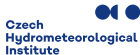


*Regional Cooperation for
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TKE-based turbulence length scale in TOUCANS: the final settings and their preliminary validation

Mario Hrastinski, Ján Mašek, Ivan Bašták Ďurán, Branko Grisogono and Radmila Brožková



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- ▶ Introduction: an overview of the TOUCANS scheme
- ▶ Development of the new Turbulence Length Scale (TLS) formulation
- ▶ Results:
 - ▶ 1D model validation against LES data
 - ▶ The high-impact weather situations in 3D model
 - ▶ The objective verification
- ▶ Conclusion

- TOUCANS is a two prognostic energy scheme (Bařtak Ďuran et al. 2014, 2018):

$$\frac{de_k}{dt} = -g \frac{\partial}{\partial p} \left(\rho K_{e_k} \frac{\partial e_k}{\partial z} \right) + I + II - \frac{e_k^{\frac{3}{2}}}{\tau_k} \quad (1)$$

$$\frac{de_t}{dt} = -g \frac{\partial}{\partial p} \left(\rho K_{e_t} \frac{\partial e_t}{\partial z} \right) + I - \frac{e_t^{\frac{3}{2}}}{\tau_t} \quad (2)$$

$$I = -\overline{u'w'} \frac{\partial \bar{u}}{\partial z} - \overline{v'w'} \frac{\partial \bar{v}}{\partial z}, \quad II = E_{s_{sL}} \overline{s_{sL}'w'} + E_{q_t} \overline{q_t'w'} \quad (3)$$

$$s_{sL} = c_{pd} \left(1 + \left[\frac{c_{pv}}{c_{pd}} - 1 \right] q_t \right) T + gz - (L_v q_l + L_s q_i), \quad q_t = q_t + q_l + q_i \quad (4)$$

- Computation of turbulent fluxes above the surface layer:

$$\overline{u'w'} = -K_M \frac{\partial u}{\partial z}, \quad \overline{v'w'} = -K_M \frac{\partial v}{\partial z} \quad (5)$$

$$\overline{s'_L w'} = -K_H \frac{\partial s_L}{\partial z} + TOMs, \quad \overline{q'_t w'} = -K_H \frac{\partial q_t}{\partial z} + TOMs \quad (6)$$

$$K_M = C_K L_K \chi_3 \sqrt{e_k}, \quad K_H = C_K L_H \phi_3 \sqrt{e_k}, \quad L_H = C_3 L_K \quad (7)$$

- Computation of turbulent fluxes in the surface layer:

$$\overline{(w'\phi')} = C_\phi \sqrt{(u^2 + v^2)} [\phi(z) - \phi_s], \quad C_{M/H} = C_{M/HN} F_{M/H}(Ri) \quad (8)$$

$$C_{MN} = \left[\frac{\kappa}{\ln \left(1 + \frac{z}{z_{0m}} \right)} \right]^2, \quad C_{HN} = \left[\frac{\kappa^2}{\ln \left(1 + \frac{z}{z_{0m}} \right) \ln \left(1 + \frac{z}{z_{0h}} \right)} \right] \quad (9)$$

- ▶ TLS is an essential quantity in the TKE – L type of closure representing the size (dimension) of the most energetic turbulence eddies
- ▶ TOUCANS distinguishes several TLS: $L_{K/H}$ and L_ϵ (related via the master TLS - L_n)
- ▶ Following [Redelsperger et al. \(2001\)](#), the relationship between L_K , L_ϵ and L_n is stability-dependent (cf. [Mašek et. al. \(2022\)](#) for details):

$$L_K = L_n F_\epsilon^{\frac{1}{3}}, \quad L_\epsilon = \frac{L_n}{F_\epsilon}, \quad F_\epsilon = \left[\frac{1 - Ri_f}{\chi_3(Ri_f)} \right]^{\frac{3}{4}} \quad (10)$$

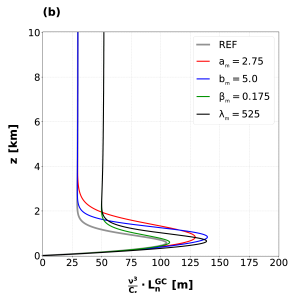
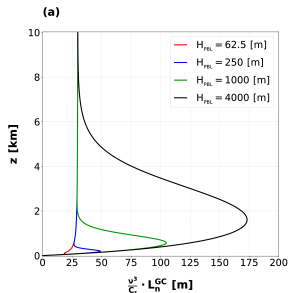
- ▶ For consistency with previous pTKE scheme, it is assumed:

$$l_m = \frac{\nu^3}{C_\epsilon} (L_K^3 \cdot L_\epsilon)^{\frac{1}{4}} \longrightarrow L_n = \frac{C_\epsilon}{\nu^3} l_m, \quad \nu = (C_K C_\epsilon)^{\frac{1}{4}} \quad (11)$$

l_m - Prandtl type mixing length; $C_\epsilon/\nu^3 \approx 6$ (prone to tuning)

- Currently, the Geleyn-Cedilnik formulation is a default choice in TOUCANS:

$$L_n^{GC} = \frac{C_\epsilon}{\nu^3} \frac{\kappa Z}{1 + \frac{\kappa Z}{\lambda_m} \left[\frac{1 + \exp(-a_m \sqrt{\frac{z}{H_{PBL}} + b_m})}{\beta_m + \exp(-a_m \sqrt{\frac{z}{H_{PBL}} + b_m})} \right]} \quad (12)$$



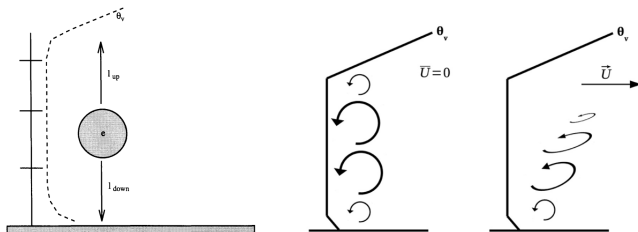
1. Very sensitive to the H_{PBL} estimation
2. $UAL=const.$ (should be regime-dependent)
3. Small variability (in practice)
4. Too strong mixing in SABL

REF: $a_m=4.5$, $b_m=3.0$, $\beta_m=0.1$
and $\lambda_m=300$

- We start from the generalized version of BL89 TLS following [Rodier et al. \(2017\)](#):

$$\int_z^{z+L_{up}} \left\{ \frac{g}{\theta_v(z')} [\theta_v(z') - \theta_v(z)] + c_0 \sqrt{e(z')} S(z') \right\} dz' = e(z) \quad (13)$$

$$\int_{z-L_{down}}^z \left\{ \frac{g}{\theta_v(z')} [\theta_v(z) - \theta_v(z')] + c_0 \sqrt{e(z')} S(z') \right\} dz' = e(z) \quad (14)$$



$$L_{AVG} = \left(\frac{L_{up}^p + L_{down}^p}{2} \right)^{\frac{1}{p}} \quad (15)$$

► Can we assign L_{AVG} directly to any of TLS options within TOUCANS?

1. Obey similarity laws in the surface layer: ~~κ~~ and ~~κ_e~~
2. Ensure numerically stable solution: ~~κ_m~~

► Initial attempt with the remaining option (L_n):

$$L_n = \min \left(\frac{C_\epsilon}{\nu^3} \kappa Z, L_{AVG} \right) \quad (16)$$

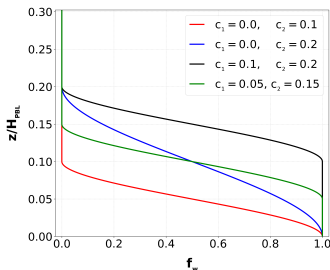
Problems:

1. Obeying MOST is not ensured
2. Possibility of a "jumpy" solution
3. Insufficient mixing (overall)

- Smoothing the transition between two solutions in the surface layer:

$$f_w = 3 \cdot f_w'^2 - 2 \cdot f_w'^3; \quad f_w' = \max \left[0, \min \left(1, \frac{c_2 - \frac{z}{H_{PBL}}}{c_2 - c_1} \right) \right], \quad c_2 > c_1 \quad (17)$$

$$L_n = \frac{C_\epsilon}{\nu^3} f_w \kappa z + (1 - f_w) L_{AVG} \quad (18)$$



1. Ensures obeying MOST and smooth and stable solution
2. Partly solves the problem of insufficient mixing
3. Problems in convective conditions and near PBL top

- Few more items are needed to finalize the TLS formulation:

1) Regime-dependent minimum TLS near the PBL top (Bechtold and Marquet 2020):

$$L_{\text{BLT}} = \min \left\{ \max \left[L_{\text{BLT}2} + \frac{L_{\text{BLT}2} - L_{\text{BLT}1}}{C_{\Delta 2} - C_{\Delta 1}} (\Delta \theta_s - C_{\Delta 2}), L_{\text{BLT}1} \right], L_{\text{BLT}2} \right\} \quad (19)$$

$$\Delta \theta_s = \theta_s(z = 1.5 \cdot H_{\text{PBL}}) - \theta_s(z = 0) \quad (20)$$

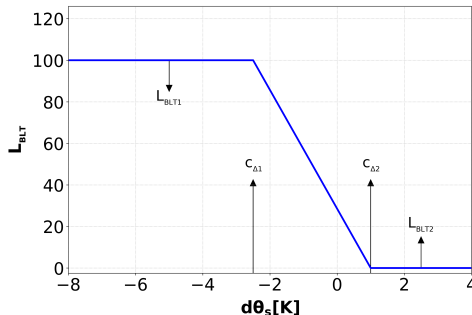
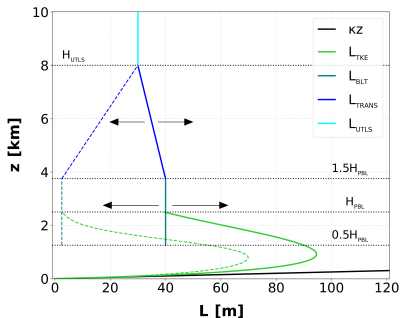
2) Minimum allowed upper-air TLS ($L_{\text{UTLS}} + L_{\text{TRANS}}$):

$$L_{\text{TKE}} = \max(L_{\text{AVG}}, L_{\text{MIN}}), \quad L_{\text{MIN}} = f(L_{\text{BLT}}, L_{\text{TRANS}}, L_{\text{UTLS}}) \quad (21)$$

3) Introduction of global scaling with κ (even smoother transition in the surface layer):

$$L_n^{\text{H24}} = \frac{\kappa C_\epsilon}{\nu^3} [f_w z + (1 - f_w) L_{\text{TKE}}] \quad vs. \quad L_n^{\text{H24}*} = \frac{C_\epsilon}{\nu^3} f_w \kappa z + (1 - f_w) L_{\text{TKE}} \quad (22)$$

► Scheme of the final solution:

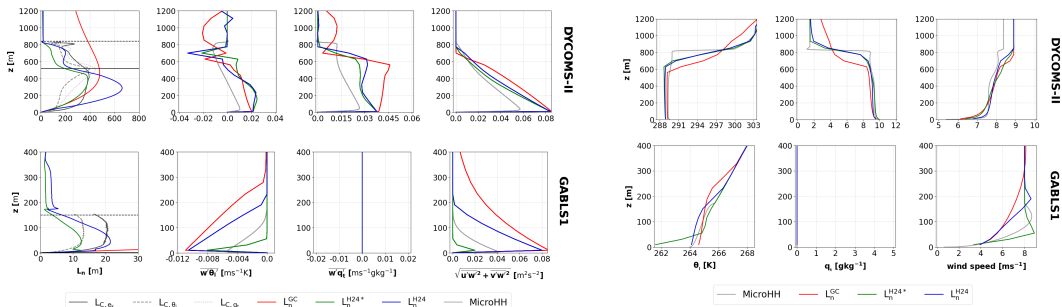


We need a reliable "tool" to validate this → LES-based TLS diagnostics

MicroHH DNS and LES model (van Heerwaarden et al. 2017)

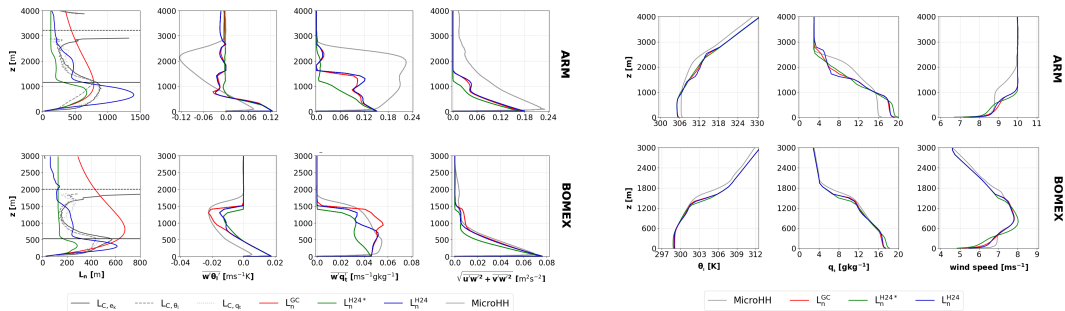
Results: 1D model validation against LES data

- ▶ The evaluation in idealized cases (1-h averaged profiles): (i) more stable PBL (DYCOMS-II and GABLS1; at +3-h and +4.5-h) and (ii) convective PBL (ARM and BOMEX; at +10-h and +7-h)
- ▶ TLS is diagnosed from LES budgets of e_k , s_{sL} and q_t after Bašták Ďurán et al. (2020) and Reilly et al. (2022)
- ▶ An improvement over the L_n^{GC} in DYCOMS-II and GABLS cases, particularly near the surface and the PBL top (key roles of κ scaling and L_{BLT})

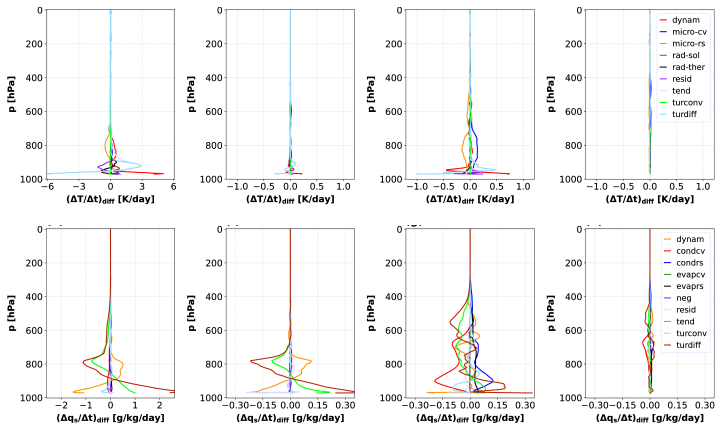


Results: 1D model validation against LES data

- ▶ In ARM and BOMEX cases, L_n^{H24} and L_n^{GC} perform similarly, while L_n^{H24*} underestimates TLS and fluxes
- ▶ Within the cloud layer and ARM case, L_n^{GC} and L_n^{H24} considerably underestimate mixing. However, unlike L_n^{H24*} , their H_{PBL} is comparable to LES
- ▶ The ARM case results point to the importance of counter-gradient terms (TOMs)



► What is the contribution of components we added to the L_n^{H24} solution?



κ scaling

SL smoothing

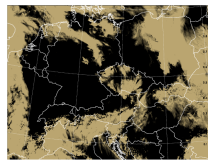
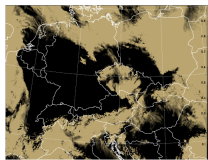
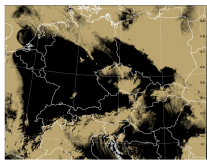
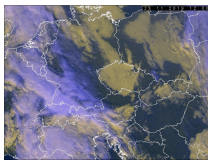
LBLT

$L_{UTLS} + L_{TRANS}$

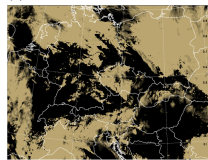
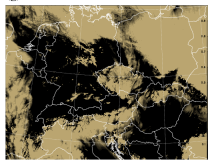
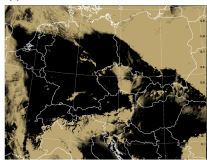
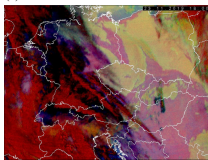
1. The κ scaling is crucial
2. Smoothing diminishes the impact of L_{up}
3. The impact of L_{BLT} is important in all cond.
4. The impact of $L_{UTLS} + L_{TRANS}$ is small
5. The impact of TLS tuning is moderate

- ▶ ALADIN-CZ at $\Delta x = 2.3125$ km and 87 levels, NH-dynamics and ALARO-1 physics
- ▶ Anticyclonic period with persistent inversion over Czechia (23 November 2019 case)

12 UTC



18 UTC



MSG

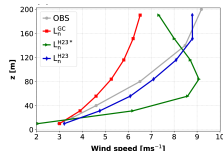
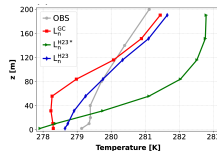
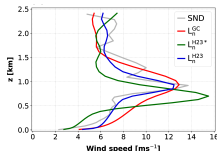
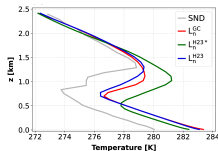
LGC

LH24

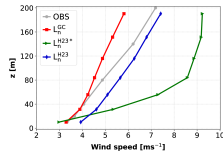
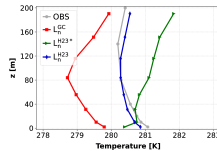
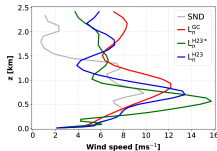
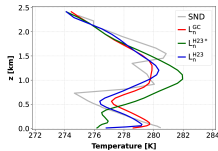
LH24*

- ▶ Slightly improved representation of inversion (Prague-Libuš)
- ▶ Improved averaged daytime and nighttime temperature and wind profiles at Cabauw

12 UTC



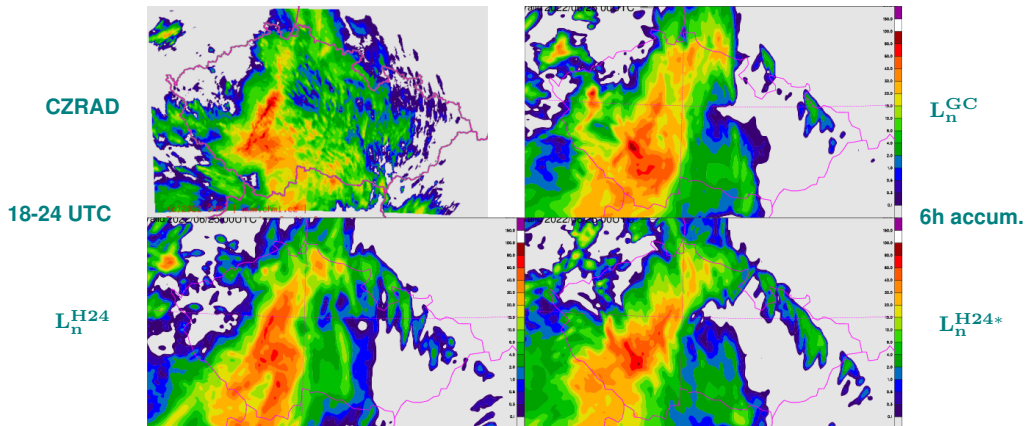
24 UTC



Prague-Libuš

Cabauw

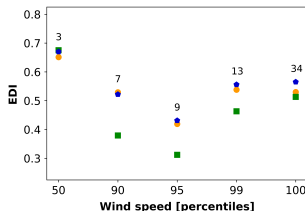
- Mesoscale Convective System 24 June 2022 (> 100 mm of precipitation in Prague)



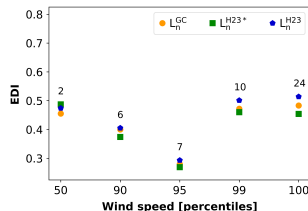
Results: the objective verification

- ▶ The L_n^{H24*} formulation (close to the initial implementation) has 15-25% larger RMSE for most of the surface and upper-air scores (up to 850-700 hPa) for two 10-day periods
- ▶ Due to κ scaling (mainly) and "internal" tuning (C_0 , c_2 , $c_{\Delta 1}$, $c_{\Delta 2}$ and L_{BLT1}), the statistical scores for L_n^{H24} are nearly neutral to L_n^{GC} ; confirmed on additional 3-5 day periods
- ▶ However, there are also some improvements:

1. BIAS and RMSE of cloudiness ($\sim 2\%$)
2. STD of T2m ($\sim 1.5\text{-}2\%$)
3. Extreme 10-m wind (FB and EDI $\sim 3.5\text{-}6.5\%$)
4. Upper-air rel. humidity



winter



summer

- ▶ The settings of L_n^{H24} TLS formulation are more or less confirmed by LES-diagnostics and 1D model evaluation, yielding also satisfying first results within the 3D model
- ▶ As expected, the improvement is mainly seen in statically stable conditions
- ▶ Further validation and tuning of other model components/processes is needed
- ▶ Despite similar attempts (L_{MIN} and H_{PBL} method), the L_n^{GC} formulation is not improved
- ▶ Future work aims to address: (i) the scale-awareness of TLS and TOUCANS scheme, (ii) the 3D aspects of turbulence and (iii) the treatment of the stable PBL in situations with weak wind

The related manuscript is in revision at Monthly Weather Review:

Hrastinski, M., Mašek, J., Bašták Ďurán, I., Grisogono, B. and Brožková, R., 2024: Regime-dependent turbulence length scale formulation for NWP models based on turbulence kinetic energy, shear and stratification.

- ▶ Bašták Ďurán, I., J.-F. Geleyn, and F. Váňa (2014): A Compact Model for the Stability Dependency of TKE Production–Destruction–Conversion Terms Valid for the Whole Range of Richardson Numbers. *J. Atmos. Sci.*, 71, 3004–3026.
- ▶ Bašták Ďurán, I., J.-F. Geleyn, F. Váňa, F., Schmidli, J. and R. Brožková, (2018): A Turbulence Scheme with Two Prognostic Turbulence Energies. *J. Atmos. Sci.*, 75, 3381–3402.
- ▶ Bašták Ďurán, I., J. Schmidli, and R. Bhattacharya, (2020): A Budget-Based Turbulence Length Scale Diagnostic. *Atmosphere*, 11(4).
- ▶ Marquet, P. and P. Bechtold, (2020): A new Estimated Inversion Strength (EIS) based on the moist-air entropy. *Research activities in Earth system modelling. Working Group on Numerical Experimentation (Report No50. WCRP Report No12/2020. No. 50*, E. Astakhova, Ed., WMO, Geneva.
- ▶ Mašek, J., I. Bašták Ďurán and R. Brožková, (2022): Stable Numerical Implementation of a Turbulence Scheme with Two Prognostic Turbulence Energies. *Mon. Wea. Rev.*, 150(7), 1667–1688.
- ▶ Redelsperger, J. L., F. Mahé, and P. Carlotti, (2001): A Simple And General Subgrid Model Suitable Both For Surface Layer And Free-Stream Turbulence. *Bound.-Lay. Meteorol.*, 101, 375–408.
- ▶ Reilly, S., I. Bašták Ďurán, A. Theethai Jacob, and J. Schmidli, (2022): A Simple And General Subgrid Model Suitable Both For Surface Layer And Free-Stream Turbulence. *Bound.-Lay. Meteorol.*, 101, 375–408
- ▶ Rodier, Q., V. Masson, F. Couvreux and A. Paci, (2017): Evaluation of a Buoyancy and Shear Based Mixing Length for a Turbulence Scheme. *Front. Earth Sci.*, 5.
- ▶ van Heerwaarden, C. C., B. J. H. van Stratum, T. Heus, J. A. Gibbs, E. Fedorovich, and J. P. Mellado, (2017): MicroHH 1.0: a computational fluid dynamics code for direct numerical simulation and large-eddy simulation of atmospheric boundary layer flows. *Geosci. Model Dev.*, 10, 3145–3165.

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Thank you for your attention!



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