

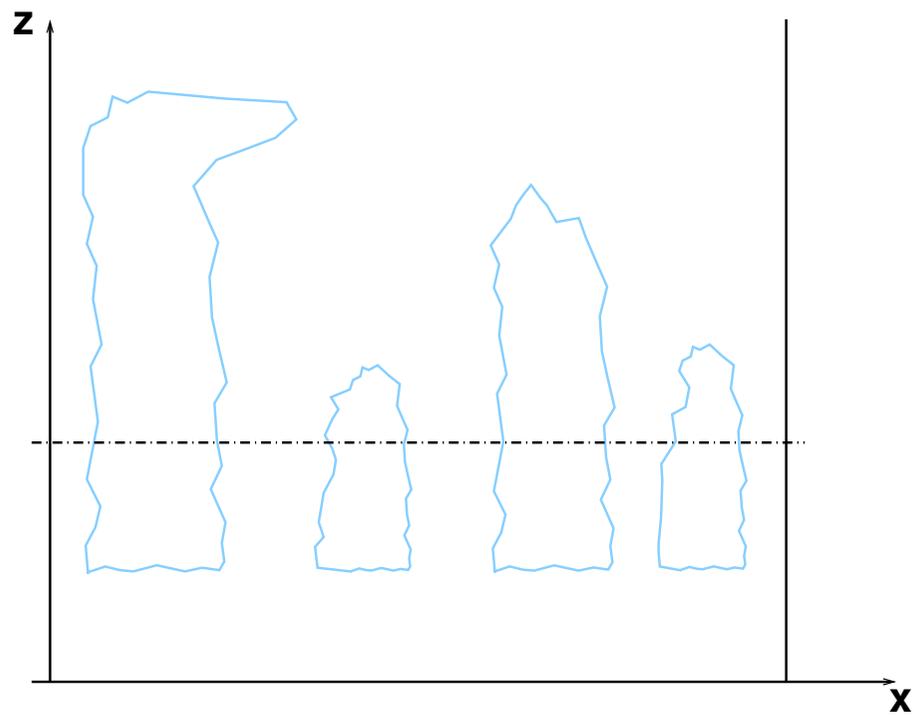


Geometrical aspects of sub-grid convective condensation and precipitation

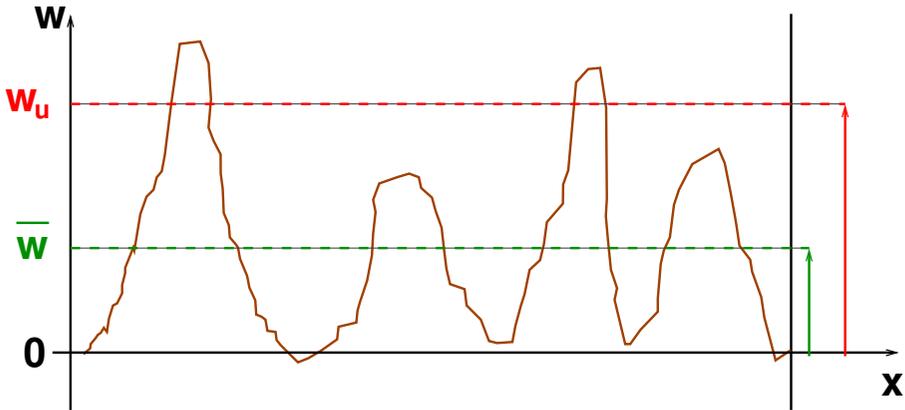
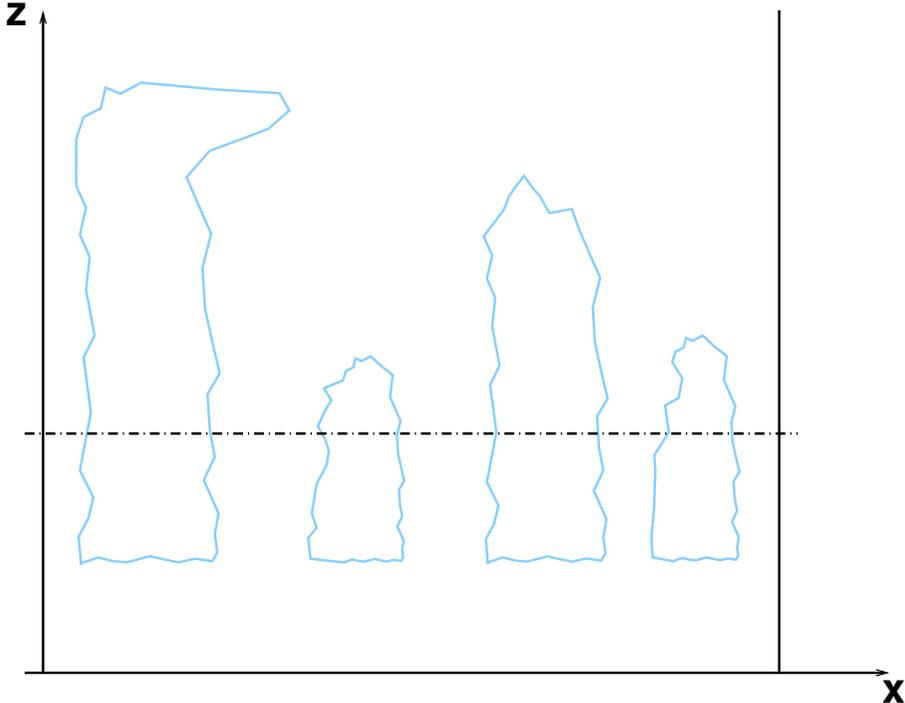
Luc Gerard

14 June 2012

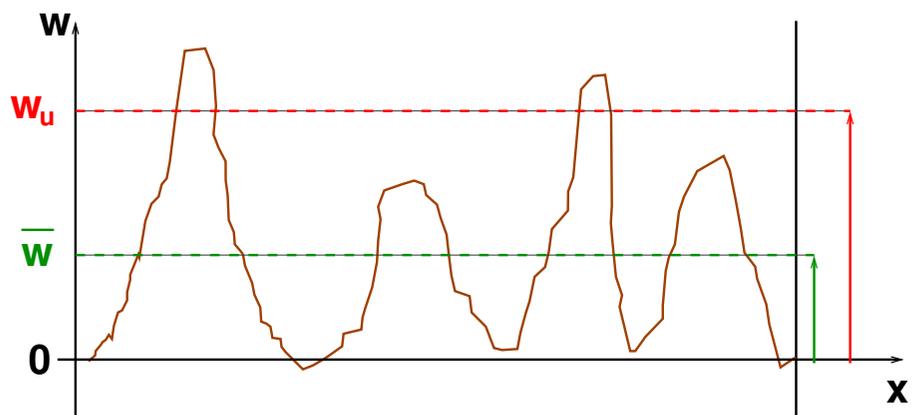
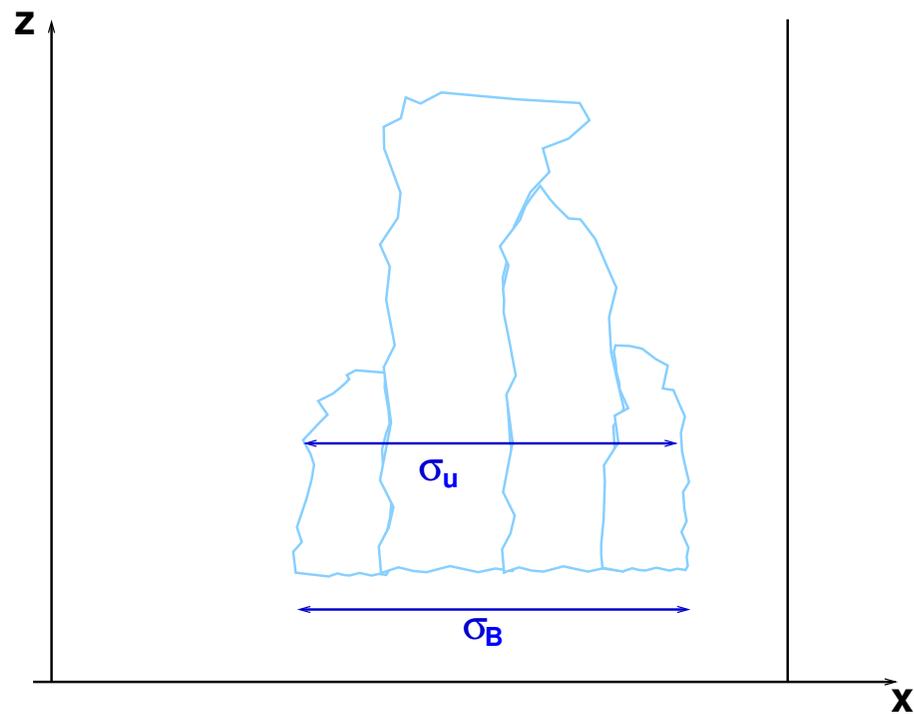
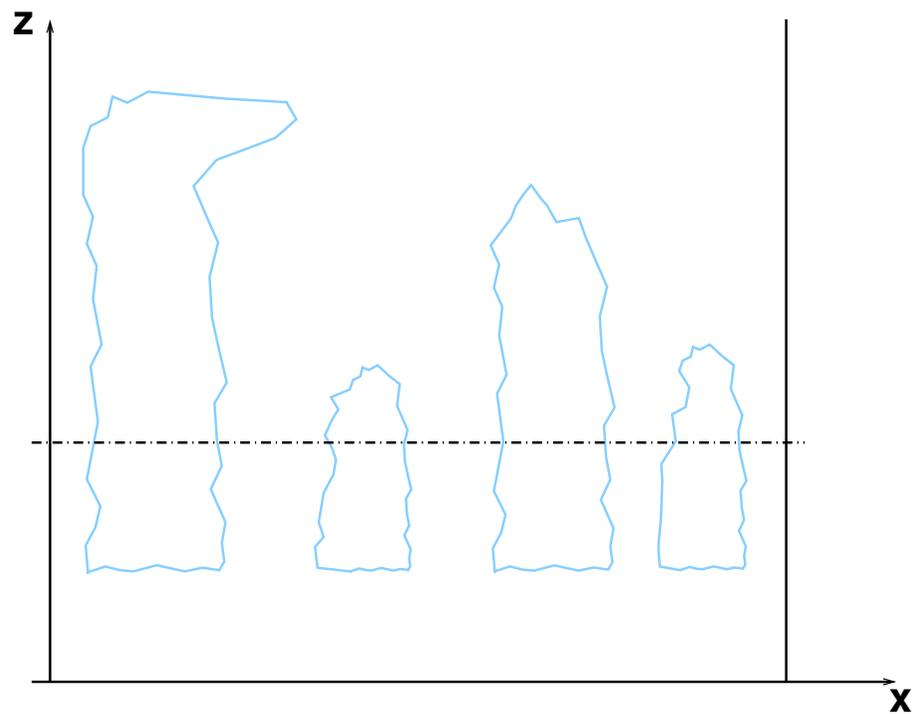
Subgrid variability



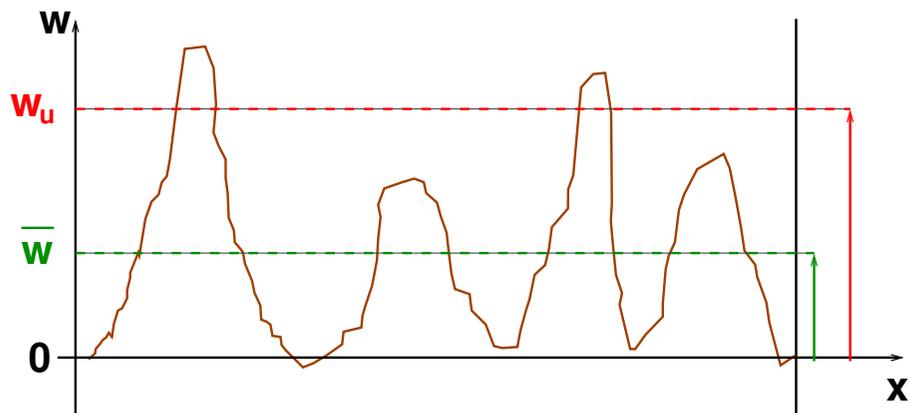
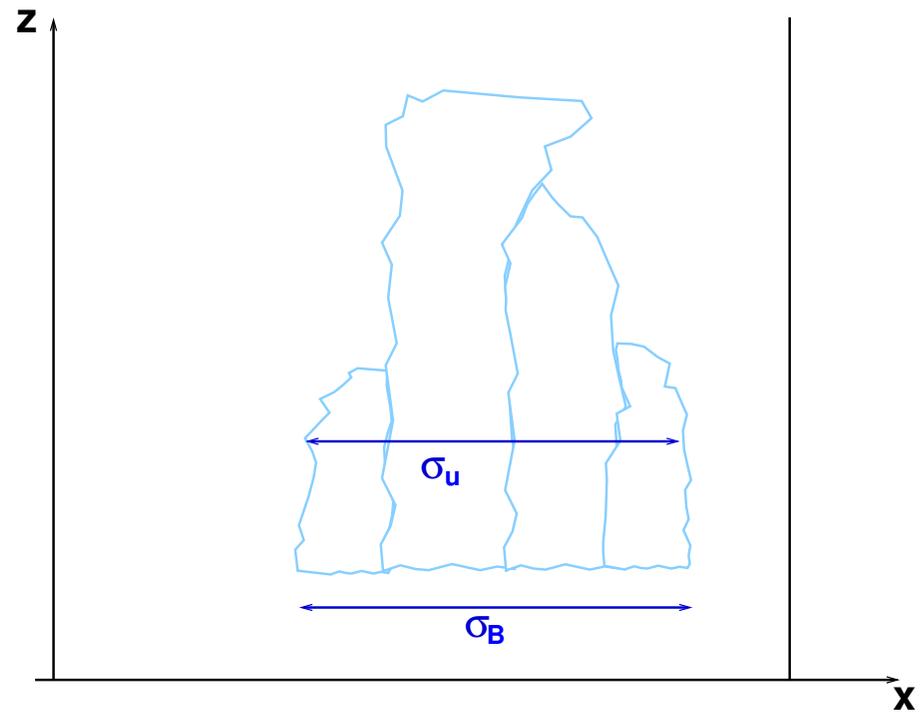
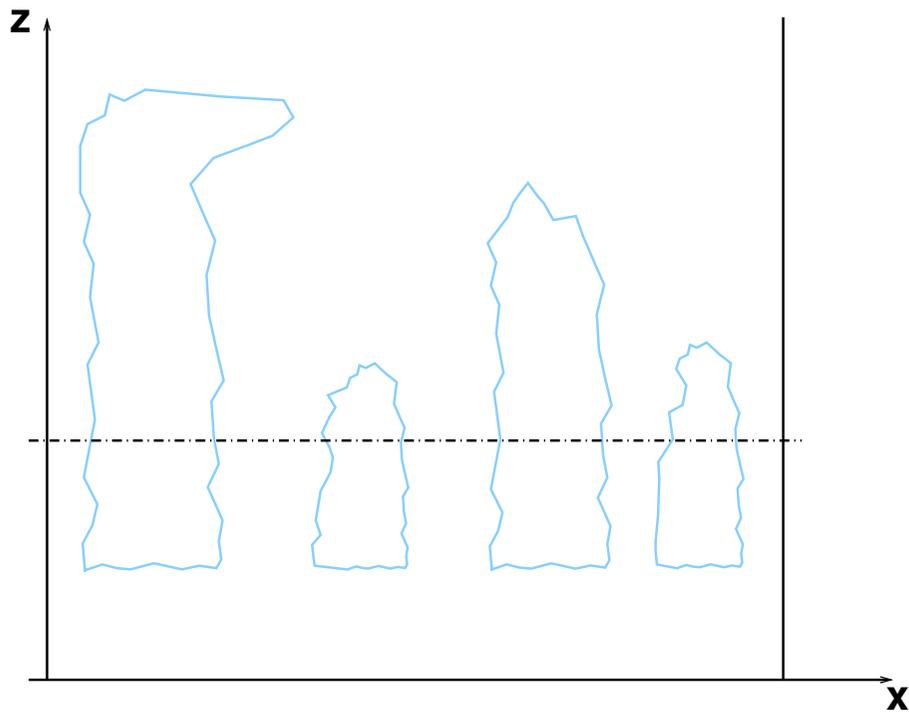
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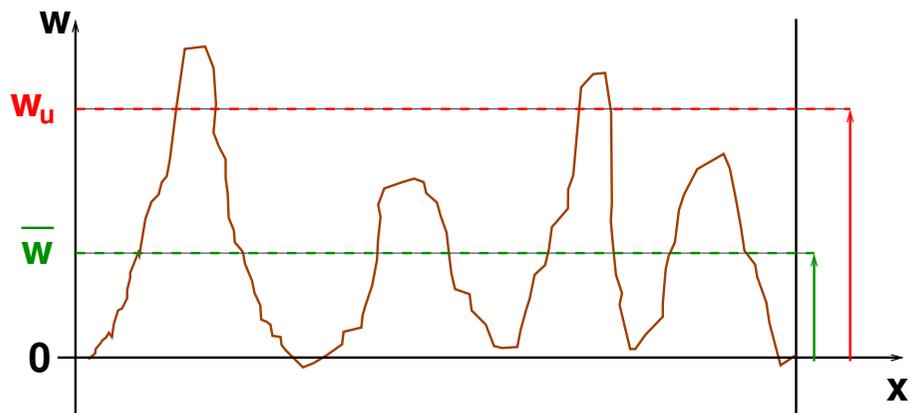
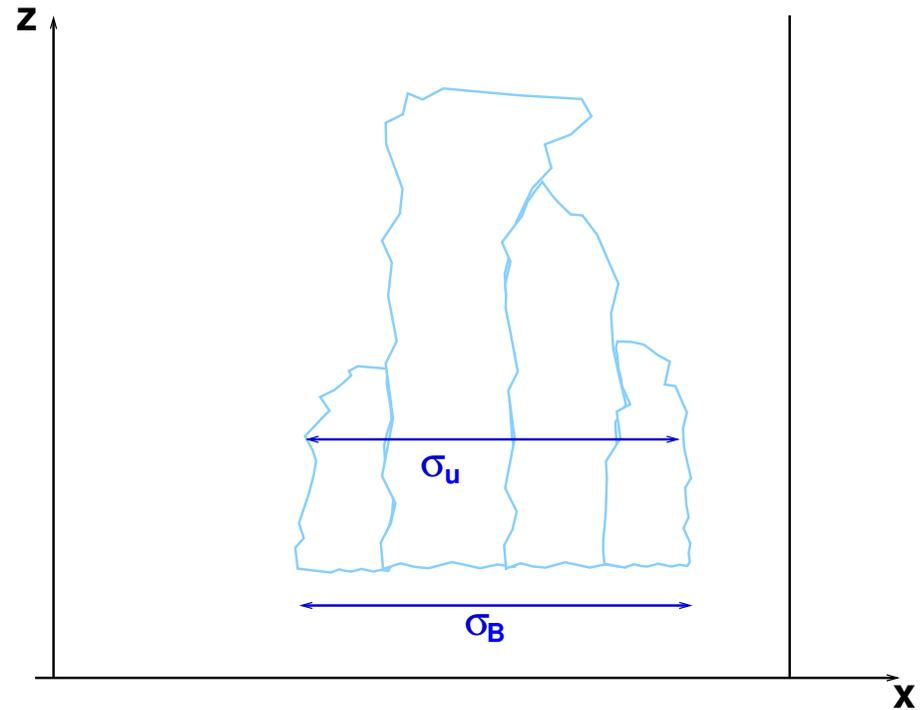
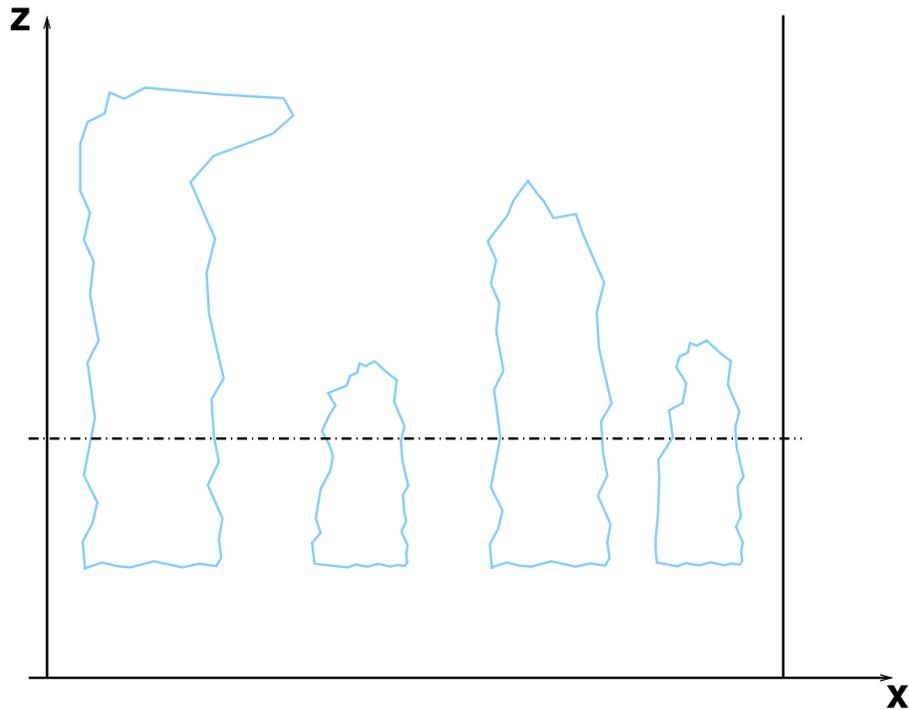


Subgrid variability



$$\begin{aligned}\bar{\psi} &= \sum_u \sigma_i \psi_i + \sum_e \sigma_j \psi_j \\ &= \sigma_u \psi_u + (1 - \sigma_u) \psi_e\end{aligned}$$

Subgrid variability



$$\bar{\psi} = \sum_u \sigma_u \psi_u + \sum_e \sigma_e \psi_e$$

$$= \sigma_u \psi_u + (1 - \sigma_u) \psi_e$$

$$\psi_u^\diamond = \psi_u - \bar{\psi}, \quad \psi_e^\diamond = \psi_e - \bar{\psi},$$

$$\Rightarrow \sigma_u \psi_u^\diamond + \sigma_e \psi_e^\diamond = 0$$

Complementary Subgrid Updraft

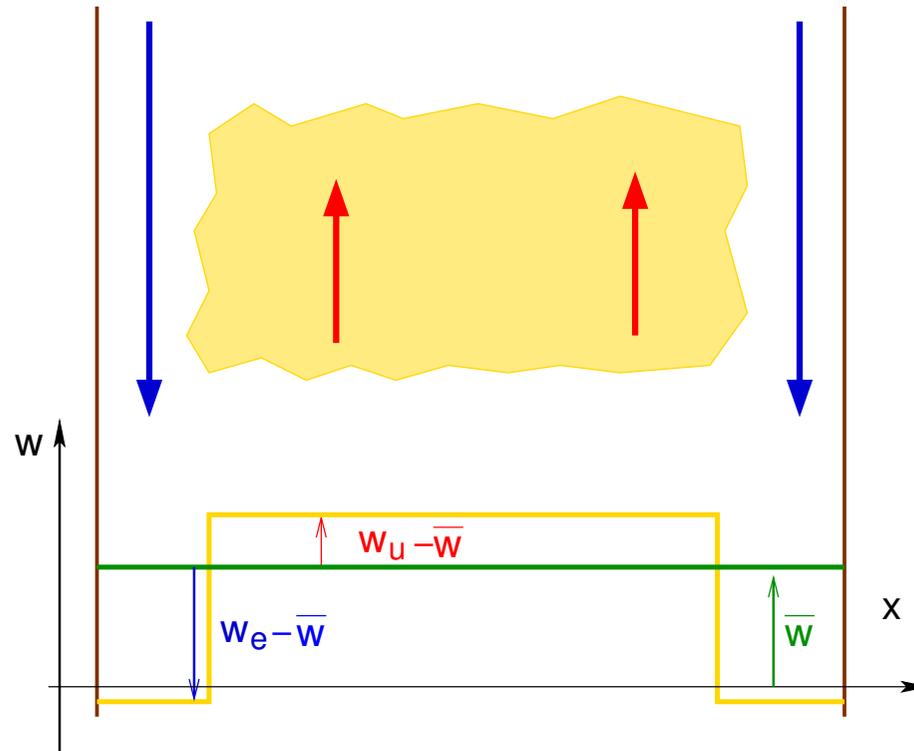
Provide complementary contribution to the resolved updraft

- Sequential physics (3MT cascade)

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- Triggering of subgrid scheme \neq triggering of convective updraft

Anelastic equation

$$\rho_0 = \rho_0(z) \quad \Longrightarrow \quad \nabla \cdot \mathbf{u} = -w \frac{d \ln \rho_0}{dz}$$

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$$\underbrace{\frac{\partial(\sigma_u \psi_u)}{\partial t}}_{\text{tendency}} + \underbrace{\frac{\mathcal{L}_b}{\rho_0 g A} \eta |w_u - w_e| (\psi_u - \psi_e)}_{\text{turb. mixing}} + \underbrace{\frac{\partial \sigma_u \omega_u}{\partial p} (\psi_u - \psi_b)}_{\text{organized entrainment}} + \underbrace{\sigma_u \omega_u \frac{\partial \psi_u}{\partial p}}_{\text{vertical transport}}$$

$$\underbrace{-\psi_b \frac{\partial \sigma_u}{\partial t}}_{\text{creation } \psi_b} \quad \underbrace{-\overline{\omega' \psi'}^b \frac{\partial \sigma_u}{\partial p}}_{\text{subplume entrainment}} \quad \underbrace{+\frac{\partial \sigma_u \overline{\omega' \psi'}^u}{\partial p}}_{\text{subplume vertic transport}} = \underbrace{\sigma_u \overline{f_\psi}}_{\text{forcing}}$$

Derivation of CSU equations

updraft:
$$\sigma_u \frac{\partial \psi_u}{\partial t} = \sigma_u \frac{\lambda_u}{\rho_0} (\omega_u - \omega_e) + \frac{\partial \sigma_u \omega_u}{\partial p} (\psi_u - \psi_b) + \sigma_u \omega_u \frac{\partial \psi_u}{\partial p} = \sigma_u \overline{f_\psi^u}$$

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resolved:
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Gathering and transforming yields

$$\frac{\partial \psi_u^\diamond}{\partial t} - \Lambda \frac{\omega_u^\diamond \psi_u^\diamond}{1 - \sigma_u} + (\omega_u^\diamond + \omega_e) \frac{\partial \psi_u^\diamond}{\partial p} + \omega_u^\diamond \frac{\partial \bar{\psi}}{\partial p} = (1 - \sigma_u) [\overline{f_{\psi sm}^u} - \overline{f_{\psi sm}^e}]$$

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$$\Lambda_w \approx \frac{\lambda_u + \mathcal{K}_{du}}{\rho_0 (1 - \sigma_u)} + \underbrace{\left(\sigma_u - \frac{\delta_{oe}}{k} \right) \frac{1}{M_u^*} \frac{\partial M_u^*}{\partial p}}_{\Lambda_{dyn}}, \quad k = \frac{\omega_u^\diamond}{\omega_u} \sim (1 - \sigma_u) \text{ if } \omega_e \sim 0$$

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$\mathcal{K}_{du} = \text{TUDFR}$ used for momentum only

Further considerations

- Source

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Final equations

$$\left. \frac{\partial \omega_u^\diamond}{\partial t} \right|_{sm} - \left(\frac{\Lambda_w}{1 - \sigma_u} + \frac{1}{k} \frac{d \ln \rho_0}{dp} \right) \omega_u^{\diamond 2} + \frac{\omega_u^\diamond}{k} \frac{\partial \omega_u^\diamond}{\partial p} = -\alpha_b \rho_0 g^2 \frac{T_{vu} - \overline{T_v}}{\overline{T_v}}$$

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$$\Rightarrow \psi_u^\diamond = \psi_b^\diamond e^{\left(\frac{\Lambda \Delta p}{1 - \sigma_u} \right)} - \frac{(1 - \sigma_u)}{\Lambda \Delta p} \left[(1 - \sigma_u) (\psi_h - \psi_b) - (\bar{\psi}^l - \bar{\psi}^{l+1}) \right] (1 - e^{\left(\frac{\Lambda \Delta p}{1 - \sigma_u} \right)})$$

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for this,

- first guess from moist adiabatic ascent $\rightarrow \{T_u^0, q_u^0, \delta q_{ca}^0\}$;
- interpolate linearly $\rightarrow \{T_u^+, q_u^+ = q_{\text{sat}}(p, T_u^+), \delta q_{ca}^+\}$

Steady-state ascent calculation

- Updraft base:
 - LCL from triggering routine
 - Top dry ascent velocity

$$\omega_{u\text{top dry}}^{\diamond} \sim \omega_{u\text{free}}^{\diamond} \exp\left[-\frac{p^{b+1} - p^b}{\Delta p_{bx}}\right], \quad \frac{1}{\Delta p_{bx}} = \text{gidpbas}$$

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- start with k and σ_B^{\bullet} guessed from $\bar{\omega}, \sigma_u^-, \omega_u^{\diamond-}$:

$$\sigma_B^* = \min\left[\text{gcvalmx}, \max\left[\text{zepsaln}, \langle \sigma_u^- \rangle, \text{sigig} \frac{\sum \frac{\bar{\omega}}{\omega_u^{\diamond-} + \bar{\omega}} \Delta p}{\sum \Delta p} + (1 - \text{sigig}) \langle \sigma_u^- \rangle\right]\right]$$

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- σ_u vertical variation: ensembling effect, entrainment-detrainment budget (effect of moisture...)

$$(1 - \sigma_u) = \mu(1 - \sigma_B), \quad \mu = f(x, \mu^{\text{TOP}}),$$

$$x = \frac{p^{\text{LCL}} - p}{p^{\text{LCL}} - p^{\text{TOP}}}, \quad \mu^{\text{TOP}} \rightarrow 1 \text{ if } \sigma_B^{\bullet} \rightarrow 1$$

p^{TOP} estimated for the non-diluted ascent.

Steady-state ascent calculation (continued)

- stationnarized ω_u^\diamond equation
⇒ Organized entr/detr (restrained to $\pm \text{gcvendymax} \cdot \Delta\phi$),

$$\Lambda_{\text{dyn}} \Delta p = \left(\sigma_u^\bullet - \frac{\delta_{oe}}{k} \right) \frac{\Delta M_u^*}{M_u^*}, \quad M_u^* = \frac{\sigma_u^\bullet \omega_u^\diamond}{1 - \sigma_u^\bullet}$$

Steady-state ascent calculation (continued)

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\Rightarrow Guess at next level.

$$(\Lambda_{\text{dyn}}\Delta p)^{l*} = (\Lambda_{\text{dyn}}\Delta p)^{l+1} \begin{array}{l} - \\ + \end{array} \delta_{\text{asc}}^{l+1} \cdot \text{gcvendy1} \cdot \Delta\phi^{\bar{l}} \cdot e^{-\max(0, \text{gcvendy2}(\phi^{l+1} - \phi_b))} \quad \text{if } \begin{array}{l} > 0 \\ < 0 \end{array}$$

Steady-state ascent calculation (continued)

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 \Rightarrow Organized entr/detr (restrained to $\pm \text{gcvendymax} \cdot \Delta\phi$),

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$$(\Lambda_{\text{dyn}}\Delta p)^{l*} = (\Lambda_{\text{dyn}}\Delta p)^{l+1} \begin{matrix} - \\ + \end{matrix} \delta_{\text{asc}}^{l+1} \cdot \text{gcvendy1} \cdot \Delta\phi^{\bar{l}} \cdot e^{-\max(0, \text{gcvendy2}(\phi^{l+1} - \phi_b))} \quad \text{if } \begin{matrix} > \\ < \end{matrix} 0$$

\Rightarrow Mixing

$$\Lambda\Delta p = \frac{\lambda_u \Delta\phi}{1 - \sigma_u^\bullet} + \Lambda_{\text{dyn}}\Delta p$$

Steady-state ascent calculation (continued)

- stationnarized ω_u^\diamond equation
 \Rightarrow Organized entr/detr (restrained to $\pm \text{gcvendymax} \cdot \Delta\phi$),

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- Steady-state for q_u^\diamond and θ_u^\diamond .

Steady-state closure in the CSU context

* Potential Energy Convertibility (Yano et al. 2005) (**LPEC=T**):

$$\text{PEC} = \int_b^t m \frac{\theta_{vu} - \overline{\theta}_v}{\overline{\theta}_v} d\phi = -R_a \int_b^t m (T_{vu} - \overline{T}_v) \frac{dp}{p}, \quad m = \frac{\omega}{\omega^*} \text{ (or 1 for CAPE)}$$

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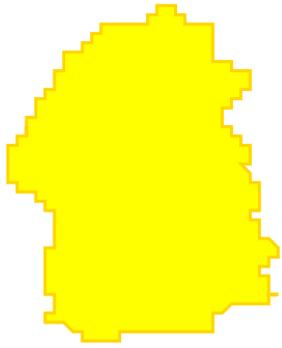
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...by the *real-world* deep convective process if it was *alone*.

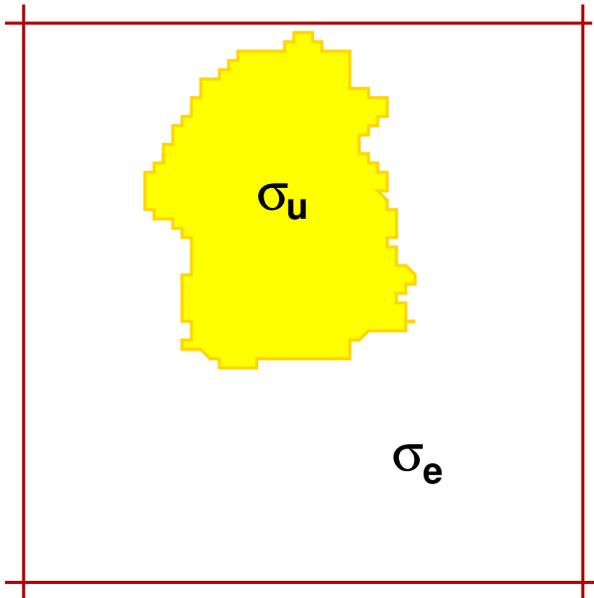
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model-column CAPE < 'environmental CAPE'

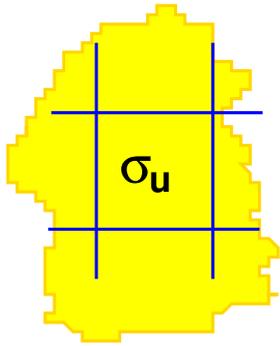


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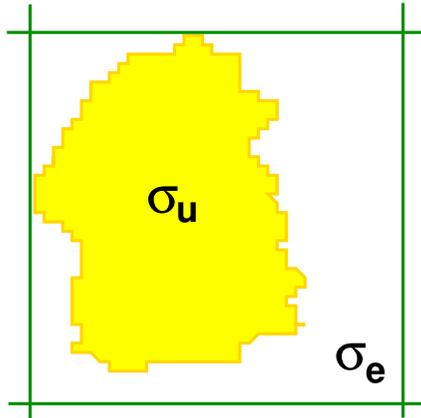
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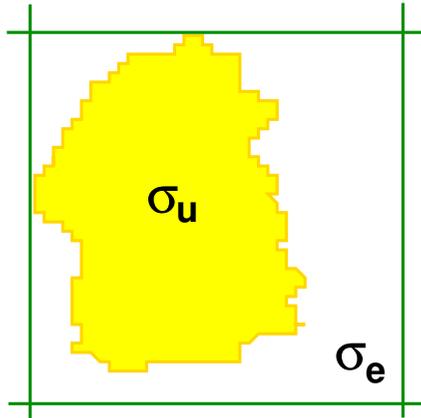
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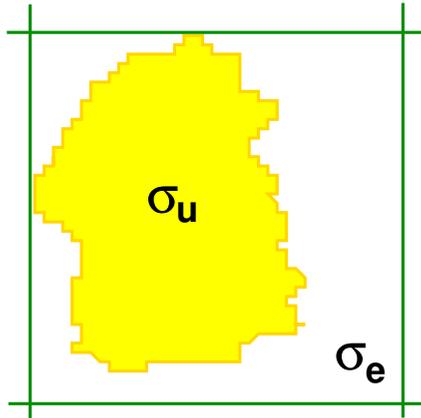
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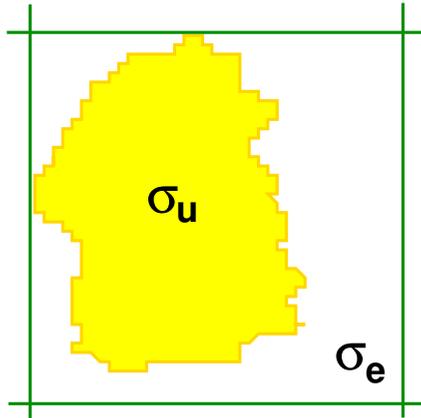
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The mean grid-box state $\bar{\psi}$ input to the updraft routine

- is already affected by resolved vertical motion $\bar{\omega}$
- has been updated after the resolved condensation (3MT cascade)

$$\bar{s}_1 = \bar{s}_0 + gL \frac{\partial F_{cs}}{\partial p} \Delta t,$$

$$\bar{q}_1 = \bar{q}_0 - g \frac{\partial F_{cs}}{\partial p} \Delta t$$

Environmental CAPE / PEC

$$\text{eCAPE} = -R_a \int_b^t m(T_{vu} - T_{ve0}) \frac{dp}{p} \approx -R_a \int_b^t m \frac{T_{vu}^\diamond - \frac{L}{c_p} \frac{\partial F_{cs}}{\partial p} \Delta t}{(1 - \sigma_u)} \frac{dp}{p}$$

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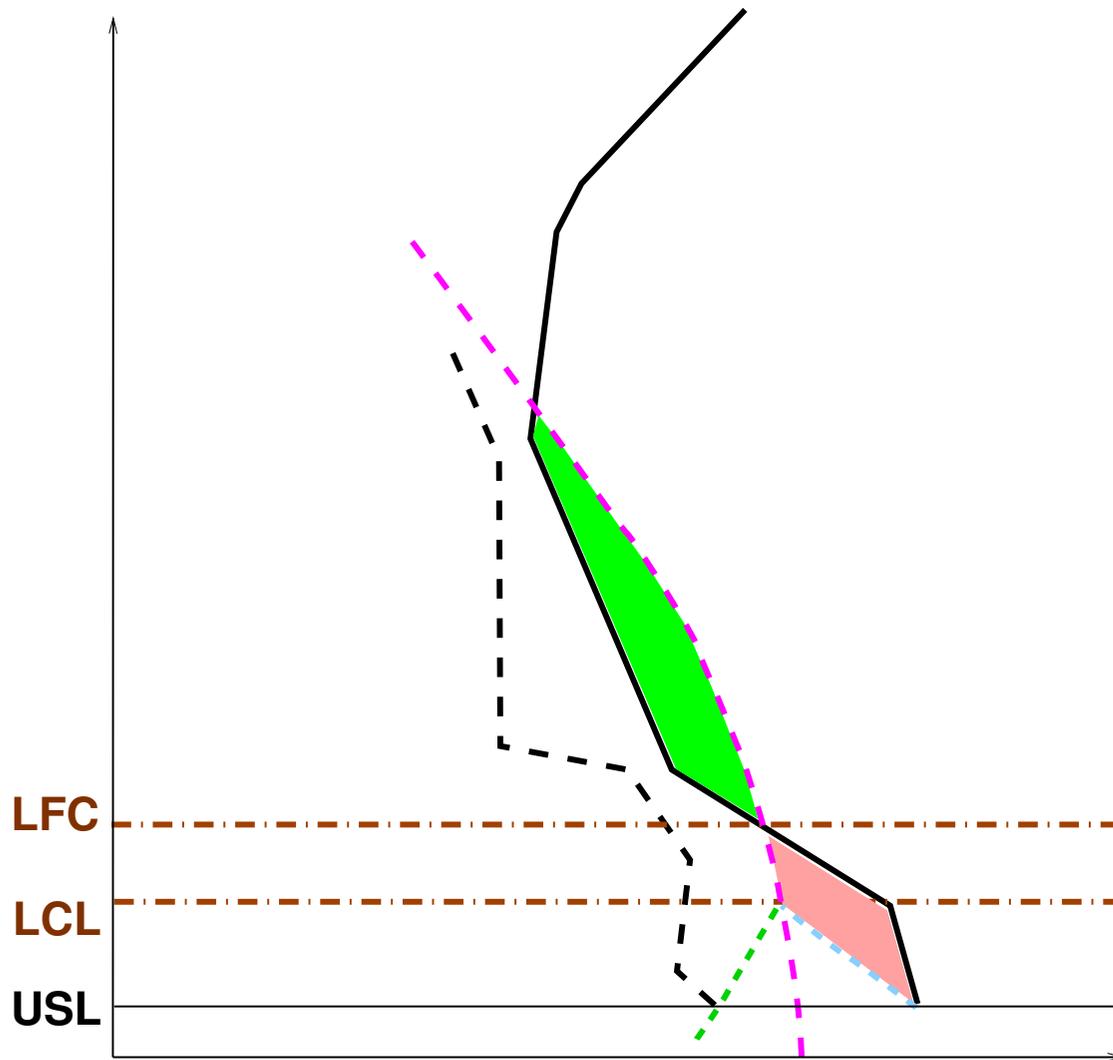
Then

$$\frac{\partial \text{eCAPE}}{\partial t} = - \frac{\text{eCAPE}}{\tau}$$

will lead $(1 - \sigma_B)$ if the profile $\mu = \frac{(1 - \sigma_u)}{(1 - \sigma_B)}$ is given.

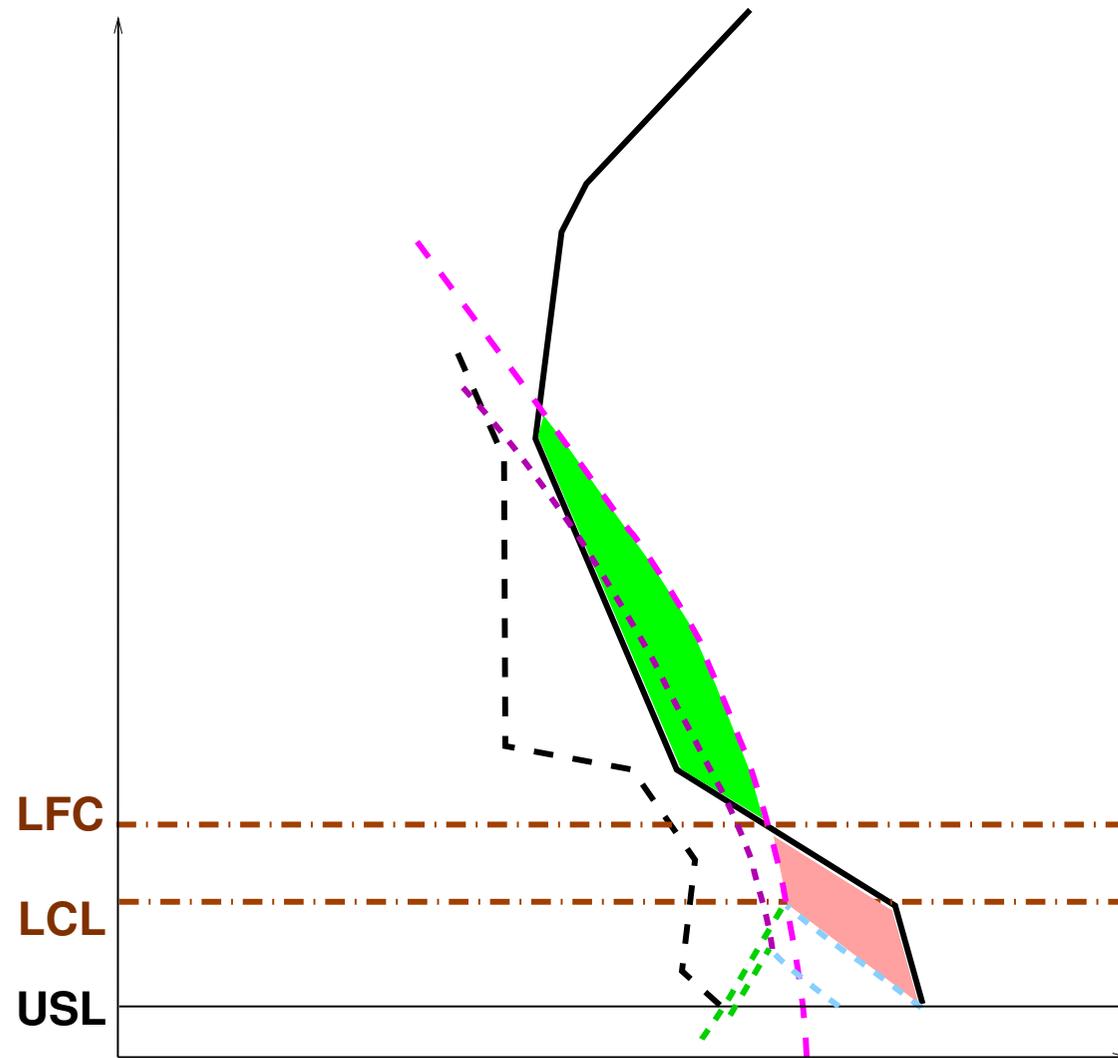
Downdraft effect on updraft closure

Main effect of downdraft is cooling and moistening USL



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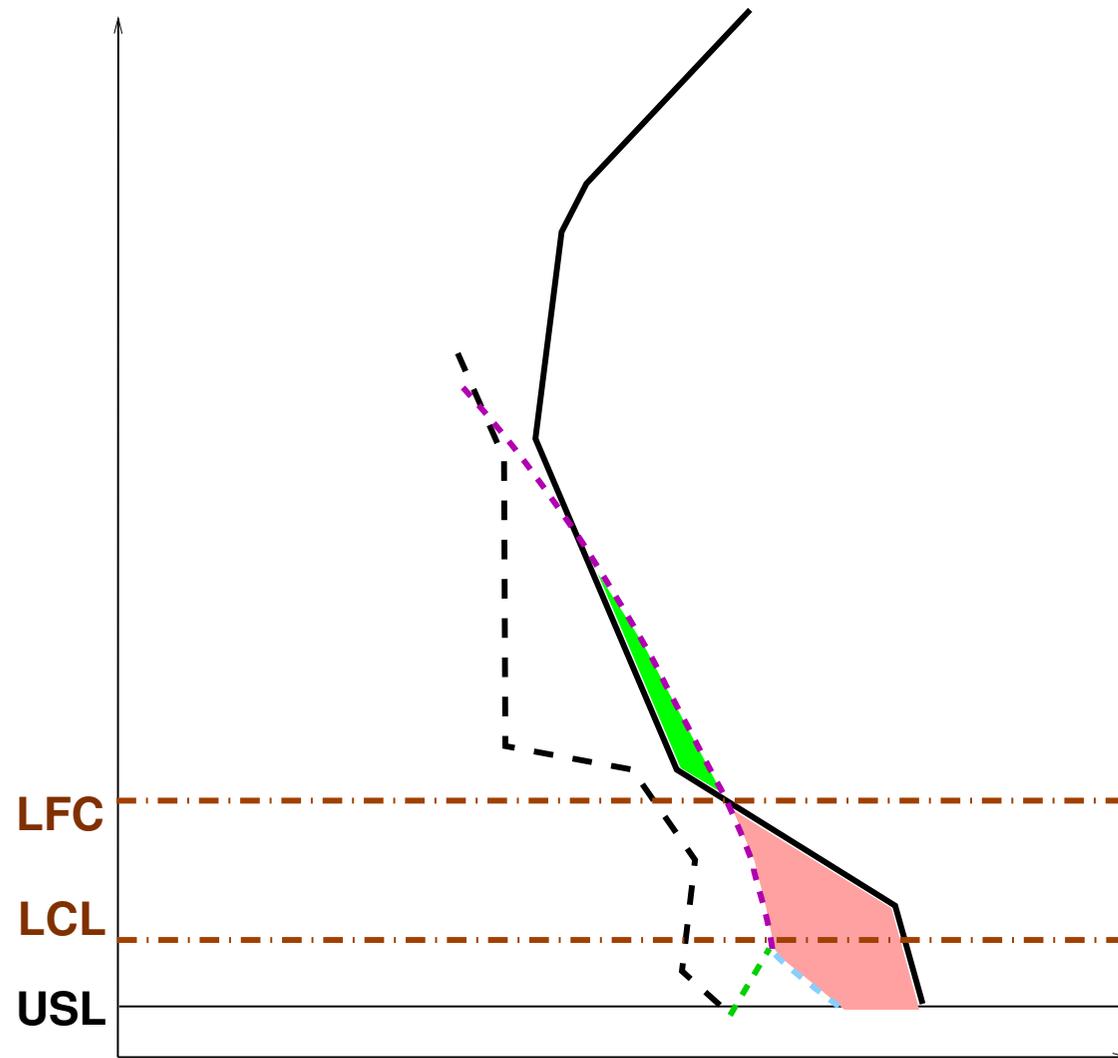
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Downdraft effect on updraft closure

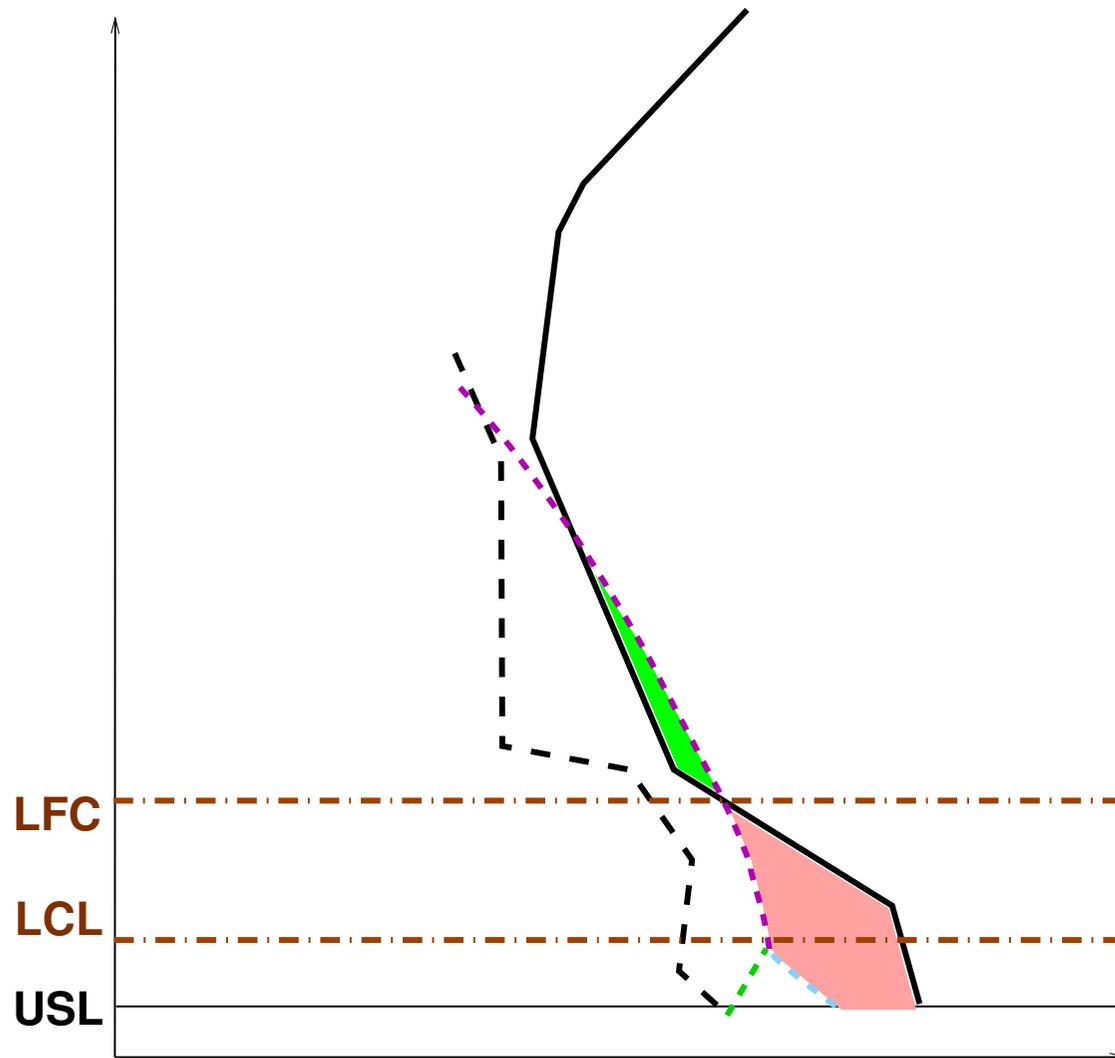
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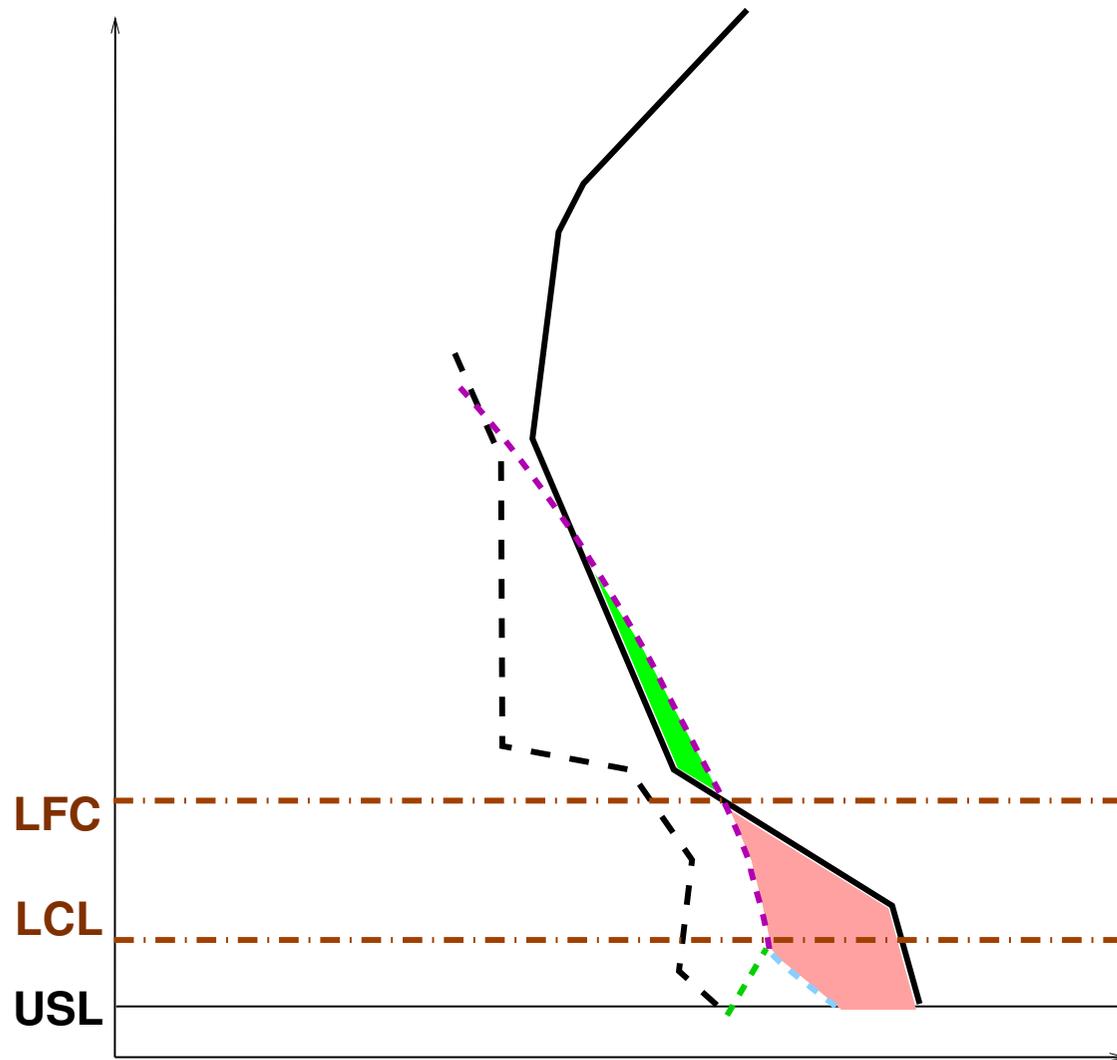
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Downdraft effect on updraft closure

M_d^* advected from previous time step

$$\left. \frac{\partial \bar{T}^{USL}}{\partial t} \right|_{dd} \approx - \frac{\sum_{uslb}^{uslt} M_d^* \frac{1}{c_p} \frac{\partial \bar{s}}{\partial p} \Delta p}{\Delta p^{mix}}, \quad \left. \frac{\partial \bar{q}^{USL}}{\partial t} \right|_{dd} \approx - \frac{\sum_{uslb}^{uslt} M_d^* \frac{\partial \bar{q}}{\partial p} \Delta p}{\Delta p^{mix}}, \quad \left. \frac{\partial p^{USL}}{\partial t} \right|_{dd} \sim 0$$

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Bolton (1980)

$$T^{\text{LCL}} = \frac{2840}{3.5 \ln T^{\text{USL}} - \ln \frac{e^{\text{USL}}}{100} - 4.805} + 55., \quad e^{\text{USL}} = \frac{q^{\text{USL}}}{q^{\text{USL}} + \frac{R_a}{R_v} (1 - q^{\text{USL}})} p^{\text{USL}}$$

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Assume a constant $\Delta \theta_{vu}$ along the updraft

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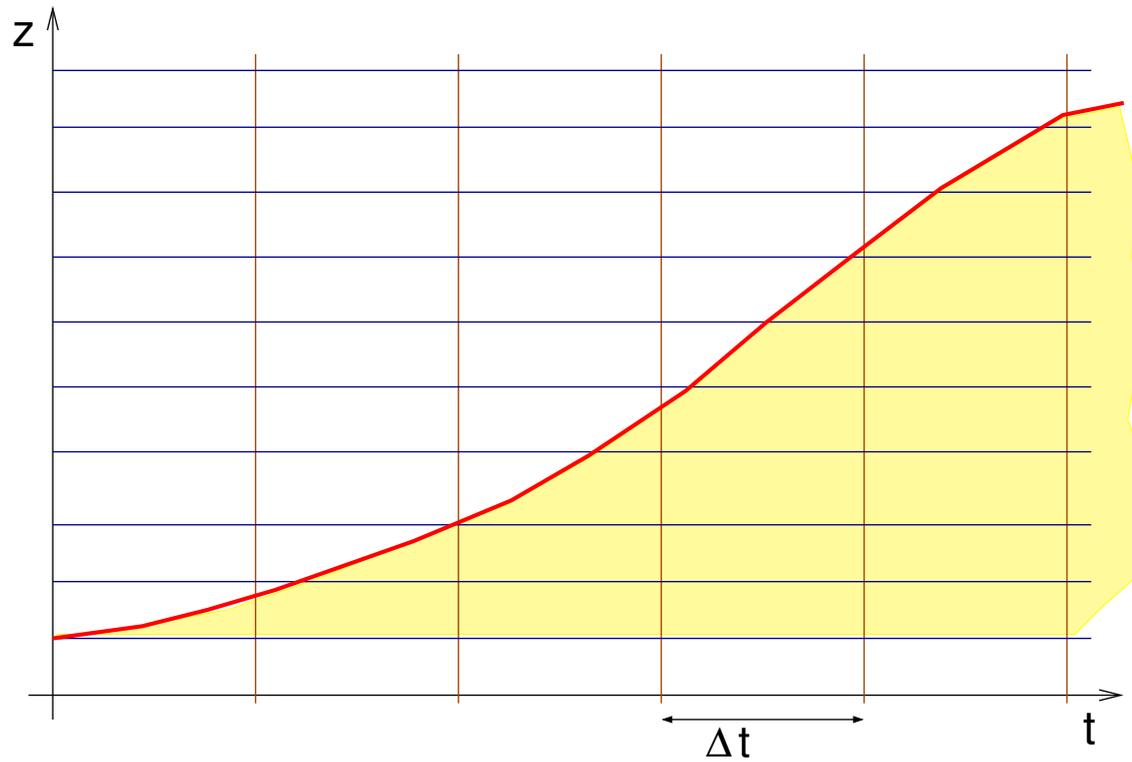
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Updraft evolution

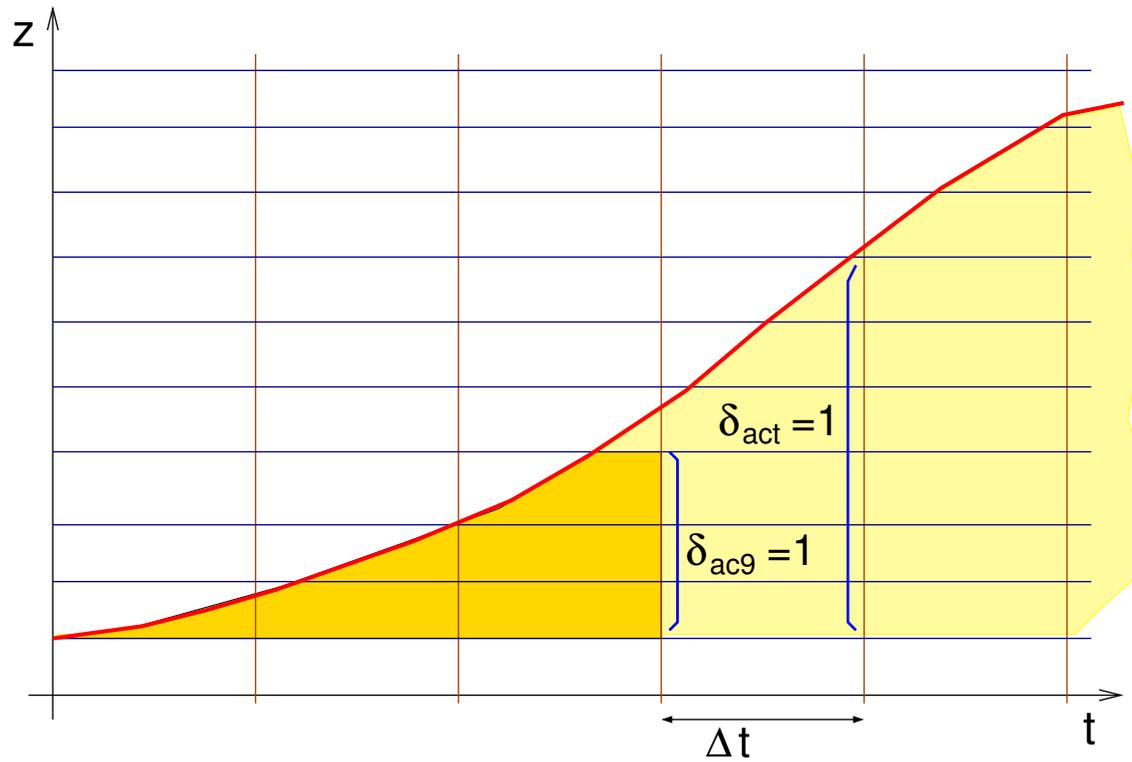
- Gradual elevation of the equivalent cloud top because:
 - The small grid box cannot contain a whole ensemble of clouds of different heights;
 - the time step is too short for the clouds to reach their full height in one time step.

Top evolution: activity index

Top evolution: activity index



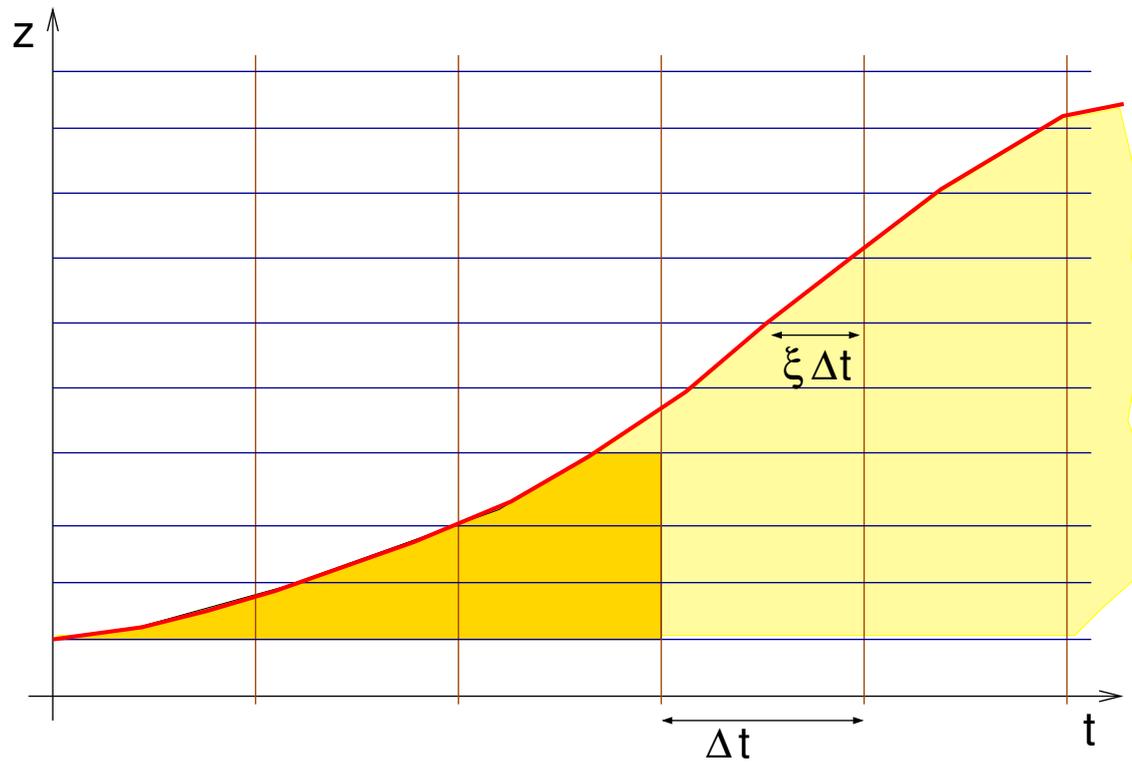
Top evolution: activity index



$\delta_{act} = 1$ at levels reached by the ascent originating at the base

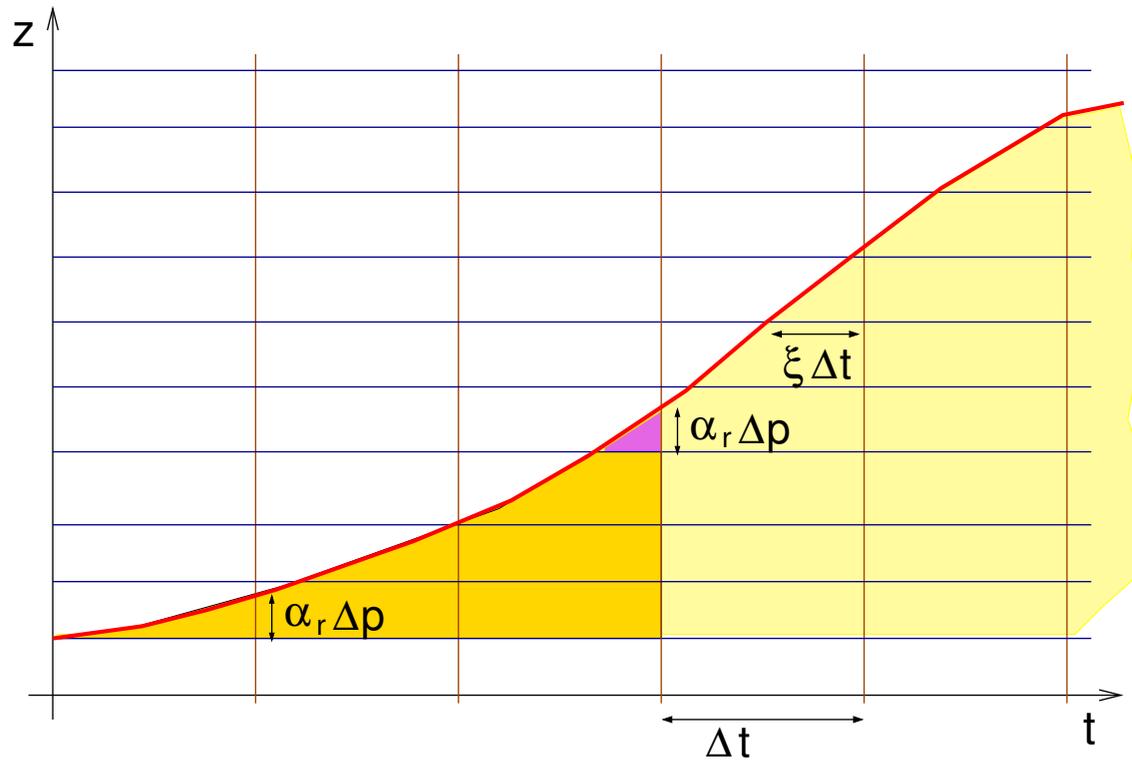
δ_{ac9} retrieved from profile of ω_u^- or σ_u^-

Top evolution: activity index



Buoyancy accelerates the fluid during $\xi \Delta t$

Top evolution: activity index

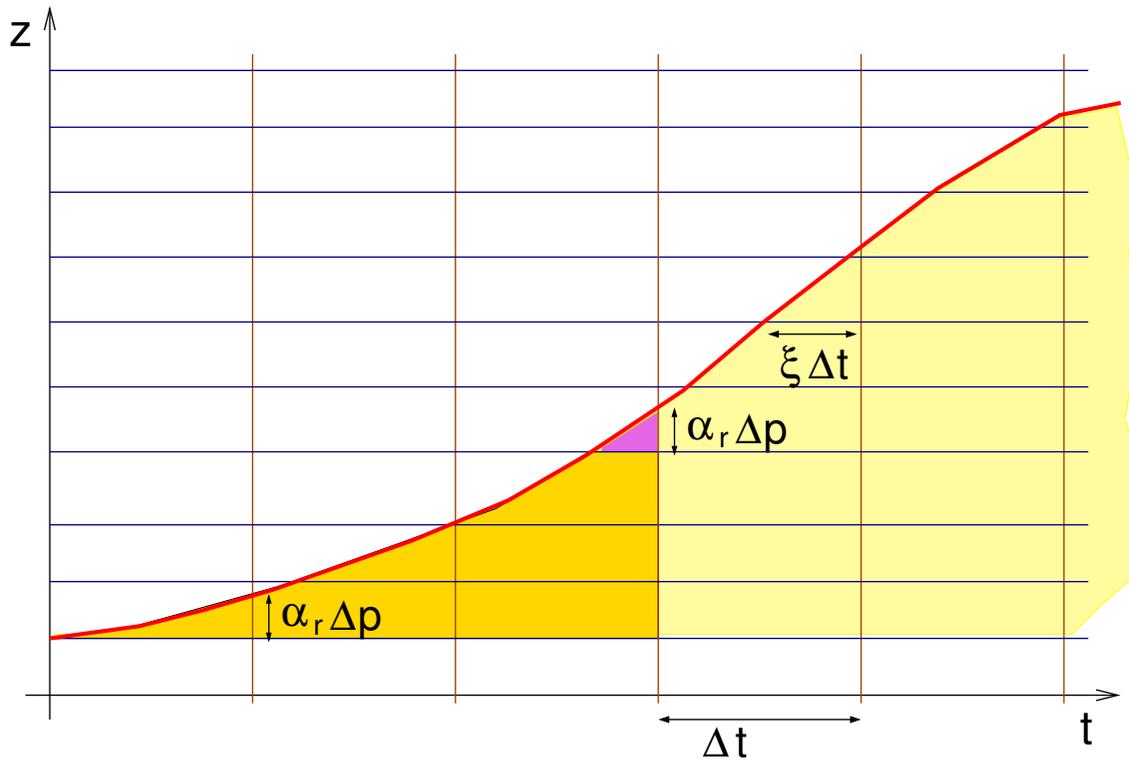


δ_{ac9} , δ_{act} record the discrete evolution of cloud vertical extension

ξ diagnosed for estimating time-averaged and final states

α_r records fractional path above upper last active level

Top evolution: activity index



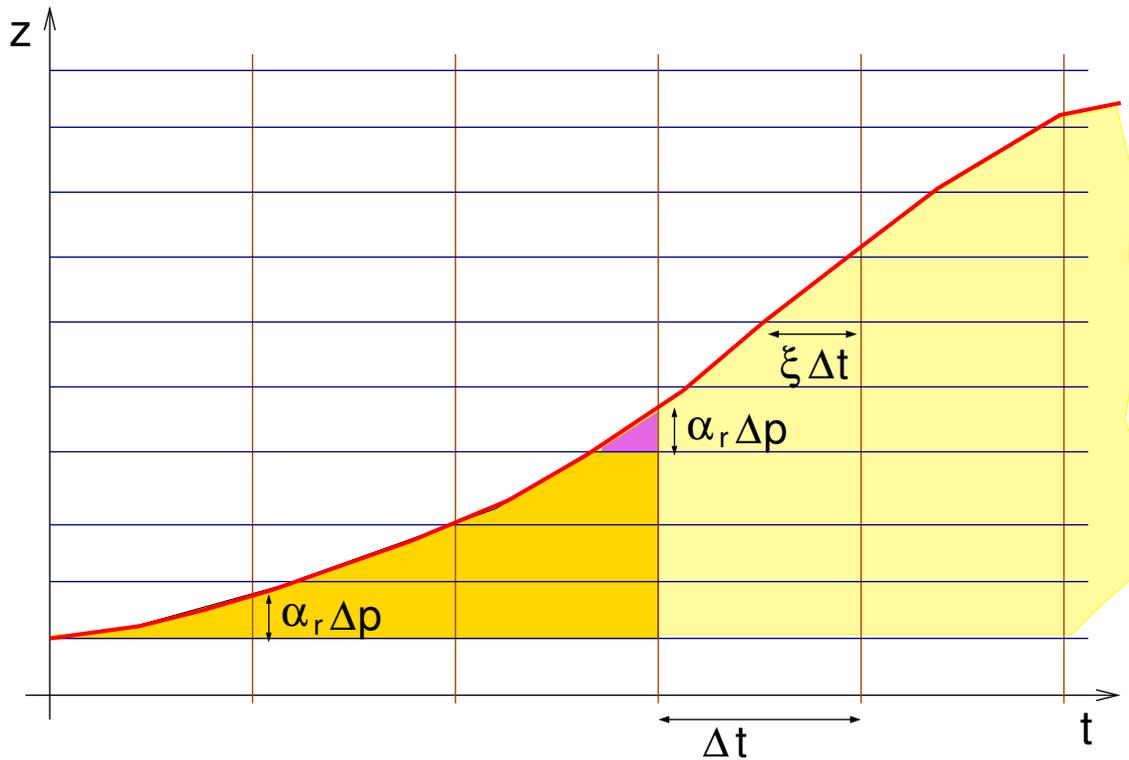
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- α_r is necessary for *initiating* an updraught with $|\omega_u|$ small;
- is necessary to compute ξ ;
- is associated to a single cloud top: top level detected in advected variables (ω_u, σ_u) , and can move its position following resolved advection.
- α_r cannot be interpolated between different columns.

Top evolution: activity index



δ_{ac9} , δ_{act} record the discrete evolution of cloud vertical extension

ξ diagnosed for estimating time-averaged and final states

α_r records fractional path above upper last active level

Idea: use a single α_r for the column, memorized in a local pseudo-historical variable:

- not advected, no interpolation;
- corresponding to the 'main' updraught segment.

Updraft evolution (continued)

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- Evolution of σ_u : relaxation towards σ_u^{\parallel}

$$\sigma_B^+ = \sigma_B^{\parallel} (1 - e^{-\Delta t/\tau}) + \sigma_B^- e^{-\Delta t/\tau}$$

$$\tau = \text{gcvtausig} \sim 300\text{s}$$

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- allows to smooth the behaviour

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- allows to smooth the behaviour
- appears better and simpler than a prognostic equation.

Updraft evolution (continued)

- Gradual elevation of the equivalent cloud top because:
 - The small grid box cannot contain a whole ensemble of clouds of different heights;
 - the time step is too short for the clouds to reach their full height in one time step.

- Evolution of σ_u : relaxation towards σ_u^{\parallel}

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- allows to smooth the behaviour
 - appears better and simpler than a prognostic equation.
- Evolution of ω_u^{\diamond} : prognostic equation, using the final σ_u^+ .

$$\text{Limit } \frac{\omega_u^{\diamond+}}{\omega_u^{\diamond}} \leq \text{gmomuss} \sim 1.5$$

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MTCS: interaction with mean flow through transport and condensation.

- Production flux $M_u^\diamond = \sigma_u(\omega_u - \bar{\omega})$
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Convective fraction $\sigma_u + \sigma_D$:

in the future, better to use a skewed distribution of moisture (Tompkins condensation scheme) and the same one to account for the intensive condensate estimation in the microphysics

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- accsu routine can be cleaned from various experimental options, final choice of relevant parameters.
- 1 additional scalar pseudo-historic field (updraft elevation between two levels).
- See triggering issues in next part.