

SURFEX
SURFEX_EKF for ALARO
SURFEX_EKF + 3DVAR
STAEKF for soil analysis

Recent activities on surface assimilation using SURFEX in the RMI

Hamdi R. with input from Annelies Duerinckx
LACE Data Assimilation Working Days, 18-20 June 2012

1. Development of a SURFEX EKF for ALARO and comparison of the offline and coupled version with the OI analysis.

2. Preliminary results on combining SURFEX EKF with 3dVar atmospheric assimilation.

3. A feasibility study of using a Short Time Augmented Extended Kalman Filter (STAEKF) for Soil Analysis.

The new land surface scheme of ALARO: SURFEX

- **Both ALADIN and ARPEGE relied on the ISBA scheme for the parameterization of the surface processes.**
(Noilhan and Planton 1989 ; Mahfouf et al. 1995 ; Noilhan and Mahfouf 1996)
- **In 2000 Valéry Masson developed a scheme for simulation the interactions with urban areas and this scheme becomes part of the meso-NH surface model.** (TEB, Masson 2000)
- **During the last decade, the surface scheme (ISBA+TEB) has been externalized from the atmospheric part of the meso-NH model following the approach of** (Best et al. 2004)
- **This led to the creation of the SURFEX scheme (SURface Externalisée) where the characteristics of the surface are specified by the ECOCLIMAP database** (Masson et al. 2003)
- **Additionally, more schemes has been added to SURFEX for: Sea and oceans (prescribed SST, ECUME, 1-D ocean model), lakes (prescribed ST, FLAKE), Surface boundary layer scheme CANOPY**
(Masson and Seity 2009; Hamdi and Masson 2008)...etc.

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STAEKF for soil analysis

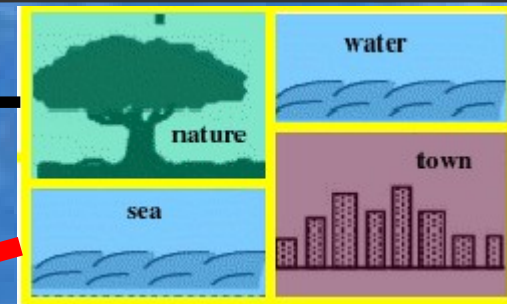
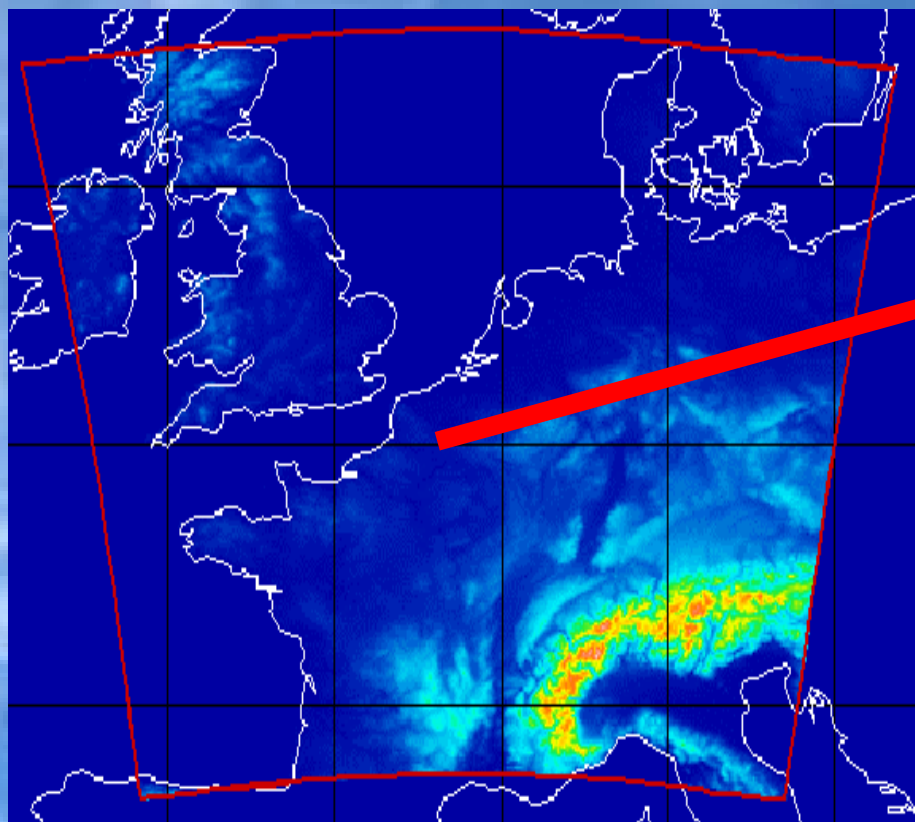
surfex

surfex with alaro

operational scores with surfex

Tiling

One important feature of the externalized surface:
each grid cell is divided into 4 elementary units called tiles according to the
fraction of covers in the grid cell



- 1: bare ground
- 2: rocks
- 3: permanent snow
- 4: deciduous forest
- 5: conifer forest
- 6: evergreen broadleaf trees
- 7: C3 crops
- 8: C4 crops
- 9: irrigated crops
- 10: grassland
- 11: tropical grassland
- 12: garden and parks

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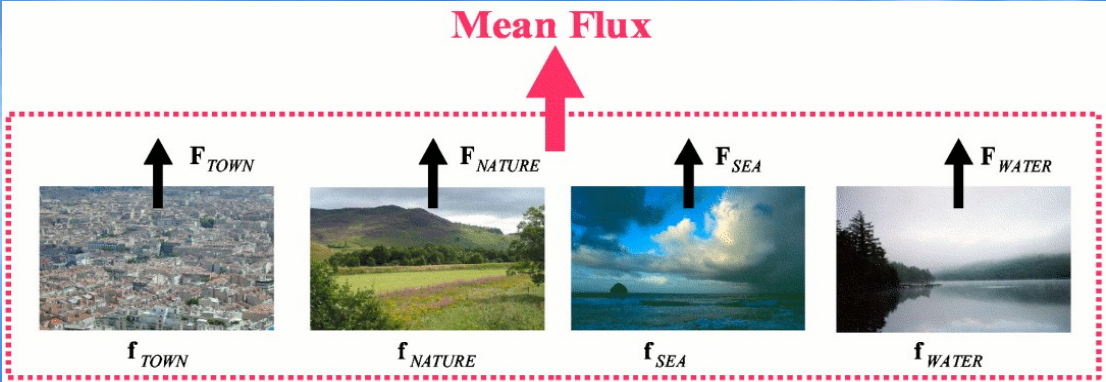
operational scores with surfex

Surfex output as surface boundary conditions for atmospheric radiation and turbulent scheme.

- albedo
- emissivity
- radiative temperature
- momentum flux
- sensible heat flux
- latent heat flux
- CO₂ flux
- chemical flux

ALARO model

- Atmospheric forcing
- Sun position
- Downward radiative flux



ALARO+FMR+SURFEX Vs ALARO+ACRANEB run over Belgium

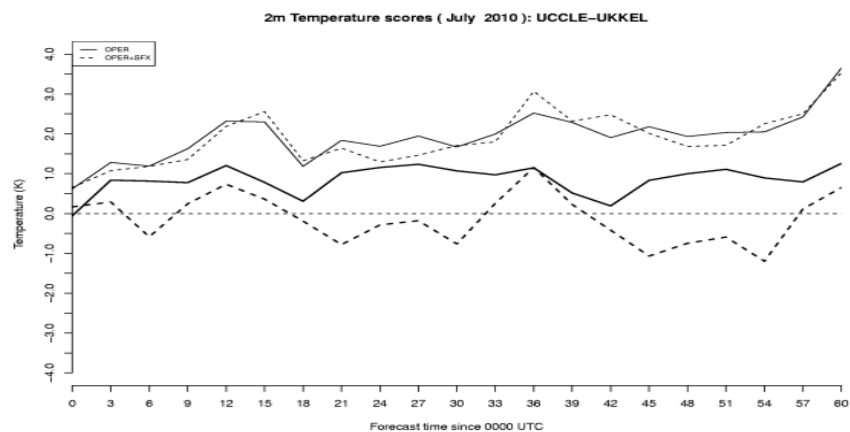
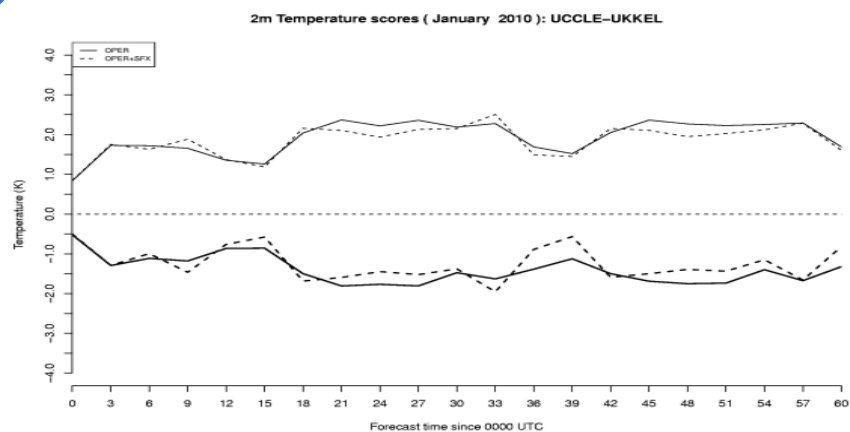


Table 2. the average daytime/nighttime scores for the flat/high topography and coastal synoptic stations, sign (+) means improvement, sign (0) means neutral effect, and sign (-) means degradation of the scores.

	<i>WinterNight</i>	<i>WinterDay</i>	<i>SummerNight</i>	<i>SummerDay</i>
Temperature at 2m				
Flat	+	0	0	0
High	0	-	0	0
Coast	0	0	+	0
Wind speed at 10m				
Flat	0	0	0	0
High	0	0	0	0
Coast	0	0	0	0
Wind direction at 10m				
Flat	0	0	0	0
High	0	0	0	0
Coast	0	0	0	0
Relative humidity at 2m				
Flat	+	+	0	0

- Optimum Interpolation (OI)
 - Coefficients are derived using simple assumptions
- Extended Kalman Filter (EKF)
 - dynamical coefficients
 - Formulation: $\mathbf{x}_t^a = \mathbf{x}_t^b + \mathbf{B}\mathbf{H}^T (\mathbf{H}\mathbf{B}\mathbf{H}^T + \mathbf{R})^{-1} [\mathbf{y}_t^o - \mathcal{H}(\mathbf{x}_0^b)]$
 - \mathcal{H} : the observation operator
 - includes a model propagation
 - \mathbf{H} : the Jacobian matrix of the observation operator:

$$\mathbf{H} = \frac{\partial \mathbf{y}_t}{\partial \mathbf{x}_{t_0}}$$

\mathbf{H} is calculated by a finite difference approach:

$$H_{ij} = \frac{y_i(\mathbf{x} + \delta x_j) - y_i(\mathbf{x})}{\delta x_j}$$

- Finite difference approach:

$$H_{ij} = \frac{y_i(\mathbf{x} + \delta x_j) - y_i(\mathbf{x})}{\delta x_j}$$

- Calculation of $y_i(\mathbf{x} + \delta x_j)$ and $y_i(\mathbf{x})$:
 - Perturb a component x_j of the control vector \mathbf{x}
 - Perform run with the perturbed surface field to calculate corresponding y_i
 - Offline: surface scheme decoupled from atmospheric model
the forcing is taken from the lowest model level
 - Coupled: surface scheme coupled to atmospheric model
interaction is possible between atmosphere and surface

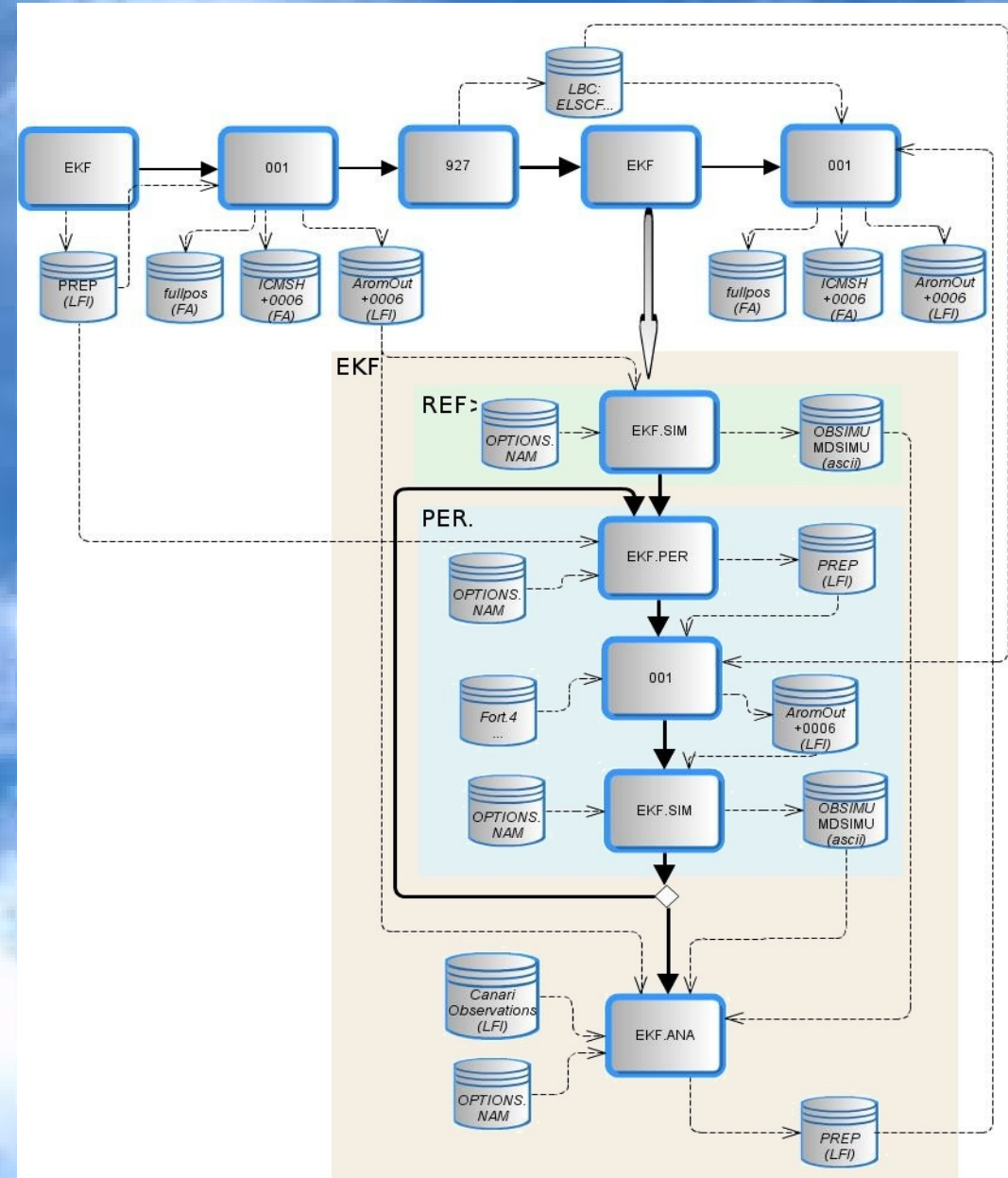
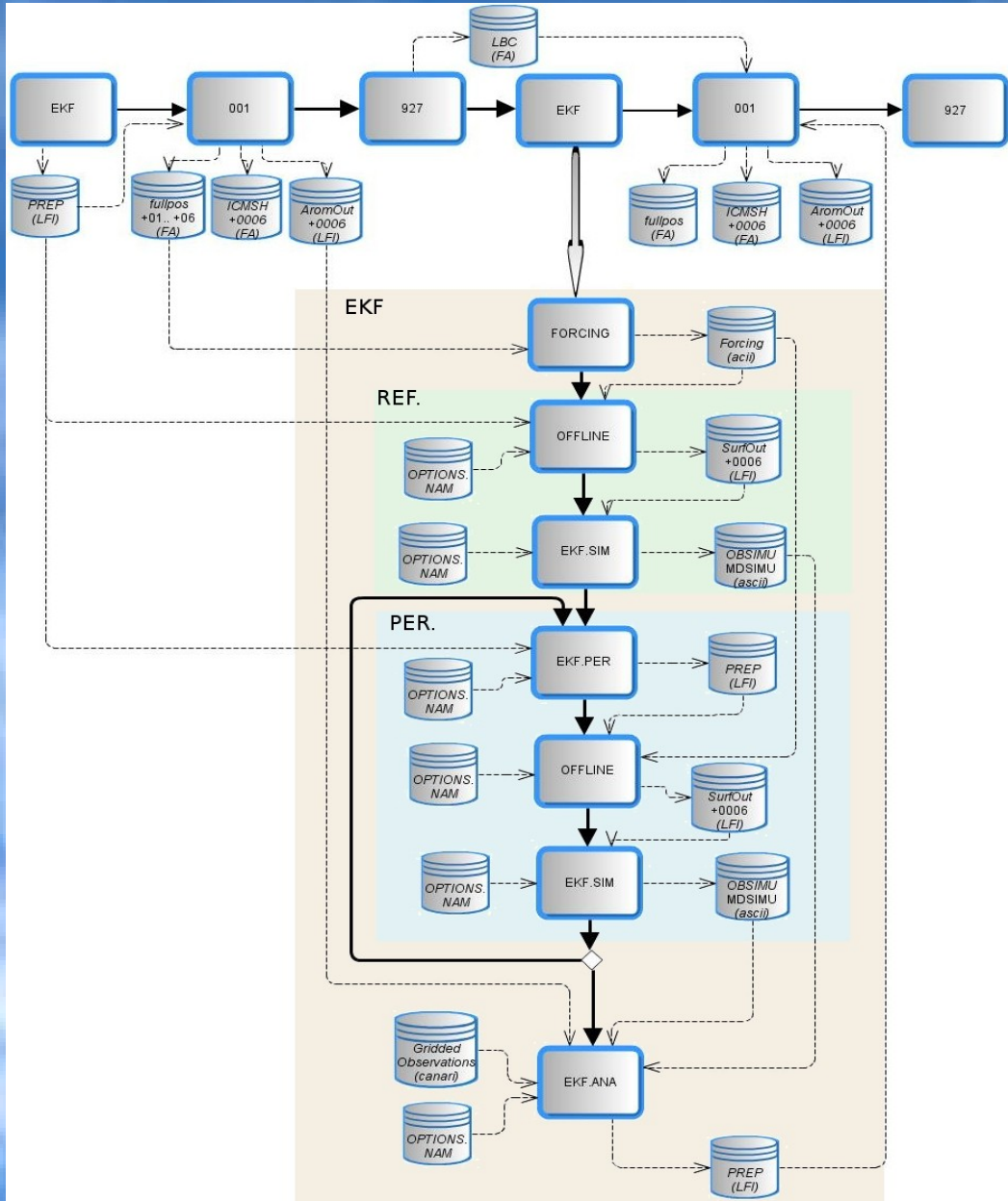
- In Balsamo et al. (2007,JH), the offline approach is compared to the coupled approach for a single day during the summer with the GEM model. The gain components are found to be smaller for the offline approach but with similar patterns as the coupled approach.
- In Balsamo et al (2004,QJRMS) it is mentioned that for coupled experiments the choice of the perturbation size is important. Small perturbations can lead to a noisy H matrix and inaccurate corrections.
- This presentation:
 - Implementation of offline and coupled SURFEX EKF for ALADIN
 - Comparison between offline and coupled EKF: jacobians, gain, increments
 - Comparison of forecast scores of EKF, OI and runs without assimilation

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Surface data assimilation
 Extended Kalman Filter
 Offline VS. Coupled EKF
 Experimental Set-up

Jacobian
 Increments (WG)
 Time evolution (WG2)
 Scores



- Prognostic variables:

- Superficial water content (wg1)
- Root zone water content (wg2)
- Surface temperature (Tg1)
- Deep soil temperature (Tg2)

- Observations:

- Screen level temperature (T2m)
- Screen level relative humidity (RH2m)

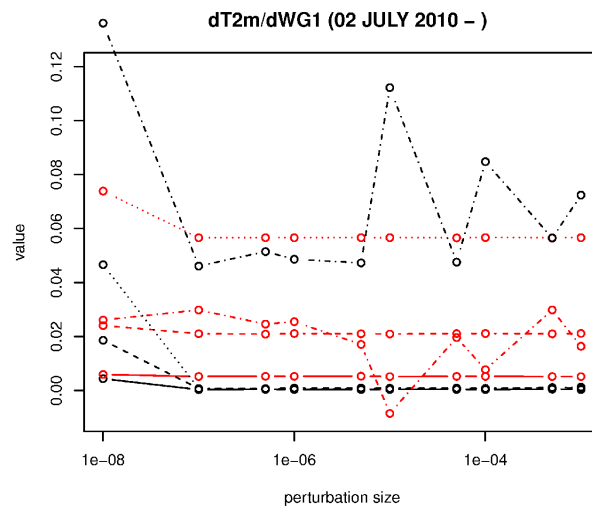
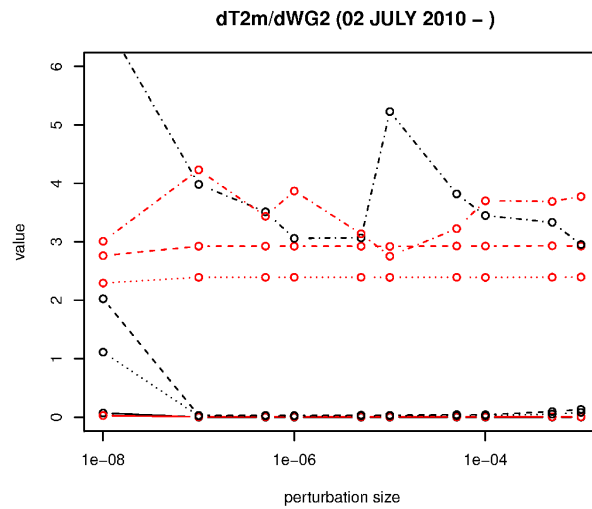
- Additional information:

- LBC data from Aladin France
- Assimilation interval $\tau = 6$ hours, with assimilation at 00,06,12,18 UTC
- Forecasts with surfex + alaro and inline fullpos (interval 1 hour)
- Background error covariance matrix **B** is kept constant (WG set to 0.1, TG set to 2 K)
- Error covariance matrix **R**: T2m set to 1K, RH2m set to 10%

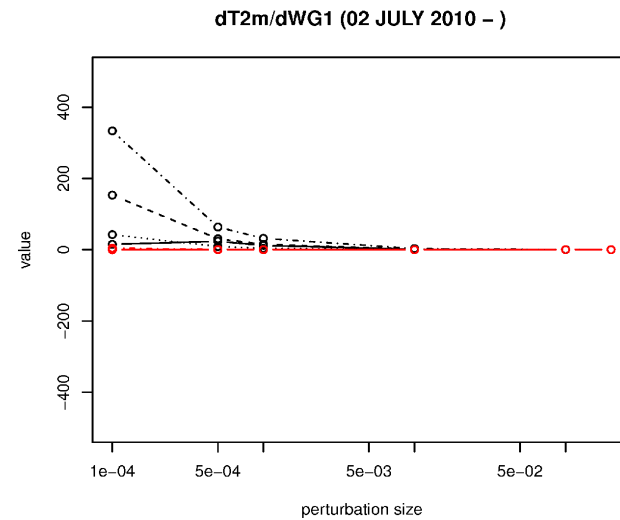
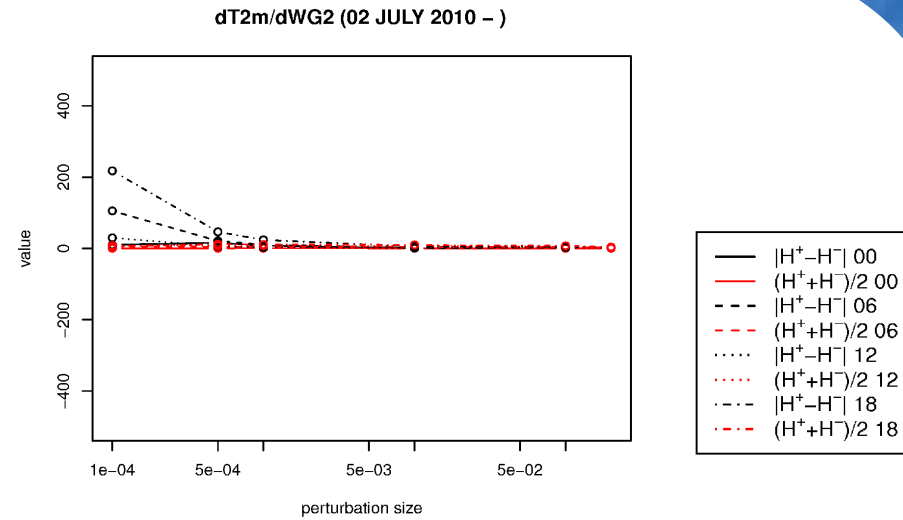
Experiments:

- Free run: (no assimilation)
 - Surface field from 6h forecast of previous run
- Open loop: (no assimilation)
 - Surface is interpolated from Arpege analysis
- Surface Assimilation runs:
 - Optimum Interpolation
 - EKF with offline jacobian calculation
 - EKF with coupled jacobian calculation
 - Surface guess for assimilation is taken from 6h forecast of previous run

EKF offline



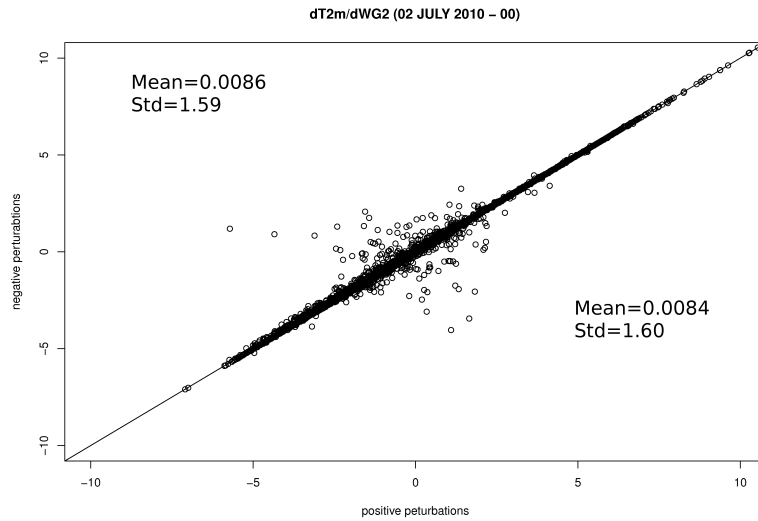
EKF Coupled



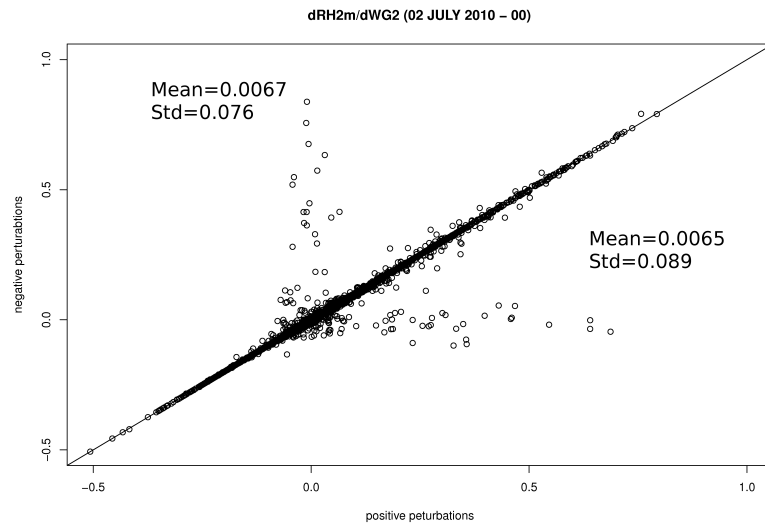
EKF offline

EKF Coupled

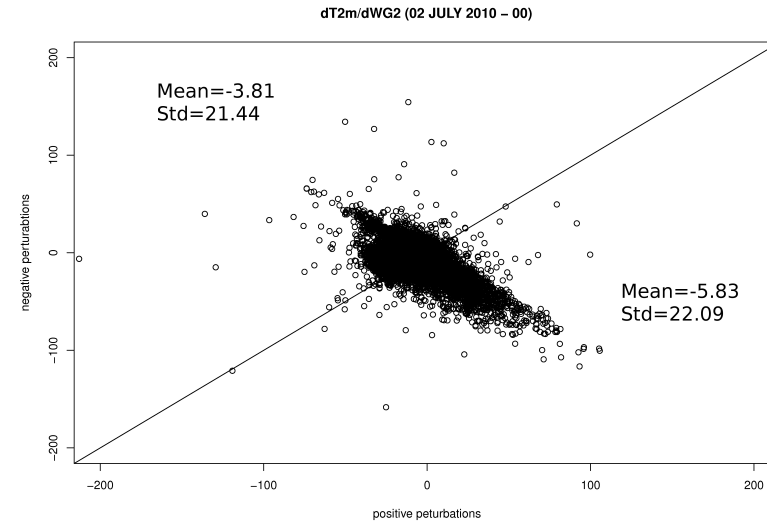
Pert. size:
 10^{-4}



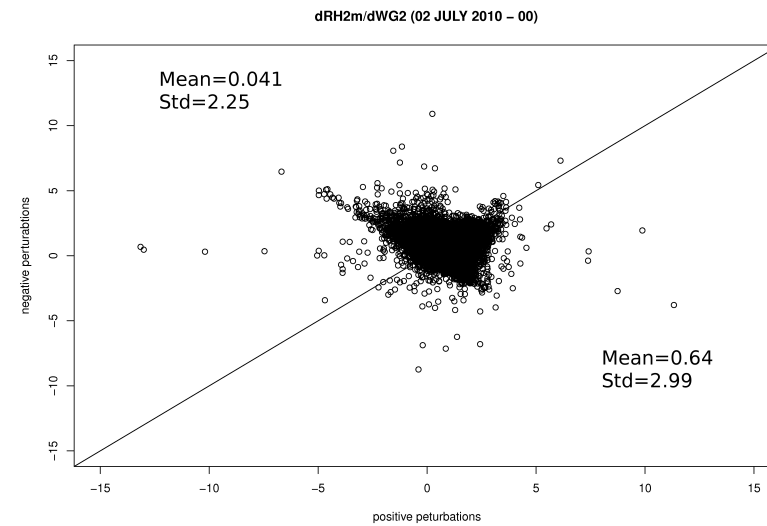
Pert. size:
 10^{-4}



Pert. size:
 10^{-1}



Pert. size:
 10^{-1}



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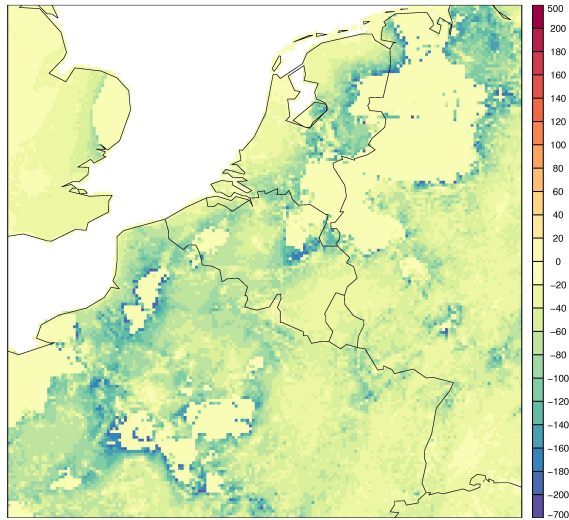
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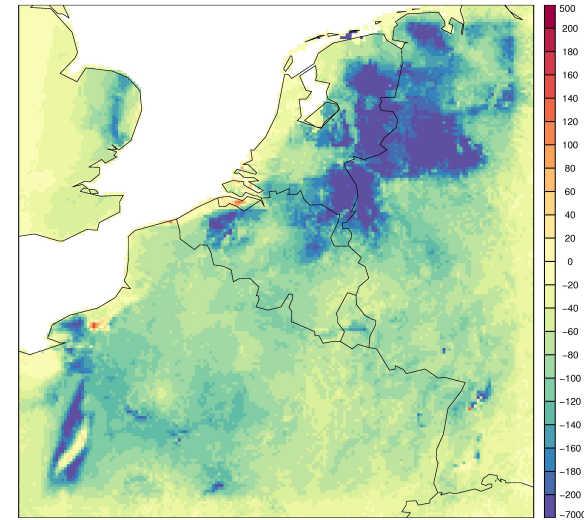
EKF offline

EKF coupled

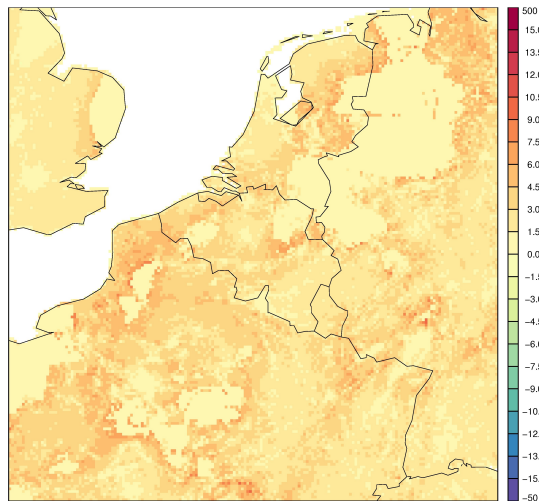
T2m-WG2 jacobians for 02 July 2010 1200



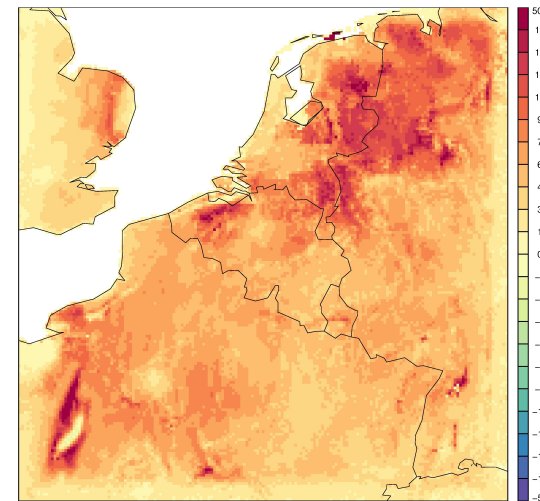
T2m-WG2 jacobians for 02 July 2010 1200



RH2m-WG2 jacobians for 02 July 2010 1200



RH2m-WG2 jacobians for 02 July 2010 1200



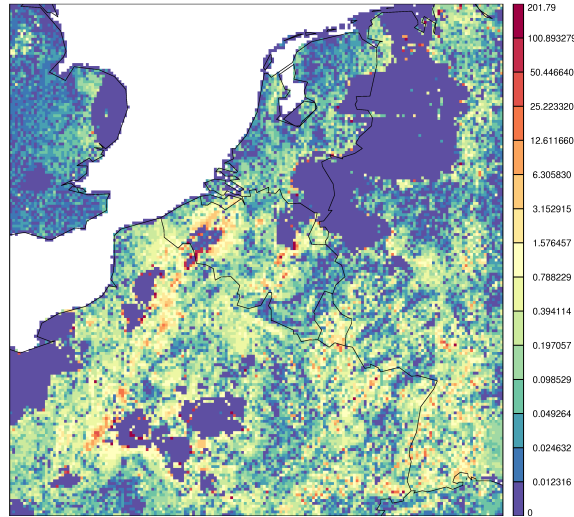
T2m-
WG2

RH2m-
WG2

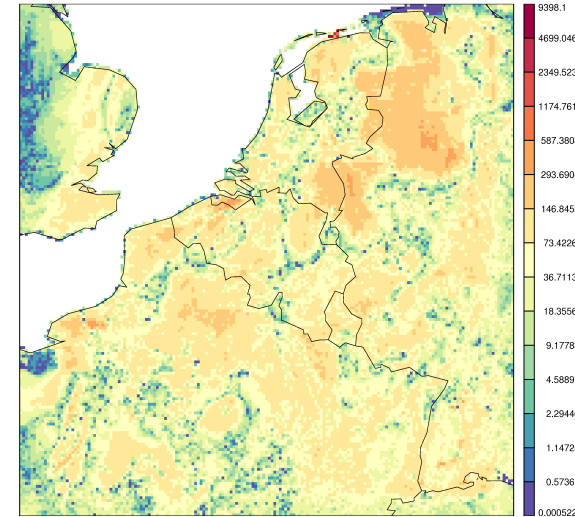
EKF offline

EKF coupled

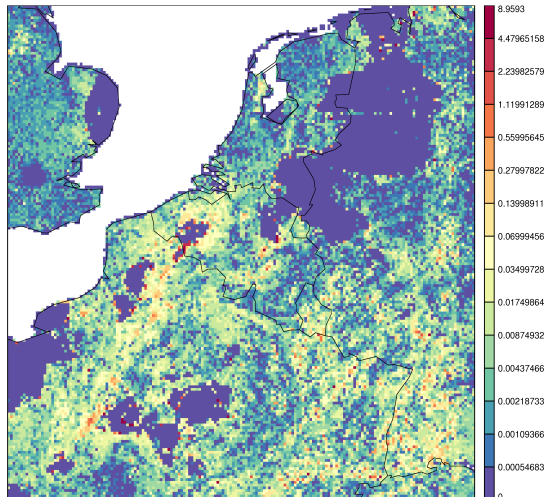
$|H+ - H-|$ for dT2m/dWG2 02 July 2010 1200



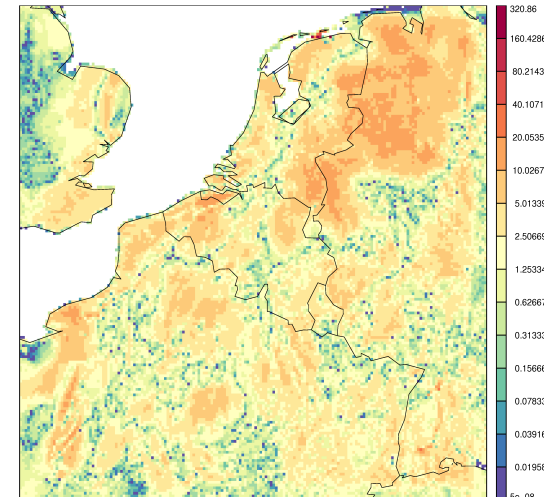
$|H+ - H-|$ for dT2m/dWG2 02 July 2010 1200



$|H+ - H-|$ for dRH2m/dWG2 02 July 2010 1200



$|H+ - H-|$ for dRH2m/dWG2 02 July 2010 1200



T2m-
WG2

RH2m-
WG2

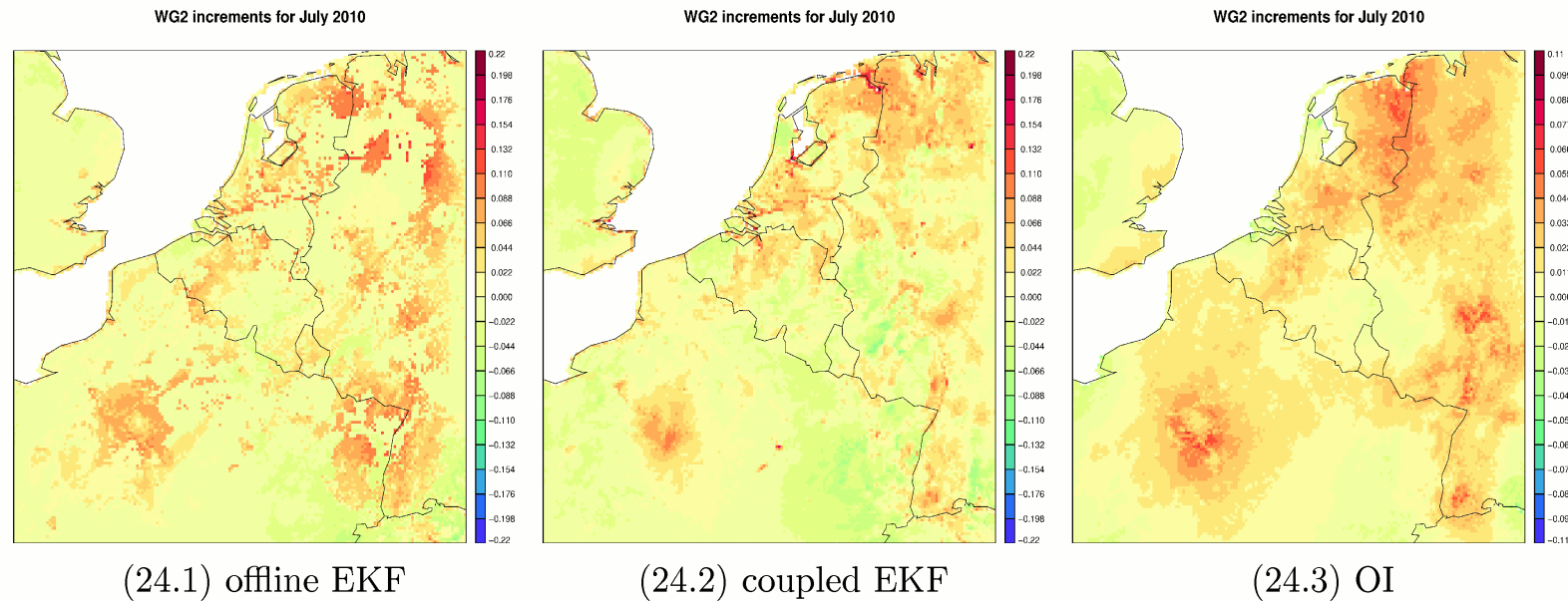


Figure 24: Cumulative increments for July 2010, WG2. The incrementscale for OI is two times smaller than for both EKF.

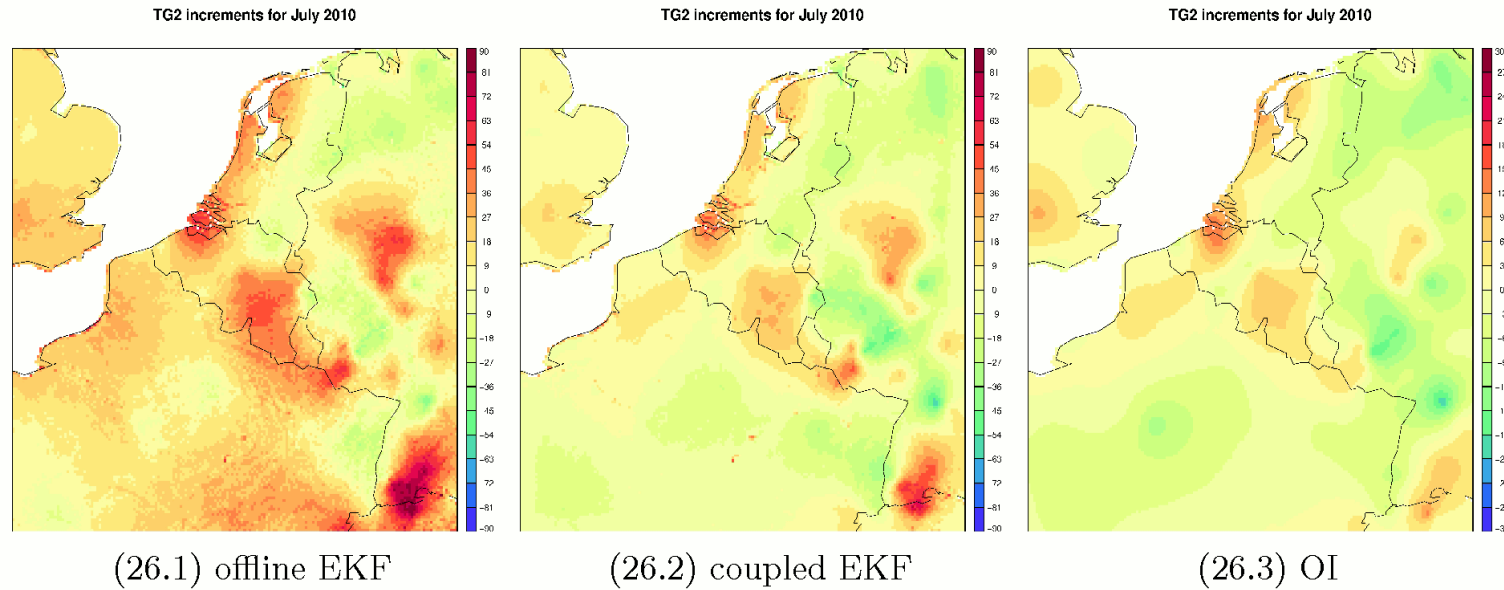
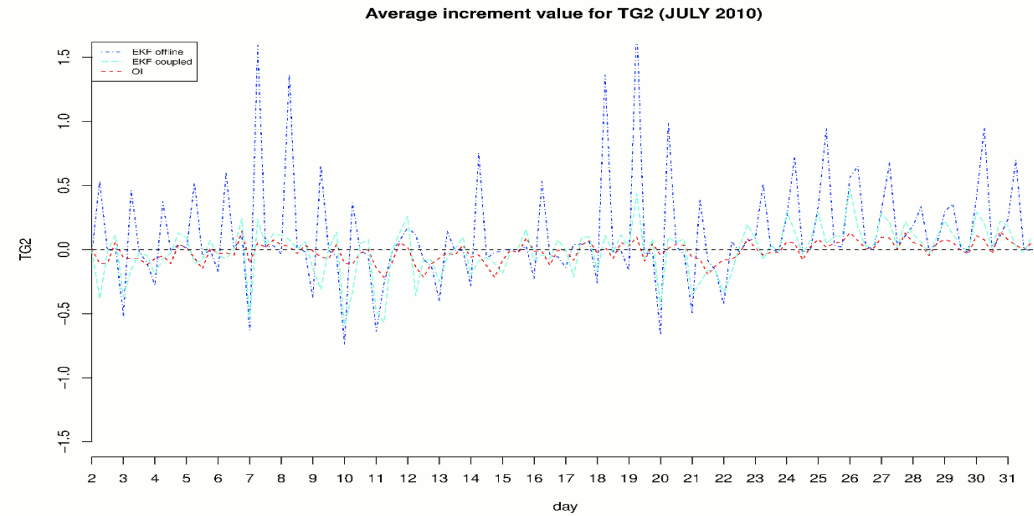
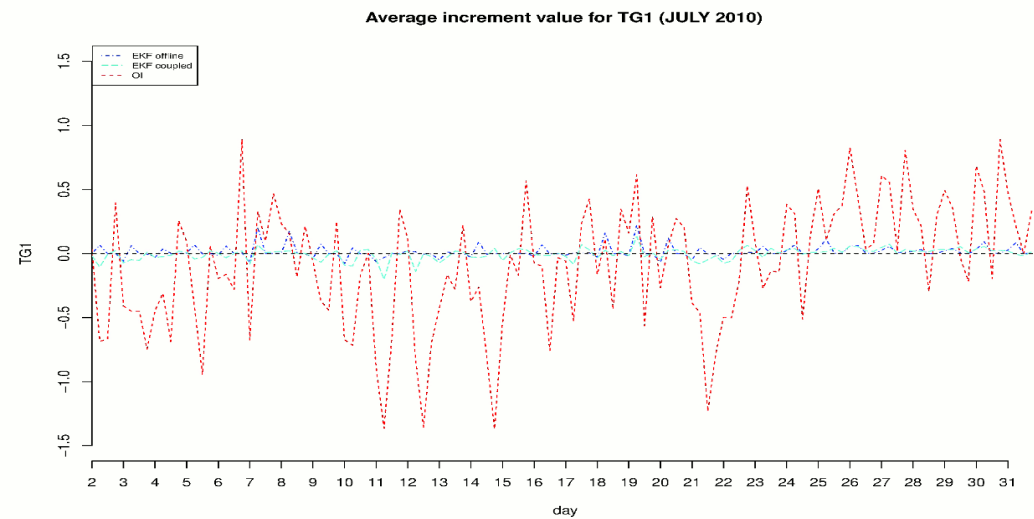


Figure 26: Cummulative increments for July 2010, TG2. The incrementscale for OI is three times smaller than for both EKF.

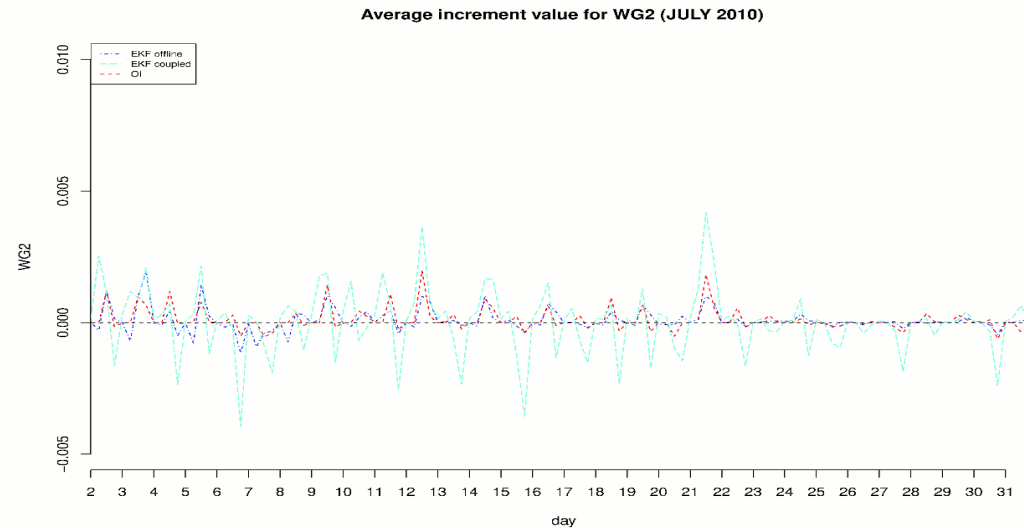


(28.1) Evolution of the TG2 increments averaged over the domain for July 2010

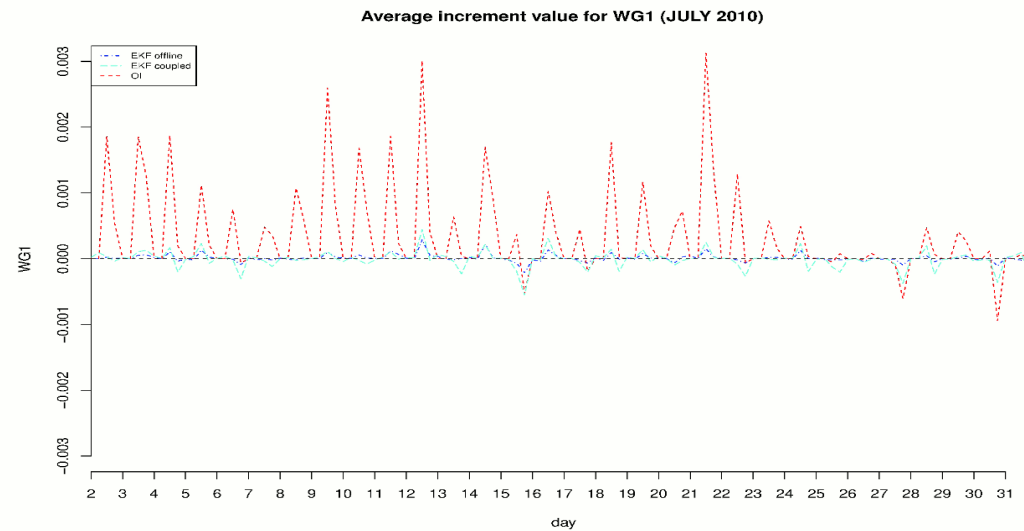


(28.2) Evolution of the TG1 increments averaged over the domain for July 2010

Figure 28



(28.3) Evolution of the WG2 increments averaged over the domain for July 2010



(28.4) Evolution of the WG1 increments averaged over the domain for July 2010

Figure 28: Evolution of the increments of the soil prognostic variables

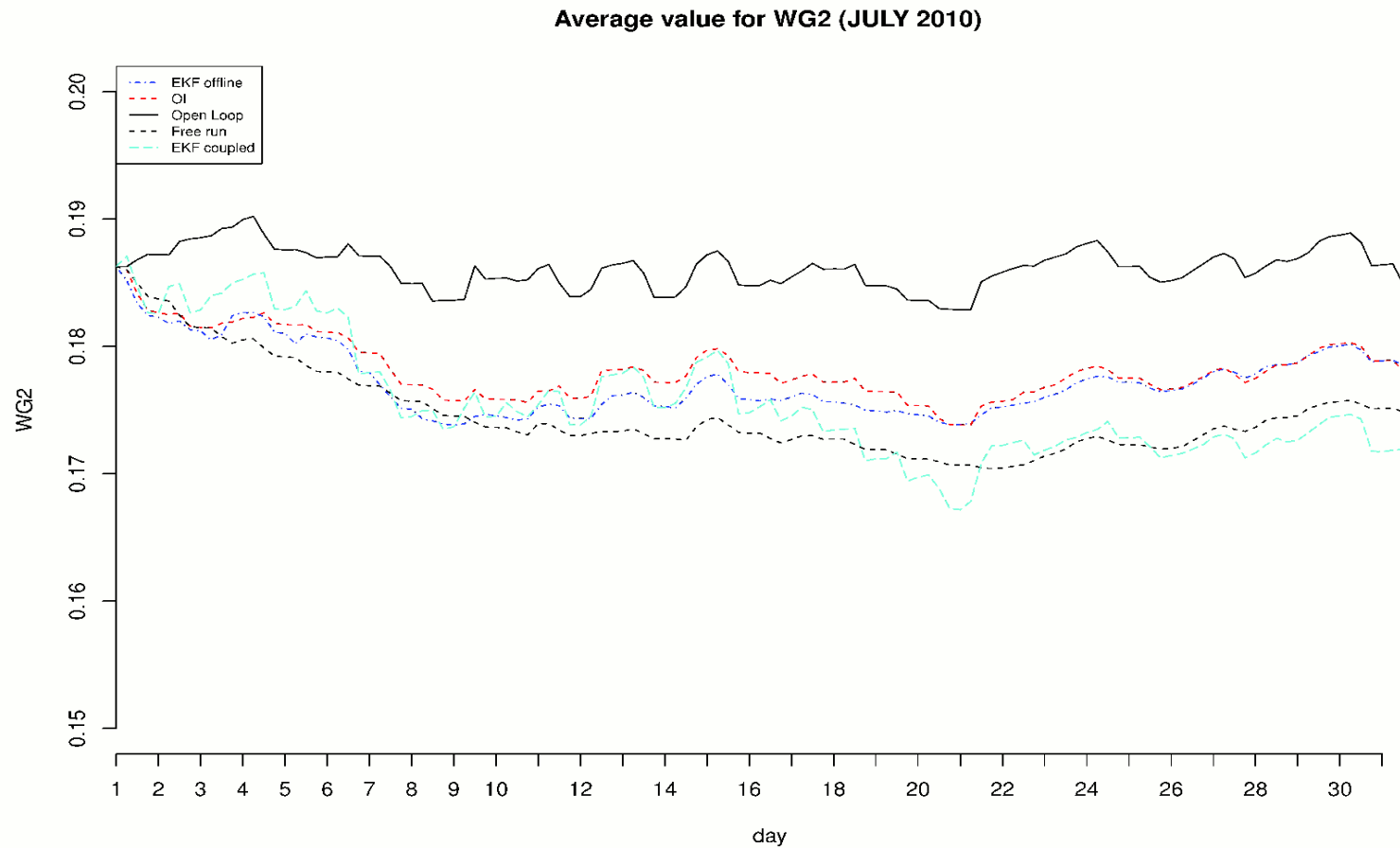
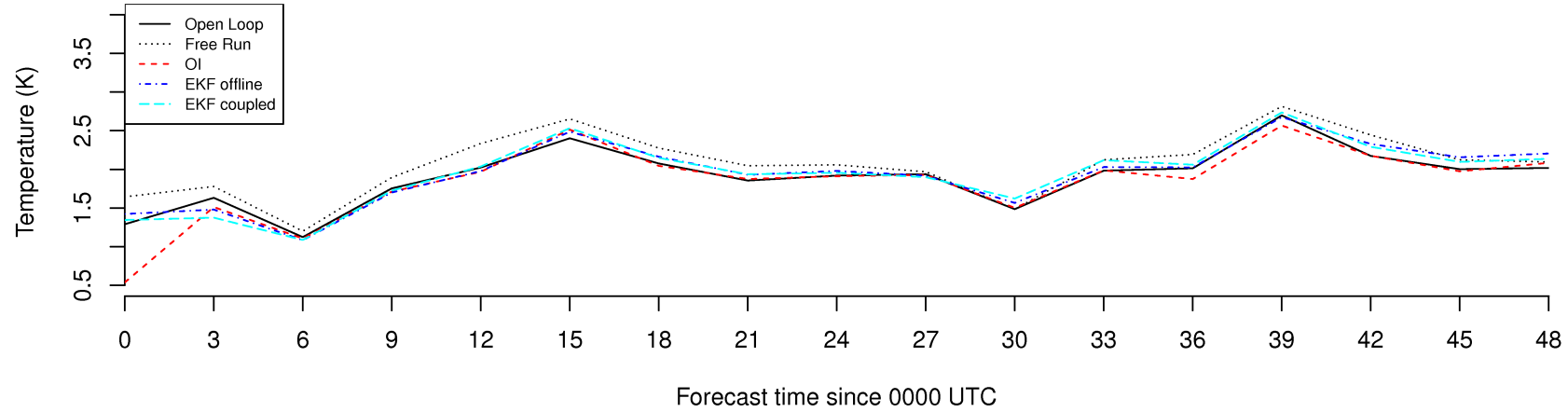
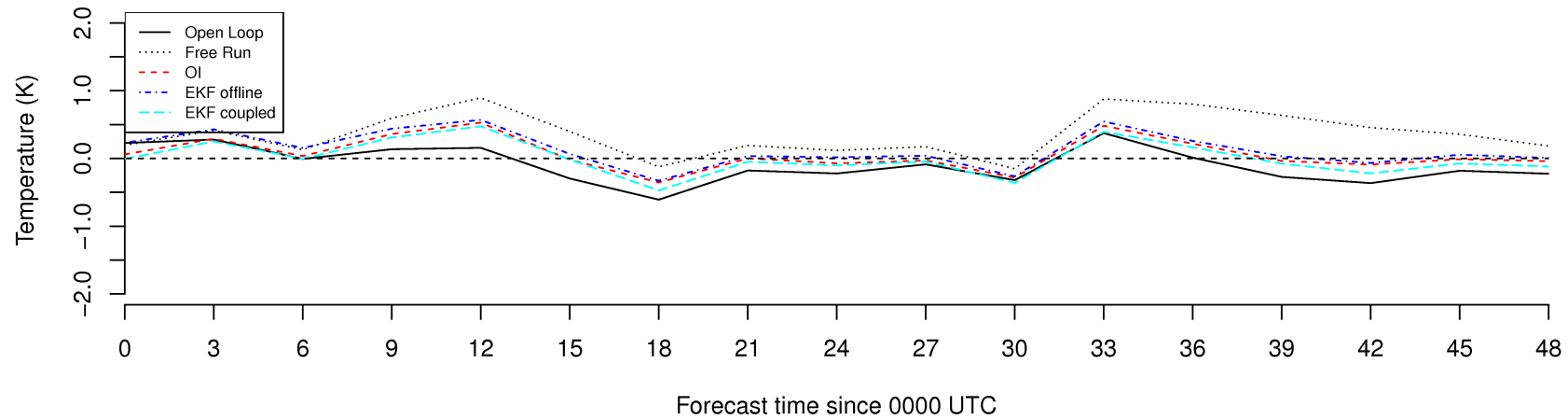


Figure 32: Evolution of the WG2 values averaged over the domain for July 2010

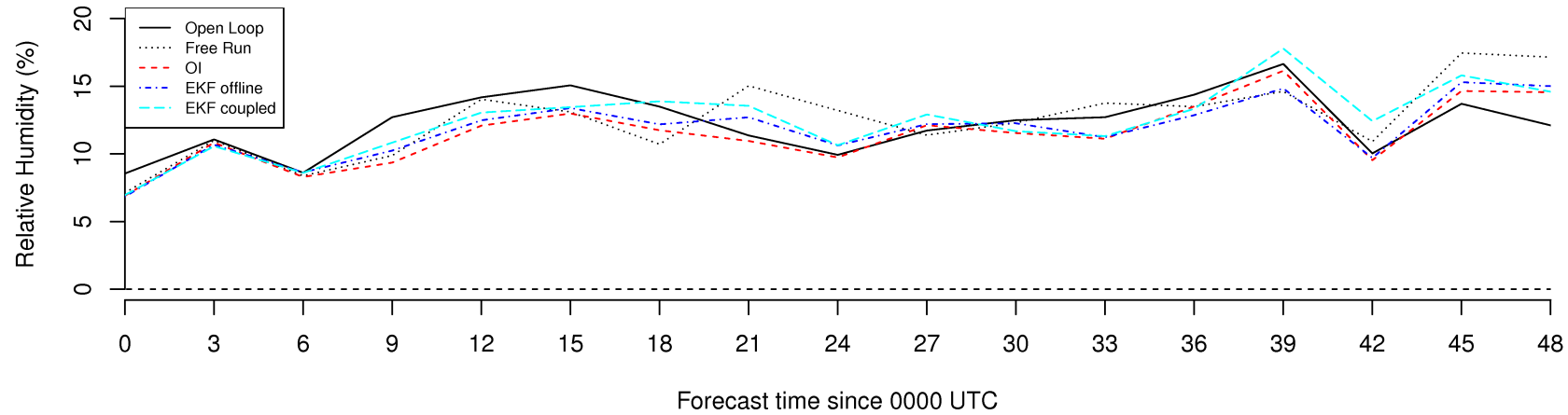
2m Temperature RMSE (JULY 2010) run 0



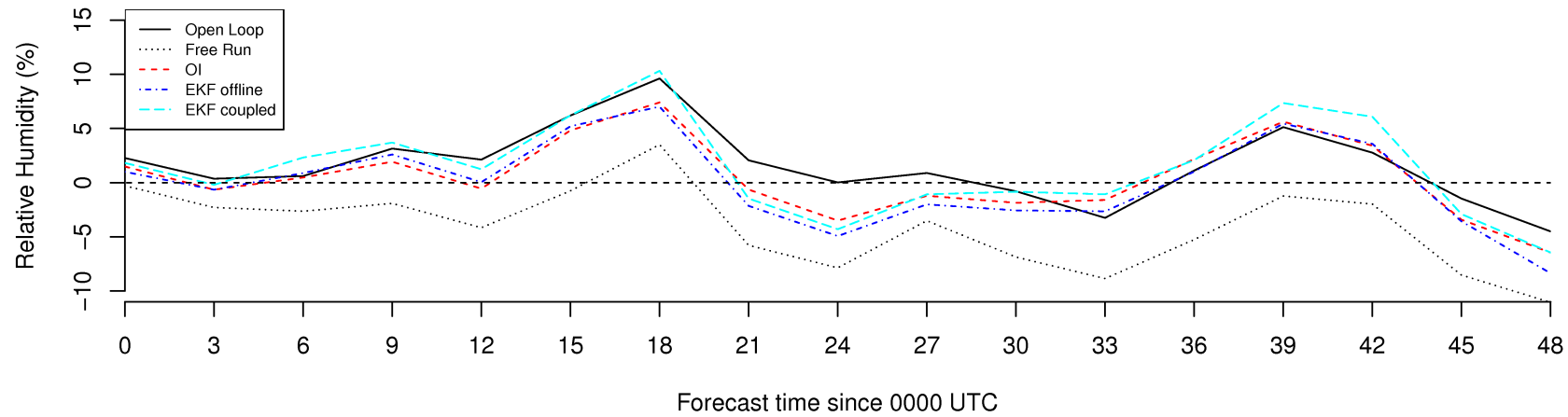
2m Temperature BIAS (July 2010) run 0



2m Relative Humidity RMSE (July 2010): UCCLE-UKKEL



2m Relative Humidity BIAS (July 2010): UCCLE-UKKEL



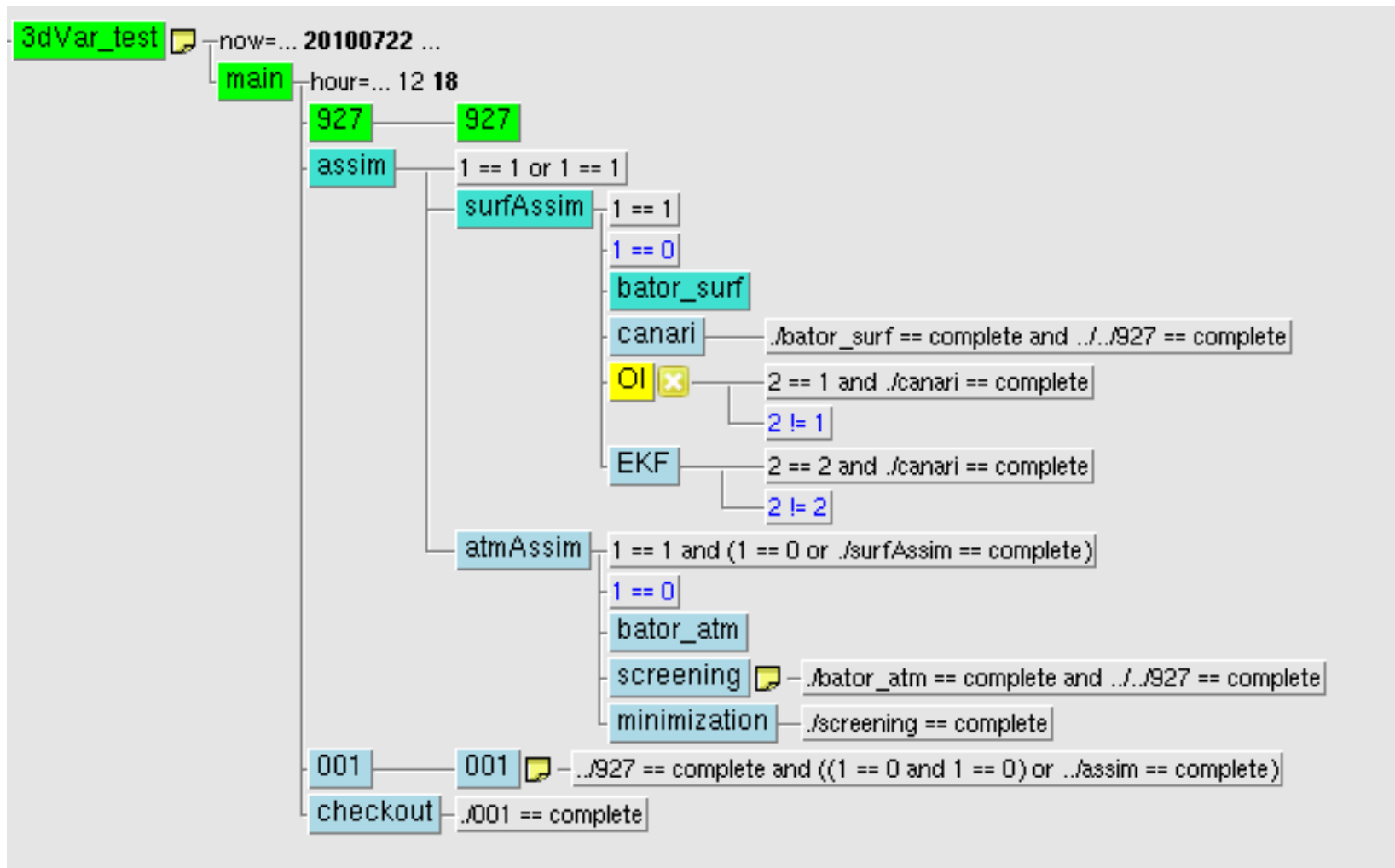
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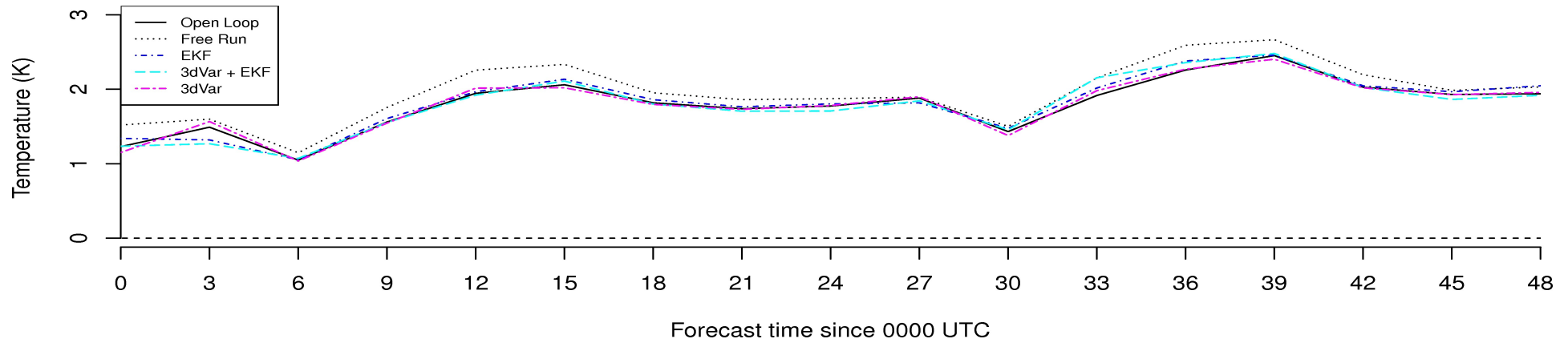
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Experiments:

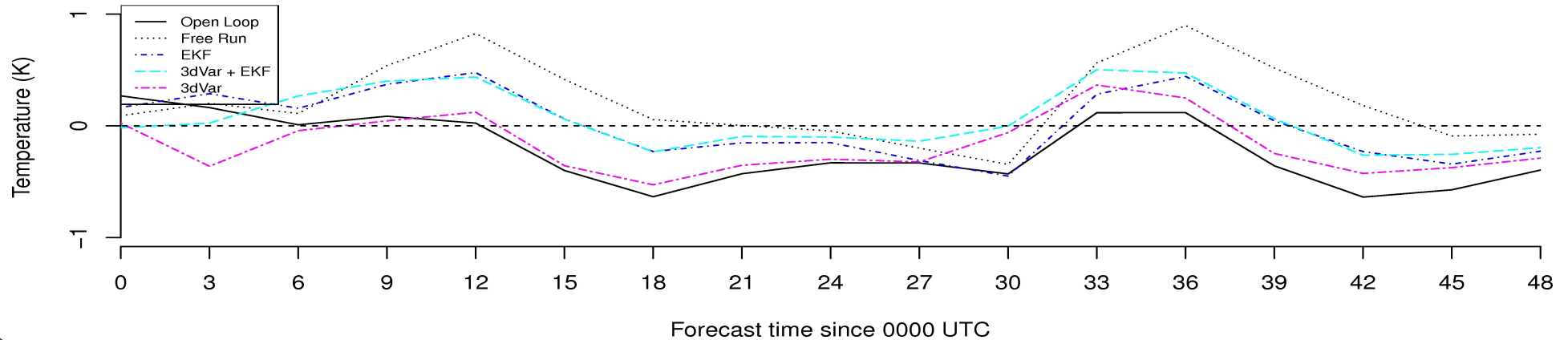
- **Free run:** (no assimilation)
Atmospheric fields from ARPEGE analysis. Surface field from 6h forecast of previous run
- **Open loop:** (no assimilation)
Atmospheric fields and surface fields are interpolated from Arpege analysis
- **EKF** with surface assimilation.
- **3DVAR** With atmospheric assimilation using only conventional observations (no satellite, no radar)
- **3DVAR+EKF** surface assimilation is done before atmospheric assimilation.



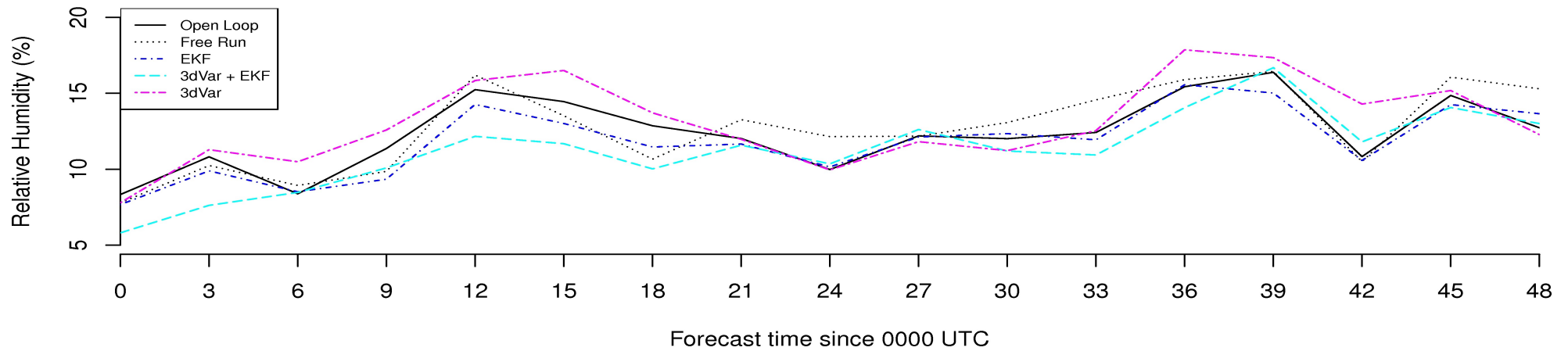
2m Temperature RMSE (01–31 July 2010) run 0



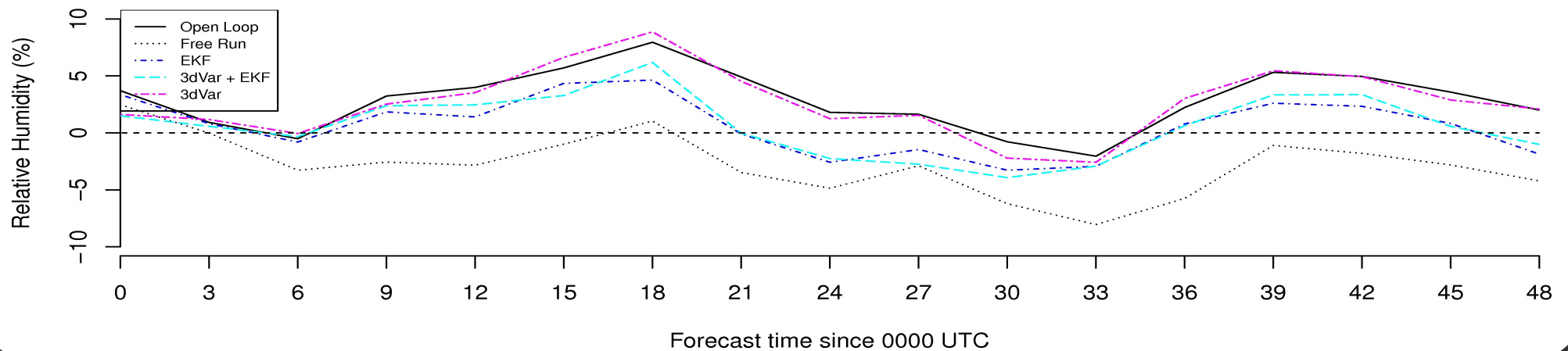
2m Temperature BIAS (01–31 July 2010) run 0



2m Relative Humidity RMSE (01–31 July 2010) run 0



2m Relative Humidity BIAS (01–31 July 2010) run 0



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- Recently, Carrassi and Vannitsem (2011, QJRMS) introduced an alternative formulation of the EKF where the uncertain model parameters are estimated along with the system state variables.
- The algorithm, Short Time Augmented Extended Kalman Filter (STAEOF), uses a deterministic formulation for the model error dynamics (Nicolis, 2003, JAS).
- The same formulation has been used for the treatment of the error arising from the unresolved scales (Carrassi and Vannitsem, 2011, IJBC) and in the context of variational assimilation (Carrassi and Vannitsem, 2010, MWR).
- We undertake here a set of numerical twin experiments designed to test the STAEOF in estimating three land surface parameters: LAI, the albedo, and the minimum stomatal resistance RS_{min} .
- Assimilation of 2m temperature and relative humidity using an offline version of ISBA.

- **The two-layers version of the land surface model ISBA.**

The model equations read:

$$\frac{\partial T_s}{\partial t} = C_T(R_n - H - LE) - \frac{2\pi}{\tau}(T_s - T_2)$$

$$\frac{\partial T_2}{\partial t} = \frac{1}{\tau}(T_s - T_2)$$

$$\frac{\partial w_g}{\partial t} = \frac{C_1}{\rho_w d_1}(P_g - E_g) - \frac{C_2}{\tau}(w_g - w_{geq})$$

$$\frac{\partial w_2}{\partial t} = \frac{1}{\rho_w d_2}(P_g - E_g - E_{tr}) - \frac{C_3}{d_2 \tau} \max[0., (w_2 - w_{fc})]$$

- **The model is available within a surface externalized platform (SLDAS, Mahfouf 2007).**
- **The state vector, $X = (T_s, T_2, W_g, W_2)$ and the equation can be formally written as a dynamical system, $\frac{dx}{dt} = g(X, \lambda)$**
- **The vector λ is taken to represent the set of model parameters.**

The forecast model, is augmented with P model parameters:

$$\mathbf{z}^f = \begin{bmatrix} \mathbf{x}^f \\ \lambda^f \end{bmatrix} = \mathcal{F}\mathbf{z}^a = \begin{bmatrix} \mathcal{M}\mathbf{x}^a \\ \mathcal{F}^\lambda\lambda^a \end{bmatrix}, \quad (1)$$

$\mathbf{z} = (\mathbf{x}, \lambda)$ is the augmented state vector. The augmented dynamical system \mathcal{F} includes the dynamical model for the system's state, \mathcal{M} , and a dynamical model for the parameters \mathcal{F}^λ . In the absence of additional information, a persistence model for \mathcal{F}^λ is often assumed so that $\mathcal{F}^\lambda = \mathbf{I}$ and $\lambda_{t_{k+1}}^f = \lambda_{t_k}^a$; the same choice has been adopted here.

The forecast/analysis error covariance matrix, $\mathbf{P}_z^{f,a}$, for the augmented system reads:

$$\mathbf{P}_z^{f,a} = \begin{pmatrix} \mathbf{P}_x^{f,a} & \mathbf{P}_{x\lambda}^{f,a} \\ \mathbf{P}_{x\lambda}^{f,aT} & \mathbf{P}_\lambda^{f,a} \end{pmatrix} \quad (2)$$

where the $I \times I$ matrix $\mathbf{P}_x^{f,a}$ is the error covariance of the state estimate $\mathbf{x}^{f,a}$, $\mathbf{P}_\lambda^{f,a}$ is the

$P \times P$ parametric error covariance and $\mathbf{P}_{x\lambda}^{f,a}$ the $I \times P$ error correlation matrix between the

state vector, \mathbf{x} , and the vector of parameters λ . These correlations are essential for the

estimation of the parameters. In general one does not have access to a direct measurement

of the parameters, and information are only obtained through observations of the system's

state.

The forecast error propagation in the STAEKF is given by $\mathbf{P}_z^f = \mathbf{C}\mathbf{P}_z^a\mathbf{C}^T$, with \mathbf{C} being the STAEKF forward operator defined as:

$$\mathbf{C} = \begin{pmatrix} \mathbf{M} & \frac{\partial g}{\partial \lambda} |_{\lambda^a \tau} \\ 0 & \mathbf{I}_P \end{pmatrix} \quad (3)$$

The short-time truncation of the dynamics

where \mathbf{I}_P is the $P \times P$ identity matrix. Equation (3) embeds the key feature of the

STAEKF; the presence of the term $\frac{\partial g}{\partial \lambda} |_{\lambda^a \tau}$ allows for accounting for the contribution of

the parametric error to the forecast error as well as to the error correlation between model

state and parameters.

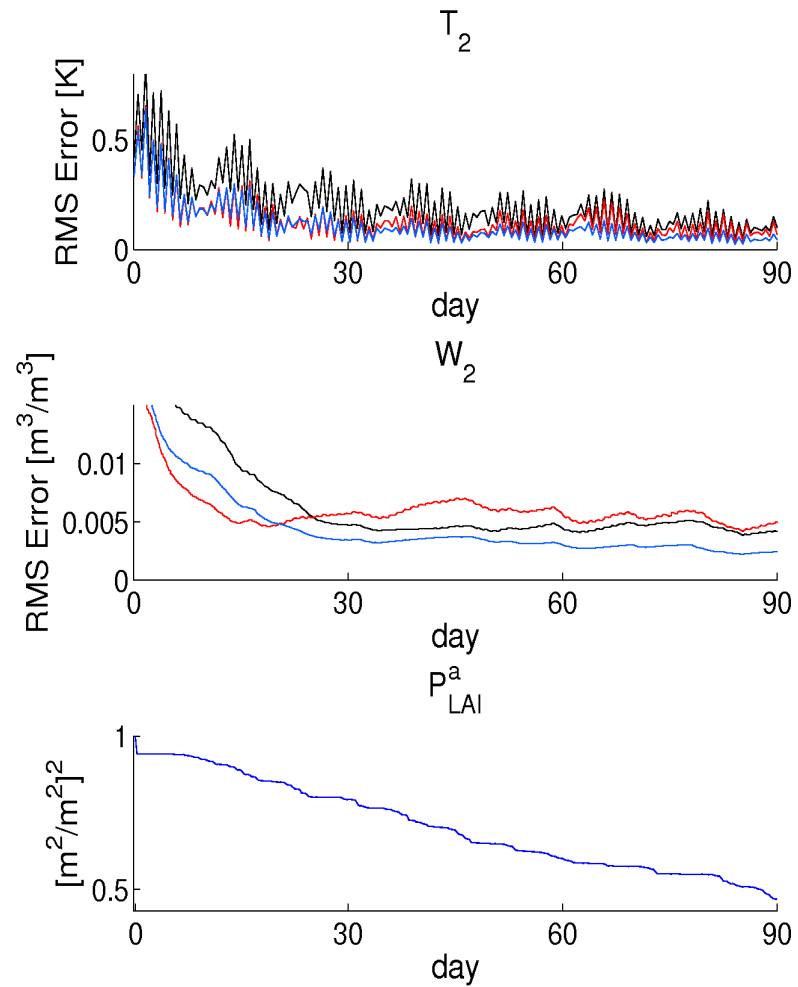
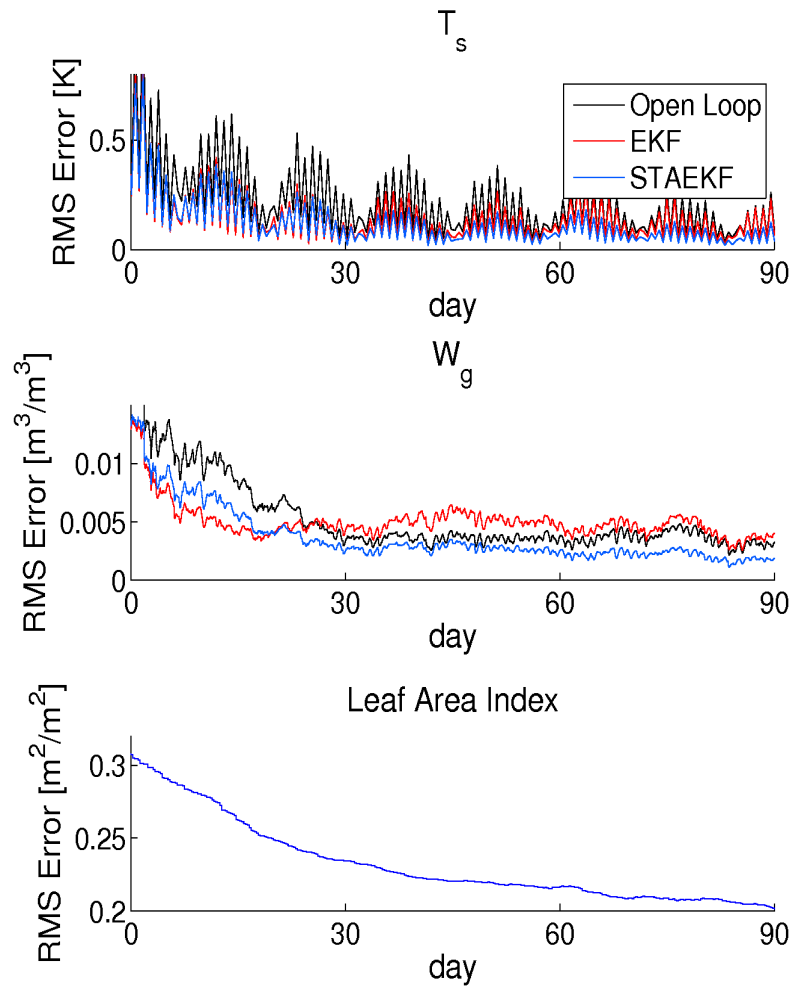
An augmented observation operator is introduced, $\mathcal{H}_z = [\mathcal{H} \ 0]$ with \mathcal{H} as for the standard EKF. Its linearization, \mathbf{H}_z is now a $M \times (I + P)$ matrix in which the last P columns contain zeros. The augmented state and covariance update complete the algorithm and are equivalent to those of the EKF except that they refer now to the augmented system, and the gain matrix has dimension $(I + P) \times M$ (see Carrassi and Vannitsem, 2011a, for details).

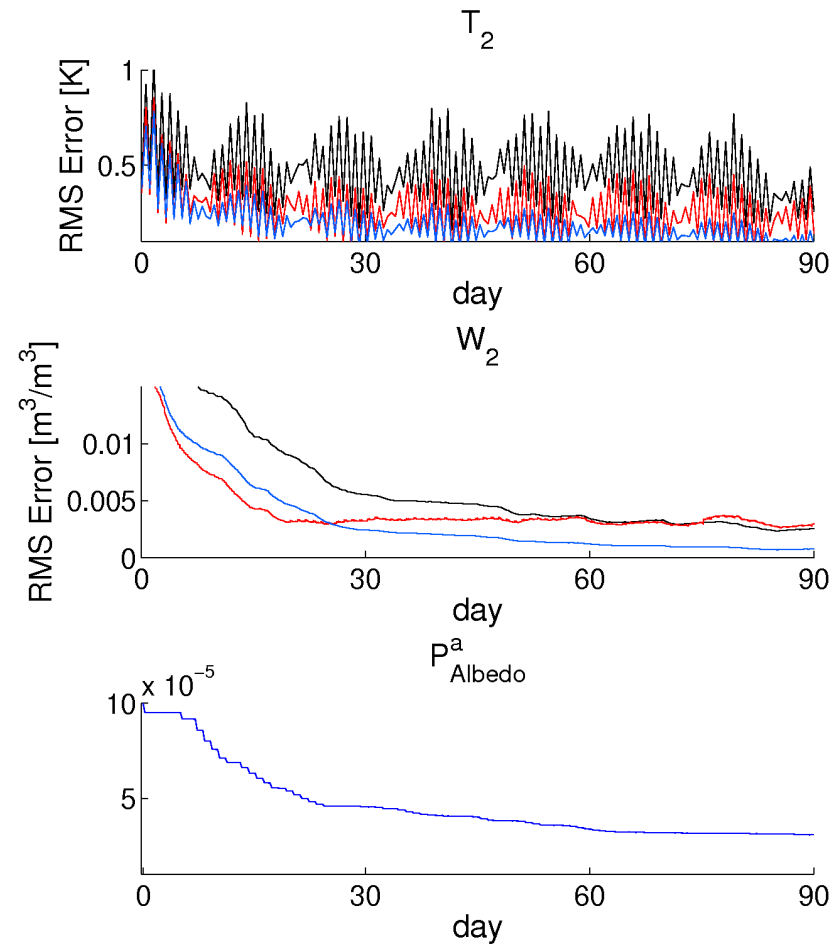
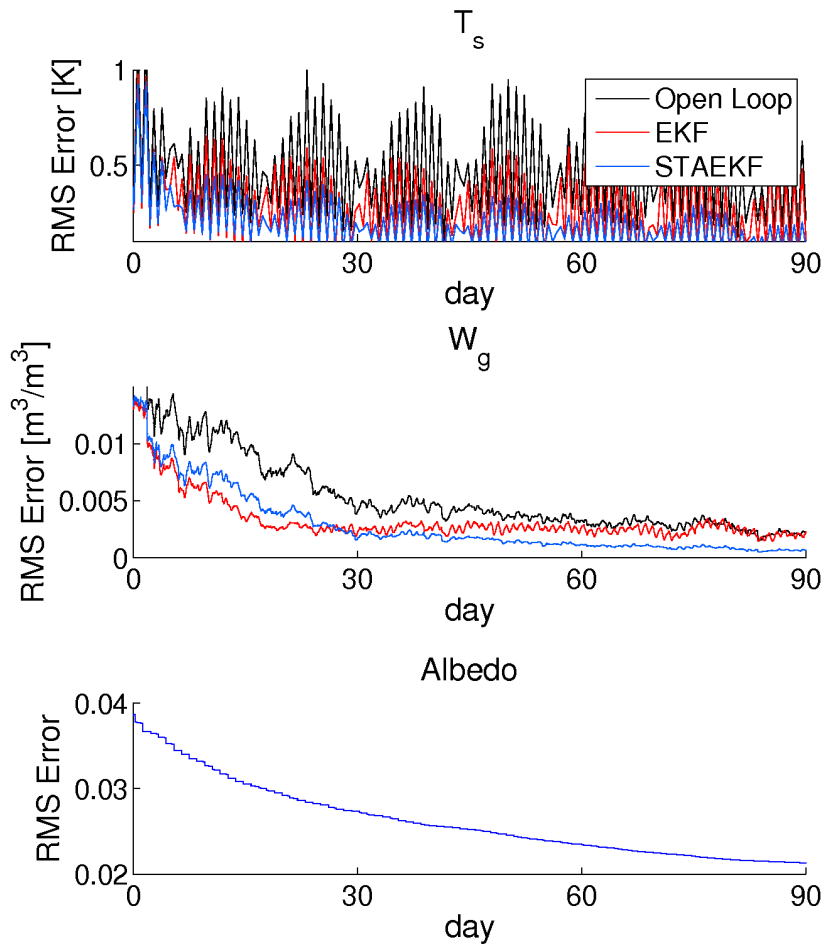
Experimental Set-up

- **Observation system simulation experiments (OSSE).**
- **The forcing consist of 1-hourly air temperature, specific humidity, atmospheric pressure, incoming global radiation, incoming long-wave radiation, precipitation rate and wind speed relative to the ten summers in the decade 1990-1999 extract from ECMWF Re-analysis ERA40.**
- **ISBA is run in one offline single column mode for a 90 day period.**
- **The simulated observations are T2m and RH2m at 00, 06, 12 and 18 UTC.**
- **The initial Pf (B) and Pm (Q) required by the EKF, are taken from Mahfouf (2007).
 $\text{diag}(R)=(1,10^{-2})$ $\text{diag}(Pf)=(1,1,10^{-2},10^{-2})$ $\text{diag}(Pm)=(25 \cdot 10^{-2}, 25 \cdot 10^{-2}, 4 \cdot 10^{-4}, 4 \cdot 10^{-4})$**
- **Parametric errors are introduced by perturbing either alternatively or simultaneously, the Leaf Area Index, LAI , the albedo, and the minimum stomatal resistance.**

Experimental Set-up

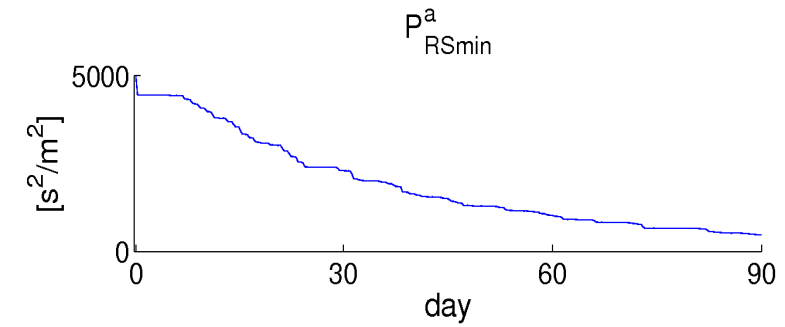
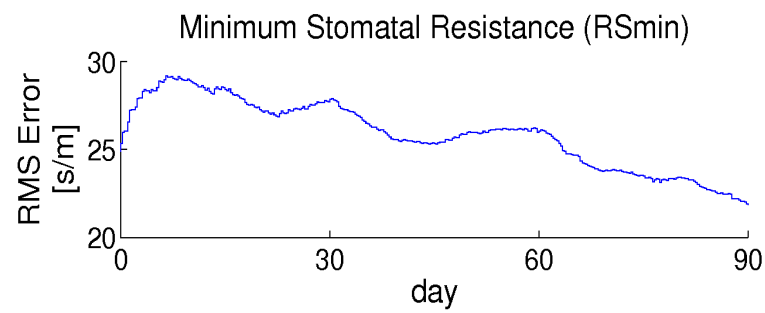
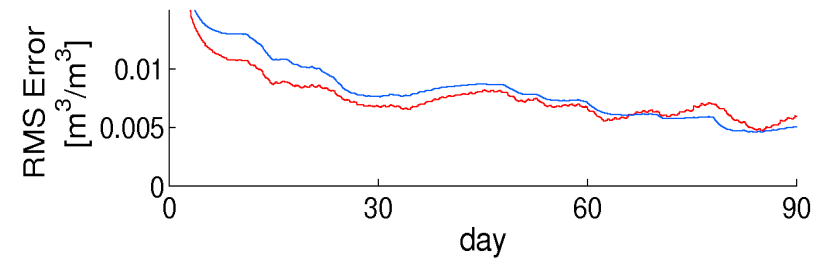
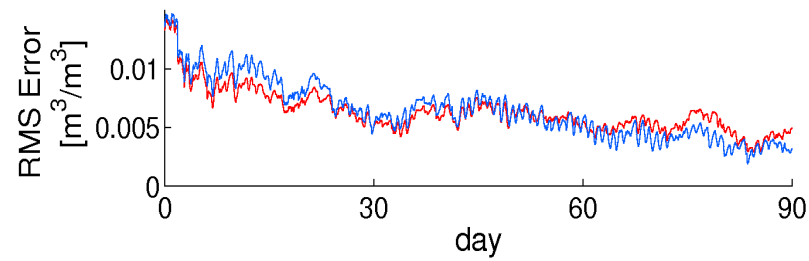
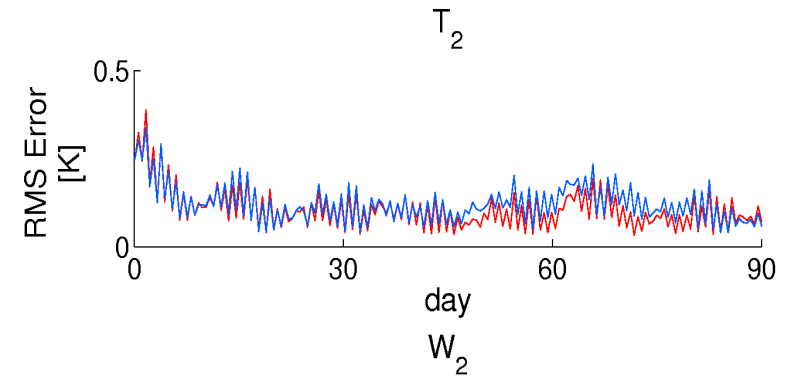
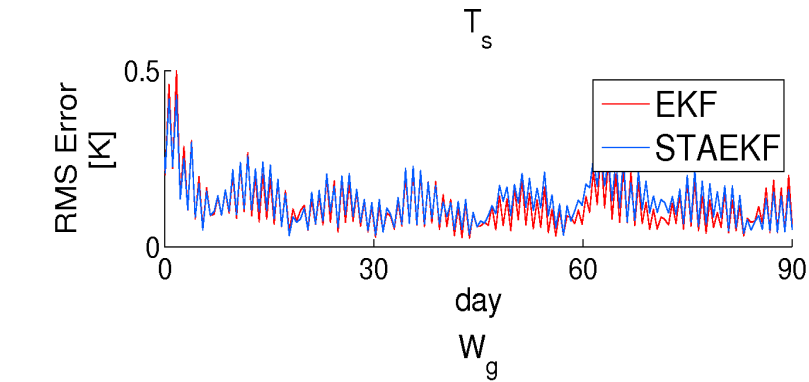
- For each summer in the period 1990 - 1999, a reference trajectory is generated by integrating the model with LAI = 1 m² /m², albedo = 0.2, and Rsmin=94 s/m.
- Around each of these trajectories, Gaussian samples of 100 initial conditions and uncertain parameters are used to initialize the assimilation cycles.
- The initial conditions are sampled from a distribution with standard deviation:
 $(\sigma T_s, \sigma T_2, \sigma w_g, \sigma w_2) = (5, 5, 1, 1)$
- LAI, albedo, and RSmin are sampled with standard deviations:
 $\sigma \text{LAI} = 0.5 \quad \sigma \text{albedo} = 0.1 \quad \sigma \text{RSmin} = 50$
- PLAI=1 Palbedo=10⁻⁴ PRSmin=5000 in the STAEKF, while Pax is taken as in the EKF; Px,λ is initially set to zero.

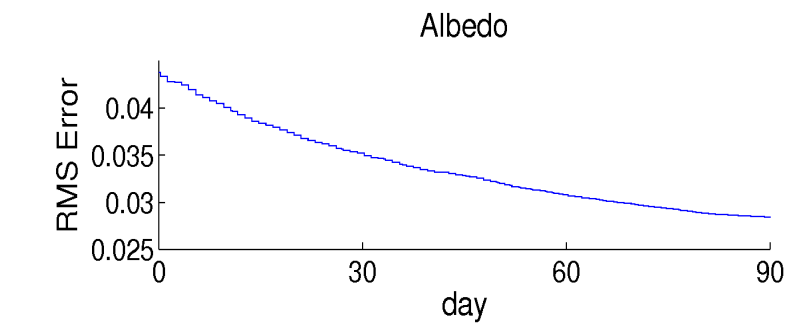
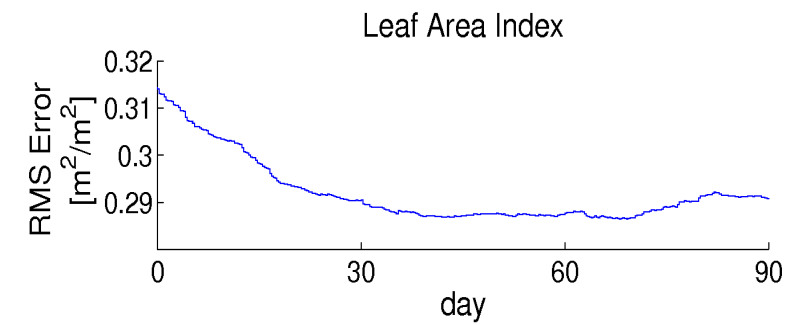
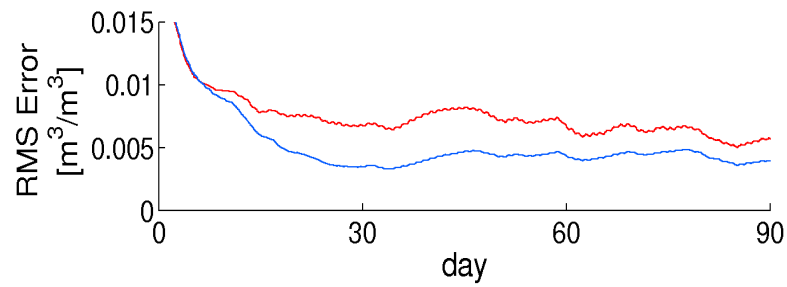
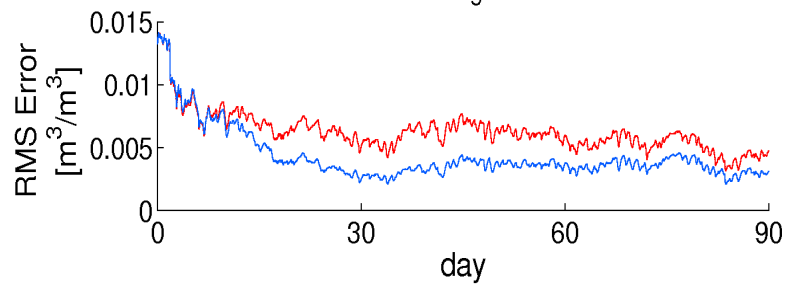
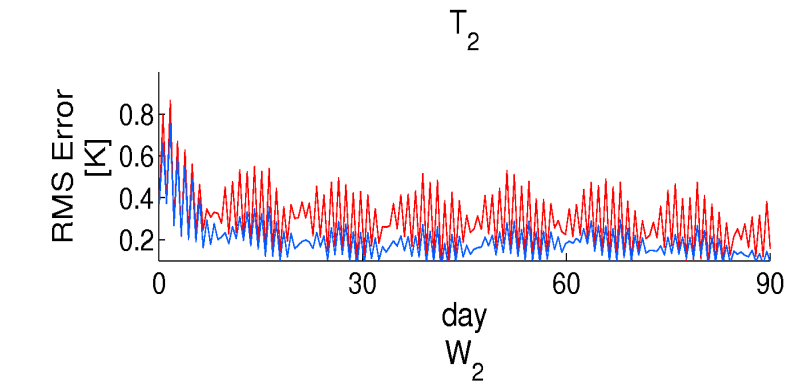
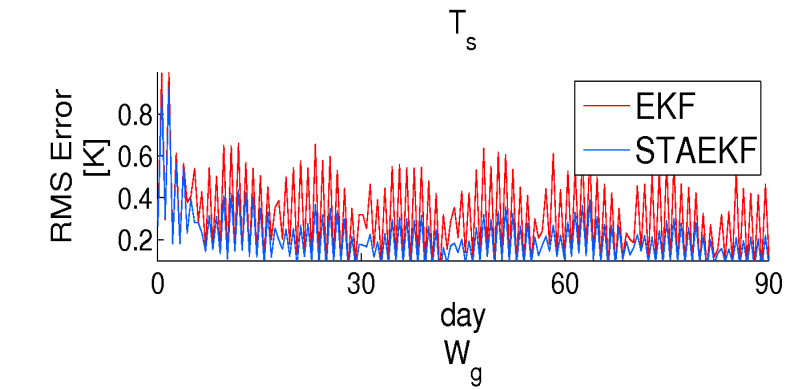




SURFEX
 SURFEX_EKF for ALARO
 SURFEX_EKF + 3DVAR
 STAEKF for soil analysis

LAI
 Albedo
 Rsmín
 LAI and Albedo





- **Continuing the evaluation of the STAEKF, a scientific paper is submitted: Short Time Augmented Extended Kalman Filter for Soil Analysis: A feasibility study.**
Carrassi A., Hamdi R., Termonia P., and Vannitsem S., 2012 ASL.
- **The STAEKF is able to reduce the parameter estimation errors.**
- **The accuracy of these estimates is inherently related to the type of parameter to be estimated.**
- **The model sensitivity to the specific parameter and the accuracy of the short-time approximation in the STAEKF.**
- **The rate of error convergence in the STAEKF is related to the initial parametric error variance.**
- **Implementing the STAEKF in SURFEX and study its behaviour within ALARO.**