

Unsaturated downdraft in Alaro-1

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Which downdraft

Knupp & Cotton 1985:

- Penetrative downdraft (non precipitating convection, width <1km, depth \sim 500m to 5km, w \sim 1-15 m/s)
- Cloud-edge downdraft (width < 5km, depth \sim 1-5 km, w<5m/s)
- Overshooting downdraft (cloud top, width \sim 500m to 5km, depth \sim 1 to 3km, w \sim 1-40 m/s))
- Precipitation-driven downdraft (Low level, width \sim 1 to 10 km, depth \sim 1-5 km , w<15 to 20 m/s).



From where and how much ?



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RMI

(GATE, time and space mean)

$$T_{r} = \frac{s}{c_{p}} = T + \frac{\phi - \phi_{\text{surf}}}{c_{p}}$$
$$T_{rd}^{*} = T_{wd} + \frac{\phi - \phi_{\text{surf}}}{c_{p}}$$
$$(assuming \ dry \ air)$$
$$T_{rs}^{*} = T_{w} + \frac{\phi - \phi_{\text{surf}}}{c_{p}}$$

- no dd in dry stable envt
- higher moisture below helps to produce negative buoyancy in dd

Subsaturation

Downdraft are usually appreciably subsaturated:

- 1. transport of condensate to be entrained and evaporated by the downdraft \rightarrow dependency to number/size of droplets or drops (especially in penetrative downdrafts)
- cooling rate (evap + melting) < adiabatic warming rate.
 Balance:
 - negative buoyancy increases downwards velocity
 - greater velocity
 - reduces residence time
 - increases the adiabatic warming rate



Precipitation-driven downdraft parametrization

Betts and Silva Dias 1979

 ψ_d follows a path of constant θ_e while remaining unsaturated.

$$\frac{dq_d}{dp} = \frac{q_w - q_d}{\Pi_E} + \frac{q_e - q_d}{\mathcal{L}_e}, \qquad \Pi_e = \frac{\rho g w_d}{4\pi DF}, \qquad \frac{1}{\mathcal{L}_e} = \frac{1}{M_d} \frac{dM_d}{dp}\Big|_e = \frac{\lambda_d d\phi}{dp}$$



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Diffusion coefficient $D(T_w, p) \approx 2.E - 5m^2 s^{-1}$

$$F = \int_{0}^{\infty} n(r)C_{v}(r)rdr, \qquad n(r) = n_{0}(2r)^{\mu}\exp(-2b\mathcal{P}^{-\beta}r)$$
$$C_{v}(r) = 1 + cr^{3/4} \qquad \text{(ventilation coefficient)}$$

 $\mathcal{P}=$ rainfall rate, r=droplet radius, μ (conv|strat) taken ~ 0

$$F(\mathcal{P}) \approx a_1 \mathcal{P}^{e_1} + a_2 \mathcal{P}^{e_2}, \qquad 2\Pi_E_{[l-1\to l]} \approx \frac{\alpha}{\mathcal{P}^{3/4} \triangle p} \widetilde{\omega_d} \triangle p = k^{\overline{l-1}} \widetilde{\omega}^{\overline{l-1}} \triangle p^{\overline{l-1}}$$









$$\frac{1}{T_{vd}} \approx \frac{1}{T_d(1+\nu q_d)} \approx \frac{(c'\,\widetilde{\omega_d}+1)}{(a'\,\widetilde{\omega_d}+b')} \cdot \Big\{ \frac{(c'\,\widetilde{\omega_{d0}}+1)}{[(c'+\nu m)\,\widetilde{\omega_{d0}}+(1+\nu n)]} \Big\}$$



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Vertical velocity equation

$$\frac{\partial \omega_d^{\diamond}}{\partial t} = -D\omega_d^{\diamond 2} - \omega_d^{\diamond} \Big(\frac{\partial \omega_d^{\diamond}}{\partial p} + \frac{\partial \overline{\omega}}{\partial p}\Big) + \frac{g^2 p}{2R_a} \Big[\frac{(c\omega_d^{\diamond} + d)}{(a\omega_d^{\diamond} + b)} - \frac{1}{\overline{T_v}}\Big]$$

D: drag, \ni entrainment, friction, surface ∇p



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$$\alpha(F_{n+1}^{\overline{l-1}})^3 + \beta(F_{n+1}^{\overline{l-1}})^2 + \gamma(F_{n+1}^{\overline{l-1}}) + \delta = 0, \qquad \qquad F_n^{\overline{l}} = \frac{\omega_d^{\diamond l} + \omega_d^{\diamond l+1}}{2} \text{ at time n}$$

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$$D^{l} = \frac{1}{2(1 - \sigma_{d}^{l})^{2}} \Big[\frac{R_{a}T_{vd}^{l}}{p^{l}} \big(\underbrace{\lambda_{d}^{l}}_{\mathsf{TENTRD}} + \frac{1}{g} \underbrace{\mathcal{K}_{dd}}_{\mathsf{TDDFR}} \big) + \frac{\mathsf{GDDDP}}{(p^{\overline{L}} - p^{l})^{\mathsf{GDDBETA}}} \Big]$$



Descent computation

• Saturated entraining moist adiabat: mixed $\psi_b \rightarrow \psi_n$

$$T_{b}^{l-1} = \psi_{n}^{l-1} + \xi^{\overline{l-1}} (\overline{T_{w}}^{l-1} - T_{n}^{l-1})$$

• parallel computation of

$$\alpha \widetilde{\omega_d}^3 + \beta \widetilde{\omega_d}^2 + \gamma \widetilde{\omega_d} + \delta = 0, \qquad q_d^l = \frac{m \widetilde{\omega_d} + n}{c' \widetilde{\omega_d} + 1}, \qquad T_d^l = \frac{a' \widetilde{\omega_d} + b'}{c' \widetilde{\omega_d} + 1}$$

- Control arrival level:
- not saturated
- remaining precipitation
- $k\widetilde{\omega_d} > 1.E 12$

$$\delta q_{ev}^{\overline{l-1}} = \frac{q_w - q_d}{\Pi_e} \triangle p^{\overline{l-1}}$$



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- M related to precipitation rate $R_i(N_0, M, z) = \frac{\overline{R}I_d}{f_c}$ (Kessler), $I_d = \text{rain intensity distribution function under a convective cloud cover } f_c$, $\overline{R} = \text{average convective rainfall in the grid cell.}$



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α_i	0.52	0.34	0.09	0.04	0.01
I_d	0	0.40	2.6	11.25	18
%	0	23.6	23.4	45	18

(Ruprecht & Gray 1976)

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maybe $\sigma_{\mathcal{P}}$ presently quite crudely estimated



1-D model profiles





CAPE closure: express a relaxation of CAPE under the effect of convective activity...

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 $\mathsf{CAPE}\xspace$ is affected by

- resolved and subgrid convergence
- resolved and subgrid precipitation
- radiative transfers
- deep convective updraft
- boundary layer transfers
- downdraft
- ...



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All these external processes (if represented by the model) intervene from one time step to the next anyway.

 \Rightarrow it could be seen as a *subjective choice* to include or not the effect of downdraft on CAPE in the updraft parameterization itself.







Main effect of downdraft is cooling and moistening USL



• Change LCL temperature and pressure level





- Change LCL temperature and pressure level
- Change LFC \Rightarrow increased CIN
- Decreased CAPE









 M_d^* advected from previous time step





 M_d^\ast advected from previous time step

$$\frac{\partial \overline{T}^{\text{USL}}}{\partial t}\Big|_{dd} \approx -\frac{\sum_{uslb}^{uslt} M_d^* \frac{1}{c_p} \frac{\partial \overline{s}}{\partial p} \bigtriangleup p}{\bigtriangleup p^{mix}}, \qquad \frac{\partial \overline{q}^{\text{USL}}}{\partial t}\Big|_{dd} \approx -\frac{\sum_{uslb}^{uslt} M_d^* \frac{\partial \overline{q}}{\partial p} \bigtriangleup p}{\bigtriangleup p^{mix}}, \qquad \frac{\partial p^{USL}}{\partial t}\Big|_{dd} \sim 0$$

Bolton (1980)

$$T^{\text{LCL}} = \frac{2840}{3.5 \ln T^{\text{USL}} - \ln \frac{e^{\text{USL}}}{100} - 4.805}} + 55., \qquad e^{\text{USL}} = \frac{q^{\text{USL}}}{q^{\text{USL}} + \frac{R_a}{R_v}(1 - q^{\text{USL}})} p^{\text{USL}}$$
$$\Rightarrow \frac{\partial T^{\text{LCL}}}{\partial t} = -\frac{(T^{\text{LCL}} - 55)^2}{2840} \Big[\frac{3.5}{T^{\text{USL}}} \frac{\partial T^{\text{USL}}}{\partial t} - \frac{1}{q^{\text{USL}}} \frac{\partial q^{\text{USL}}}{\partial t} \frac{R_a}{R_a + q^{\text{USL}}(R_v - R_a)} \Big] \quad < 0,$$



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Current status: plug and play

- LNSDO allows to choose the unsaturated downdraft (acnsdo instead of acmodo), both 3MT (ω_d) and Alaro-1 (ω_d^{\diamond}) versions available.
- Tuning parameters:
 - gddevf : σ_d/σ_P (0.33)
 - tentrd : entrainment rate (1.E-4), tddfr : drag (6.E-4, def)
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 - Icddevpro=F, jddevpro=1 (in acupm: update input profile with total evaporation)
- Profiles are realistic, further tuning should be based on model scores.
- Possible re-tuning of hard-coded constants.
- Further refinements: better estimation of $\sigma_{\mathcal{P}}$

