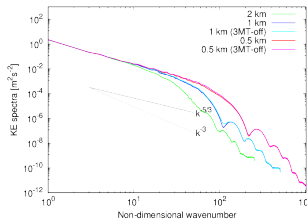
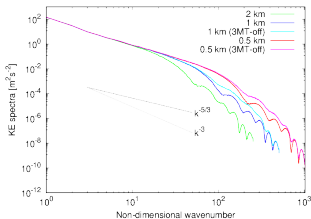
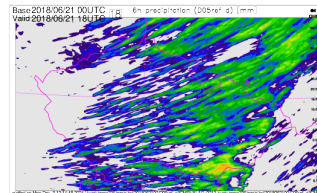
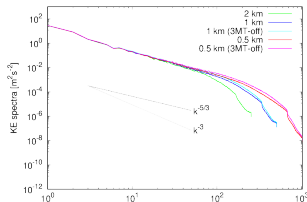
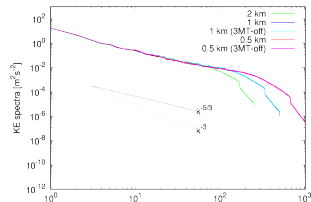


condensation fluxes
in updraught

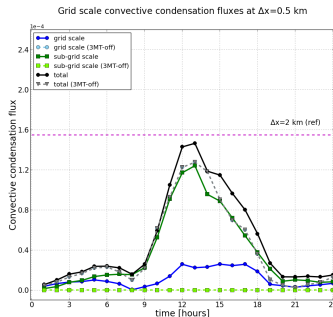
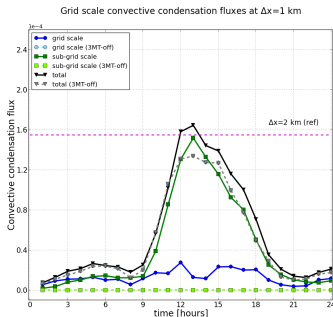
1. The maximum is shifted to one hour earlier
2. The total magnitude is comparable

Can we simply switch off the 3MT scheme?

- The impact off 3MT scheme at $\Delta x = 1.0$ km and $\Delta x = 0.5$ km:



- The impact off 3MT scheme at $\Delta x = 1.0$ km and $\Delta x = 0.5$ km:

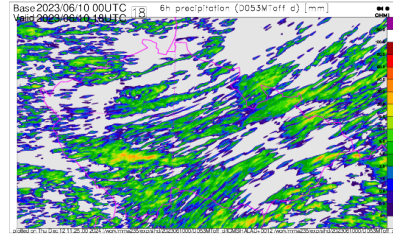
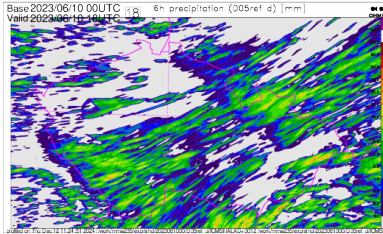


condensation fluxes
in updraught

1. The contribution from sub-grid- and grid-scale different than in case 1
2. The total flux is $\sim 20\%$ smaller without the 3MT

Results: The role of the deep convection scheme (3MT) - case 2

- The impact off 3MT scheme at $\Delta x = 0.5$ km:



- ▶ The default model settings at $\Delta x = 1.0$ km and $\Delta x = 0.5$ produce progressively increasing noisy patterns
- ▶ Stronger horizontal diffusion is needed but the existing components (SLHD + 2 spectral) seem hard to tune
- ▶ The 3D turbulence helps to reduce the noise and make precipitation forecast closer to observations; further inspection is needed on the values of closure parameters and L_K^H
- ▶ The 3MT scheme should not be switched off for $\Delta x \sim < 1.0$ km but its adjustment/tuning is required
- ▶ More extensive validation is needed

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- ▶ Smagorinsky, J. (1963): General circulation experiments with the primitive equations, *Mon. Weather Rev.*, 91(3): 99–164.
- ▶ Wang, W., Liu, B., Zhu, L., Zhang, Z., Mehra, A., and Tallapragada, V. (2021): A New Horizontal Mixing-Length Formulation for Numerical Simulations of Tropical Cyclones., *Wea. Forecasting*, 36(2), 679-695.

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Thank you for your attention!



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